

Drivers of Arctic sea ice decline in the HadGEM1 model

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1) Introduction and motivation 2) Heat budget: mean seasonal cycle · The Arctic sea ice extent and volume have been (RCP4.5 Dominant declining for a number of years. • Models submitted to the CMIP3 and CMIP5 components: intercomparisons show a wide range of rates of meltina Summer melting: decline, with a number projecting a slower decline than observed Melting at the top surface of the ice There is a large spread of estimates of the date at which the Arctic will first become seasonally ice-free during June and July to ocean to ice heat 3.0×1 flux, extending into Autumn Projections of September ice extent from 2.5×1 AR5 models (Streove et al. 2012) freezing · Global coupled models remain our best tool for predicting Winter freezing future climate change, hence it is important to understand 1.5×1 Diffusive heat flux WHY model projections differ through ice to the Here we use the heat budget of the sea ice as a tool to understand the factors driving the decline in one (CMIP3) atmosphere Components of the heat budget of the sea ice and Frazil ice formatior model: HadGEM1 overlying snow, for the HadGEM1 control · Offset by some basal · We use an ensemble of projections, forced with observed integration, expressed in terms of an equivalent melting time varying anthropogenic forcing to 2000, and the A1B or amount of heat entering or leaving the ice (per unit A2 scenario thereafter area of ice) HadGEM1 - projections of ice volume 3) Changes in components of the heat budget 4) A robust signal? ies w.r.t. LTM contro The decline in ice volume is due to extra melting during the summer, rather than reduced freezing during the winter: 2010-2019 (anomalies w.r.t. control) Ocean to ice total heat flux anomalies 9 11 (August) More melting at the sides botmel More surface melting and base of the ice during during June and (initially) August and (later) Julv 4×10 In HadGEM1 the ice 2000 1970 1980 199 201 It is likely that the extra albedo is a function of 2×10 Topmelt: June anomalies w.r.t. LTM control heating is due to in-situ surface temperature. 8×10 warming of the ocean A1B surface as the ice cover Warmer surface temperatures in June lead the extra melting is to reduced albedo and -2×10 confined to the Surface melting 4×10 extra melting 4 6 8 10 12 summer months anomalies spatially the (June) This effect is less evident increases in o2i heat 2040-2049 (anomalies w.r.t. control) in July, as in the control flux are correlated with integration the surface 1.5× decreases in ice temperature is already concentration close to the melting -2×10 temperature 1.0×1 Analysis of the heat budget of the upper ocean Other changes (eg cloud) is required to confirm this Evolution of anomalies, compared with control variability under investigation (+/- 1 and 2 standard deviations) 5.0×10 5) Summary and future work 2010-2019 response -5.0×10 Initial analysis of changes in the components of the sea ice heat budget 10 12 shows the main drivers of ice decline in the HadGEM1 model Changes in components of the heat budget of the sea ice and overlying snow, Analysis will continue to confirm conclusions and link the flux changes to for the HadGEM1, expressed in terms of an equivalent amount of heat entering wider model processes or leaving the ice (per unit area of ice) · Should be a useful approach for inter-comparing the factors driving the decline in sea ice in a range of models.

In both cases, <u>the ice albedo feedback</u> is important in determining which months have extra melting – highlighting the importance of a physically realistic representation of albedo in models used to make projections of future climate change.

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 \bullet Combined with other work to compare model budgets with observations (West et al, in preparation), can aid the assessment of model strengths and

weaknesses