

Dynamics of near-surface wind response to Arctic sea ice

1. Summary

Goal: Examine the dynamics of two near-surface wind responses, 10-meter wind (W10) and SLP-based geostrophic wind (Wg), to various Arctic sea ice conditions.

Method: Use the Polar WRF weather model forced with multiple sea ice concentration (SIC) datasets.

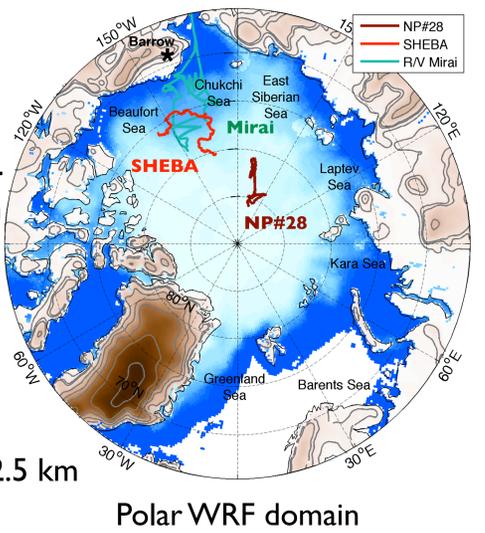
Result: Variations in W10 and Wg involve distinctive dynamical mechanisms and reflect different spatial information of sea ice condition.

Implication: An accurate representation of the dynamical process for surface winds is important for the improved understanding of air-ice-ocean interactions.

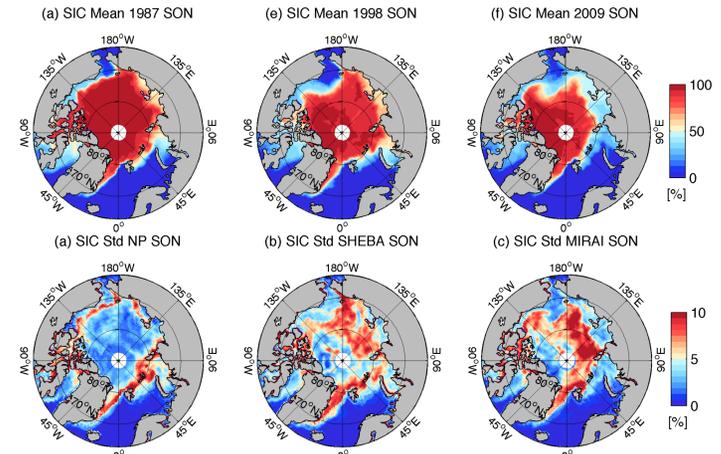
SST-wind mechanisms	vertical mixing	pressure adjustment
key process	ID turbulent momentum transport	lateral pressure gradient and ageostrophic flows
phase relationship	$\nabla \cdot \mathbf{u} \propto \nabla_d T$	$\nabla \cdot \mathbf{u} \propto \nabla^2 P \propto \nabla^2 T$
time-scale	faster (<synoptic)	slower (>synoptic)
height-scale	shallower (below PBL)	deeper (beyond PBL)
horizontal-scale	broader (the whole Arctic basin)	narrower (the ice margins)

2. Model, experiment, and data

- Model: Polar WRF (Hines and Bromwich, 2008), 25 km resolution
- Experiment: A series of 1-year integration forced with 3 SICs
 - NT: NASA Team, 25 km (Cavaliere et al., 1996)
 - BT: Bootstrap, 25 km (Comiso, 2000)
 - EU: EUMETSAT hybrid, 12.5 km (Tonboe et al., 2011)

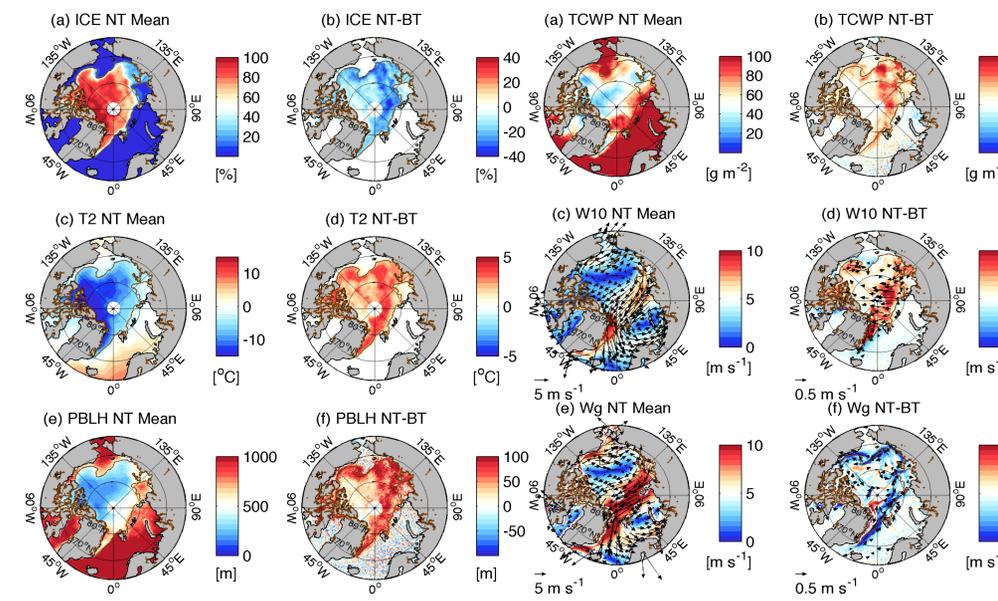


3. Sea ice concentration estimates



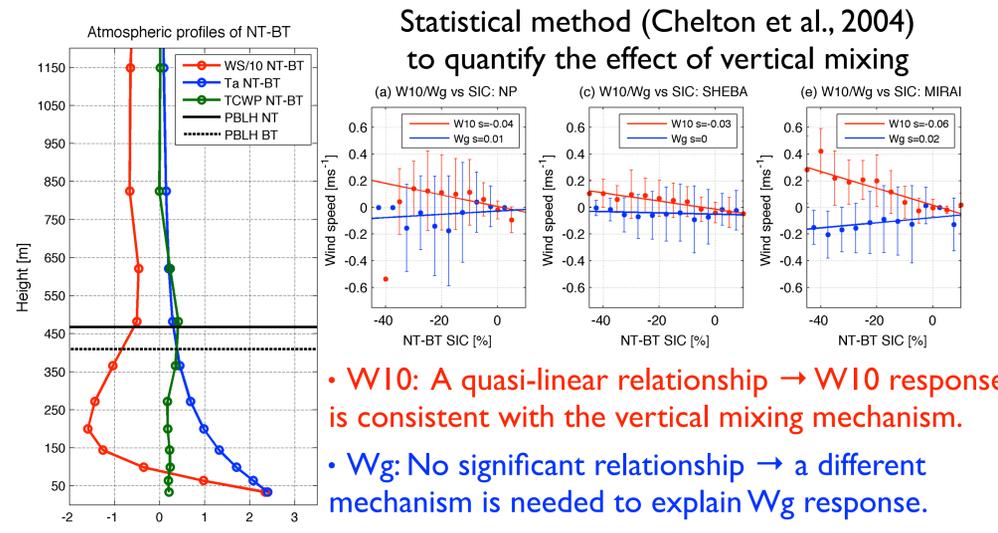
- The autumn SICs show an Arctic-wide uncertainty pattern.
- We will look for ABL response to different SIC products.

4. Pan-arctic ABL response: NT-BT



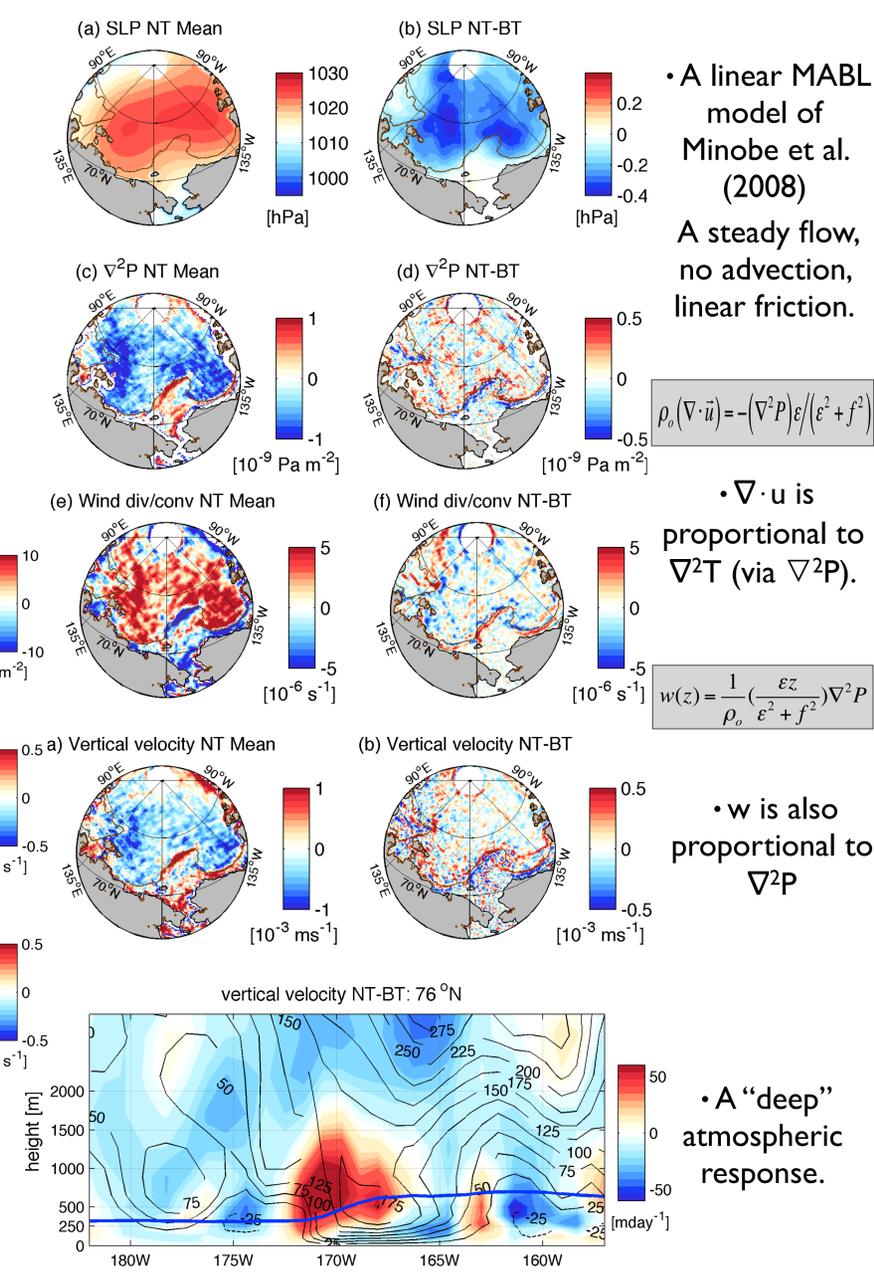
- Lower SIC produces higher T2, PBL and cloud water path, and W10, an indication of an unstable ABL.
- Comparable spatial scale of between wind and SIC
- $T' \rightarrow u'$ via ABL stability

5. Modulation of the ABL static stability



- W10: A quasi-linear relationship \rightarrow W10 response is consistent with the vertical mixing mechanism.
- Wg: No significant relationship \rightarrow a different mechanism is needed to explain Wg response.
- Unstable PBL with higher temperature and clouds below PBL.
- Increased surface wind at the expense of decreased wind aloft

6. Modulation of surface pressure



- A linear MABL model of Minobe et al. (2008)
- A steady flow, no advection, linear friction.
- $\rho_o (\nabla \cdot \mathbf{u}) = -(\nabla^2 P) \epsilon / (\epsilon^2 + f^2)$
- $\nabla \cdot \mathbf{u}$ is proportional to $\nabla^2 T$ (via $\nabla^2 P$).
- $w(z) = \frac{1}{\rho_o} \left(\frac{\epsilon z}{\epsilon^2 + f^2} \right) \nabla^2 P$
- w is also proportional to $\nabla^2 P$
- A "deep" atmospheric response.

7. Implications

- W10 and Wg, governed by different dynamics, reflect different spatial information of ice condition.
- In situ SLP-based Wg underestimates the effect of basin-scale SIC changes.
- W10 from coarse resolution reanalyses underestimates the wind variations across the ice margins.
- A more accurate representation of the surface wind variability reflecting the both effects is needed.

Seo, H. and J. Yang, 2013: Dynamical response of the Arctic atmospheric boundary layer process to uncertainties in sea ice concentration. *JGR-Atmos.*, 118, 12,383-12,402