

Coupled modeling of eddy-wind interaction
in the California Current System
— Impact on eddy kinetic energy and Ekman pumping

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Eddy-wind interaction

$$\tau = \rho C_D (U_a - U_o) |U_a - U_o|$$

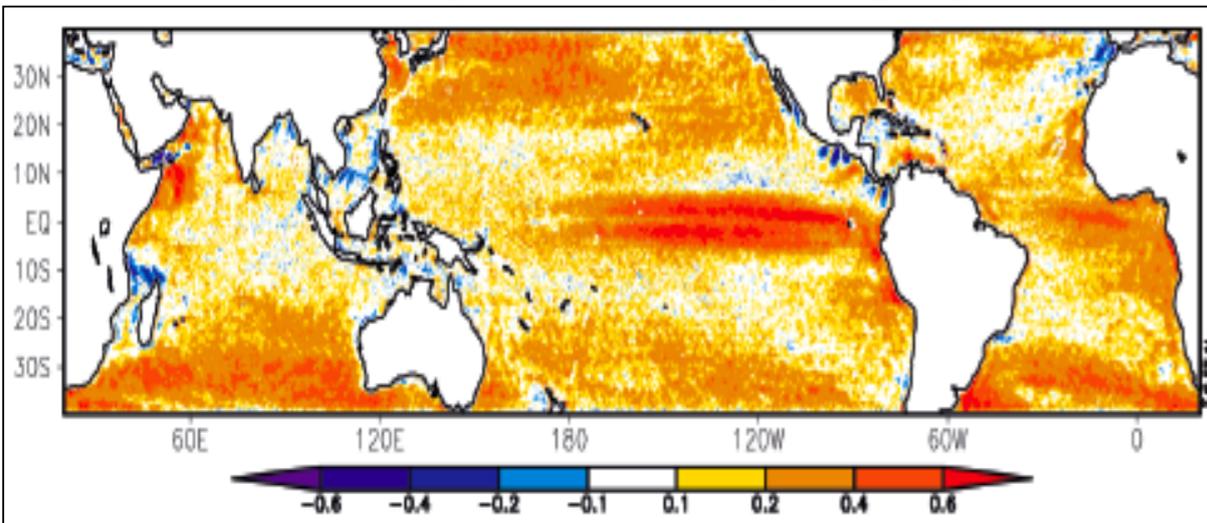
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$$U_a = U_{ab} + \underline{U_{aSST}}$$

Satellite observations: Xie 2004

Correlation (SST & wind): high-passed

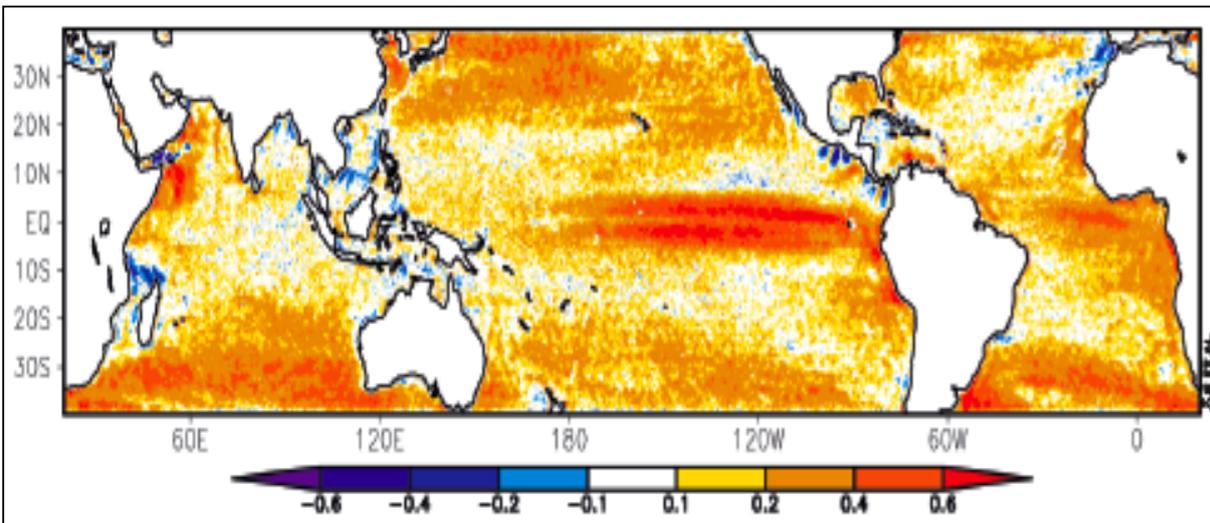


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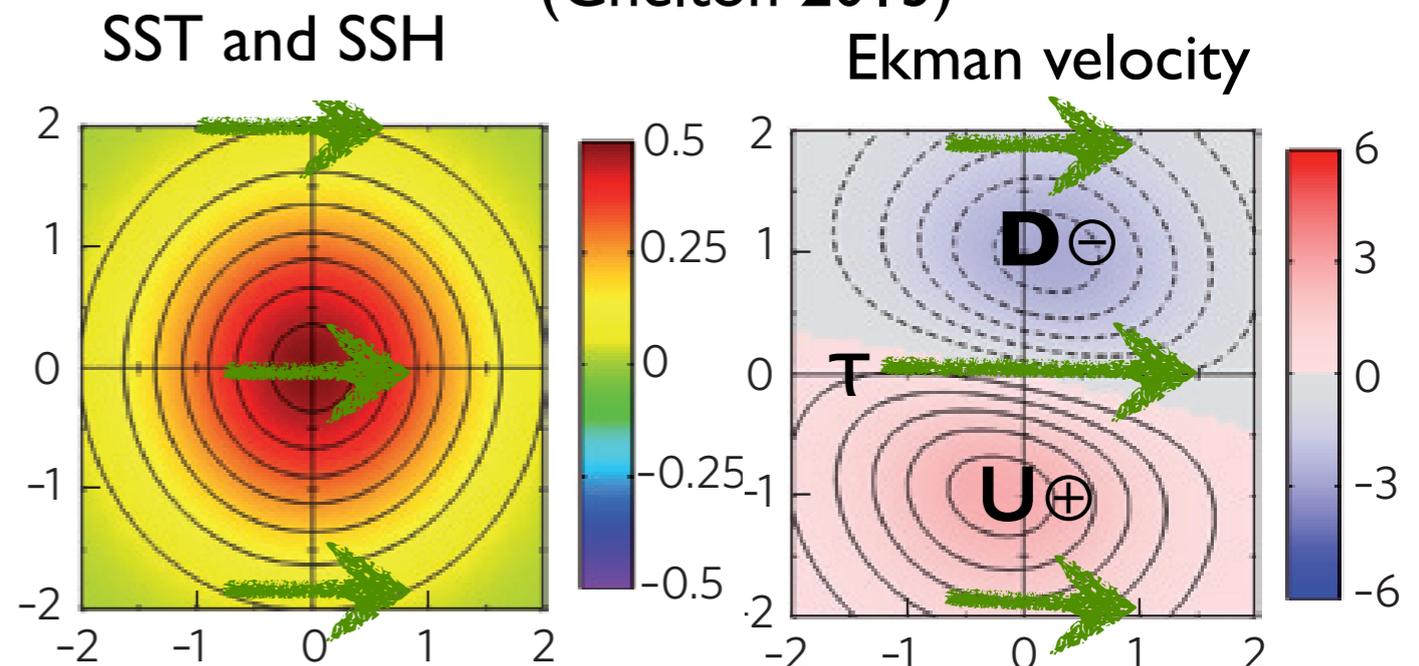
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An anticyclonic eddy in the Southern Ocean
(Chelton 2013)



northward propagation
of an anticyclonic eddy

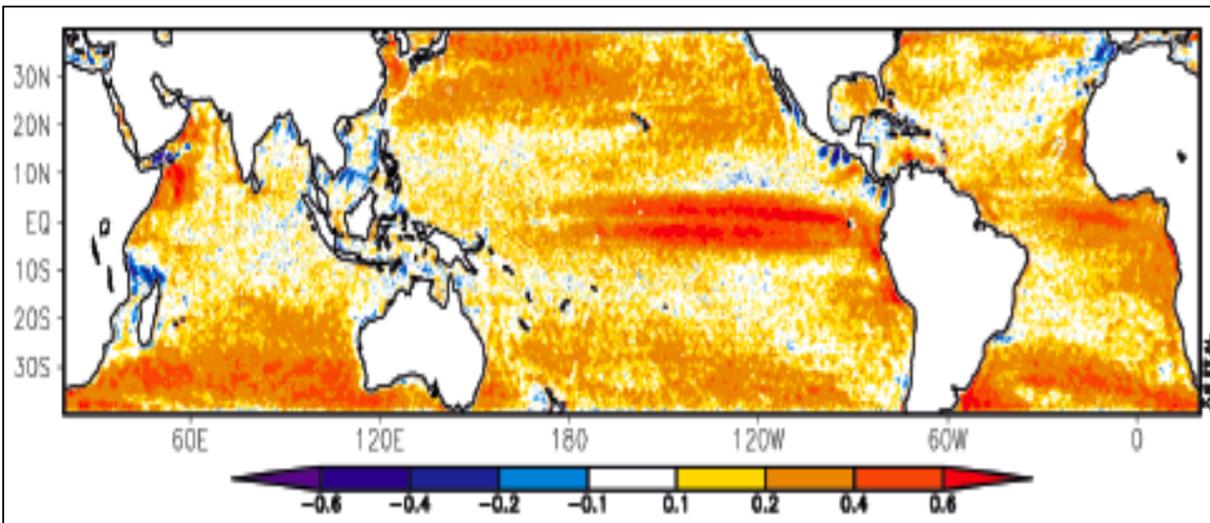
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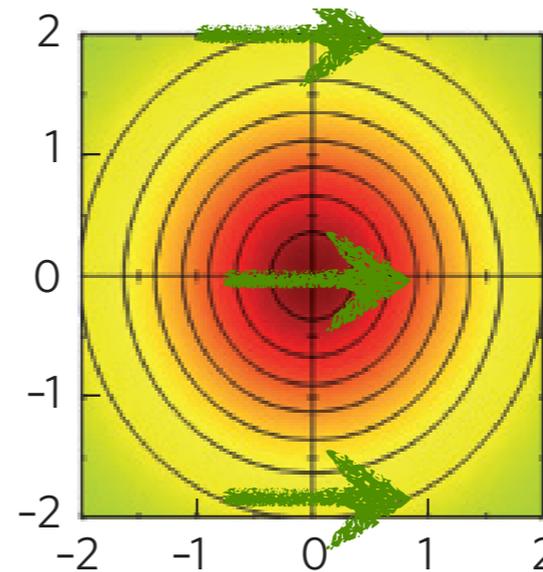
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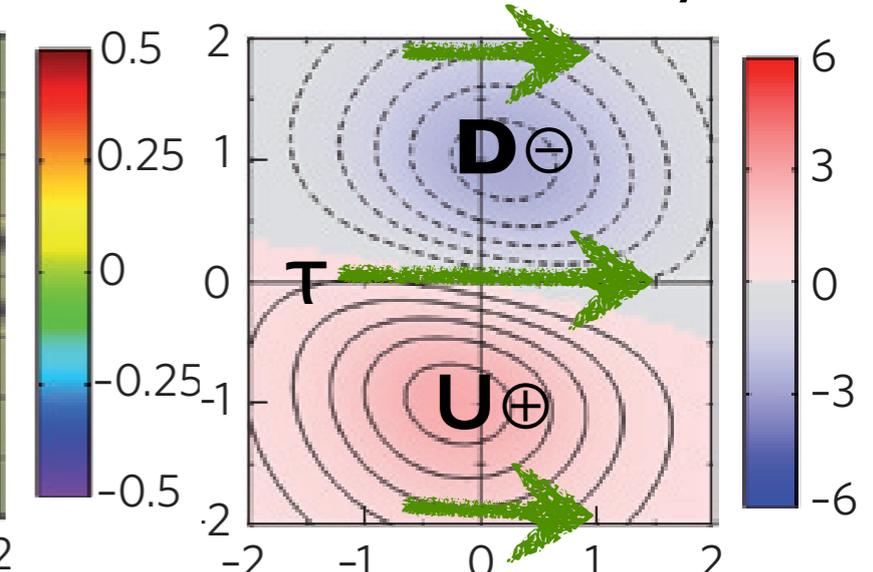


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SST and SSH



Ekman velocity



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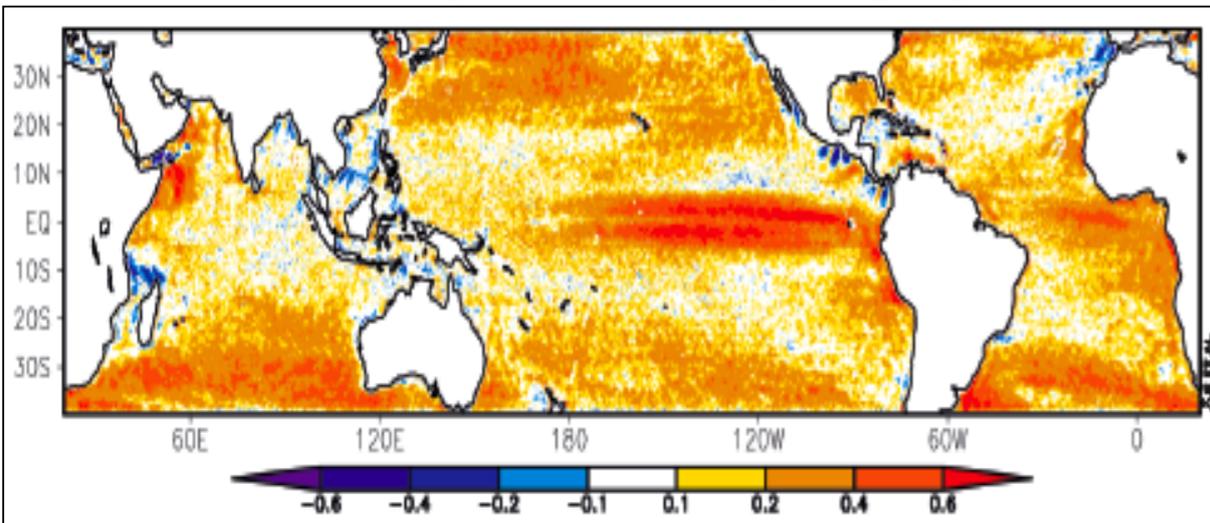
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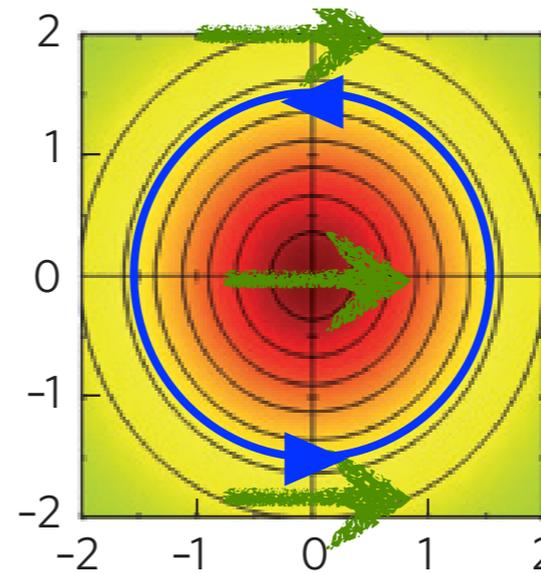
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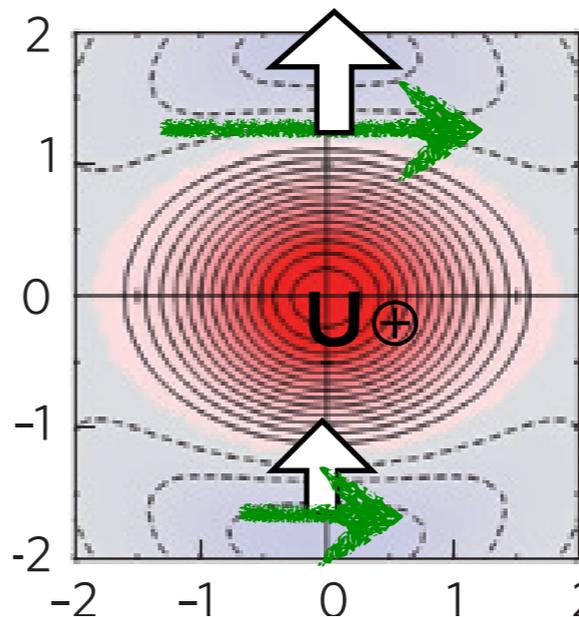
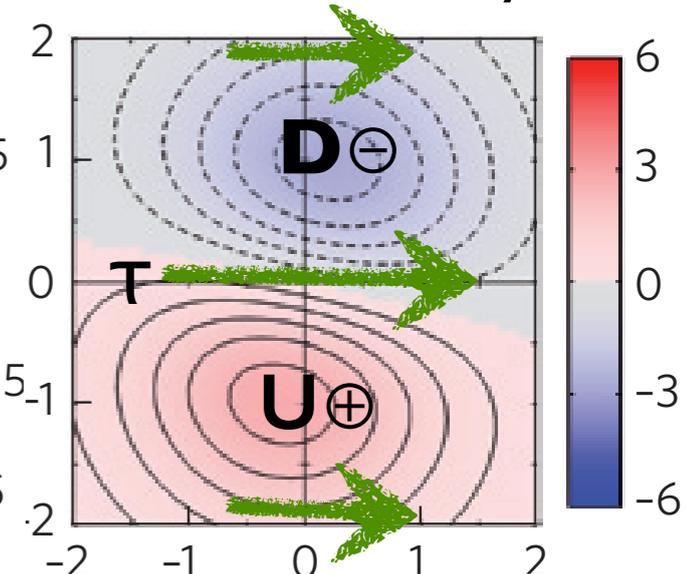


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SST and SSH



Ekman velocity



northward propagation
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Upwelling at the center:
decaying of an
anticyclonic eddy

$$W_e = \tau / [\rho(f + \zeta)]$$

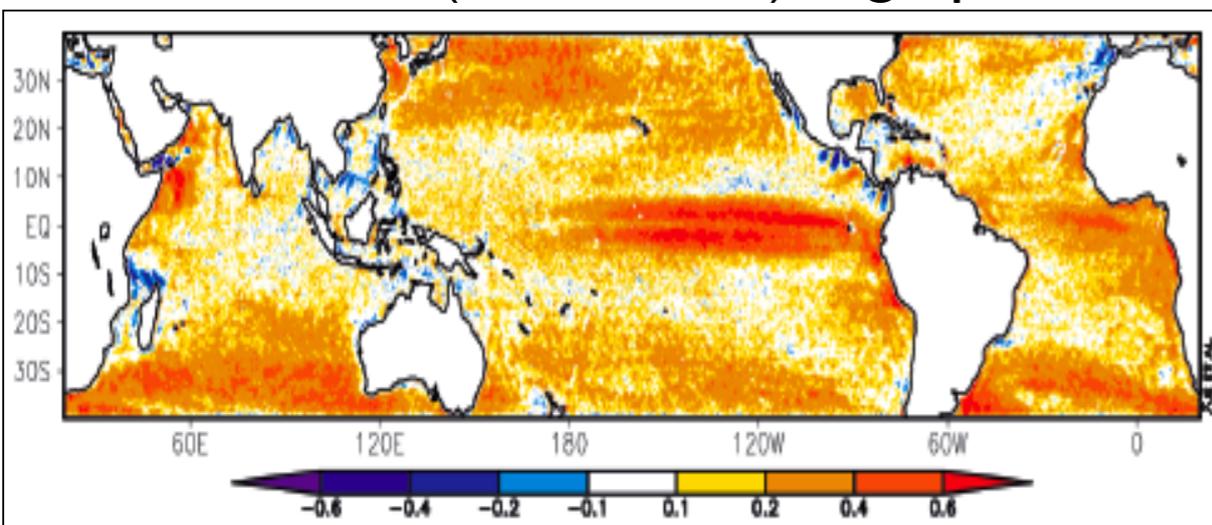
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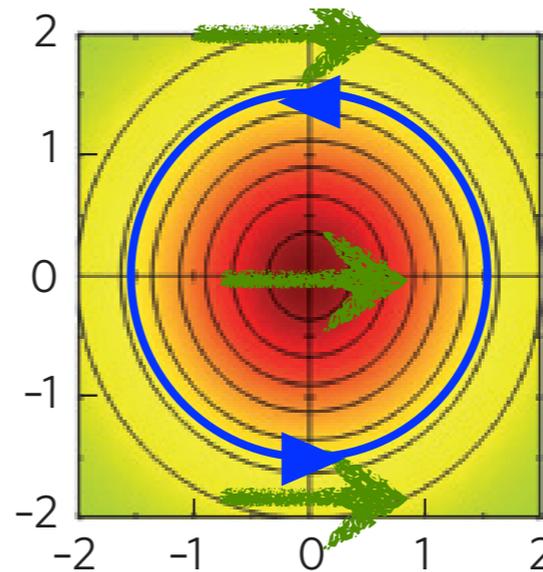
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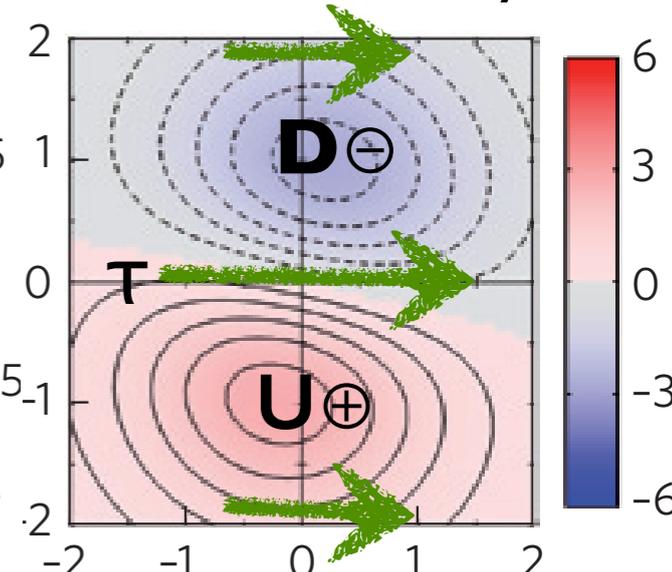


An anticyclonic eddy in the Southern Ocean
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SST and SSH

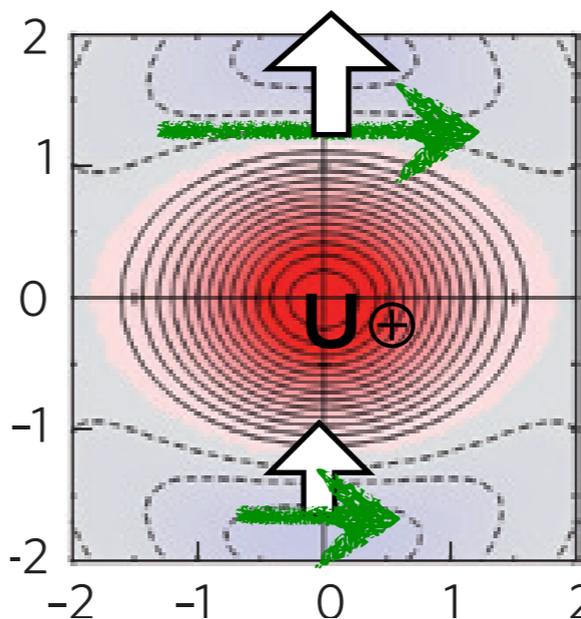


Ekman velocity



Different feedbacks due to SST- and current-induced eddy-wind interactions!

Key question:
relative impact of τ_{SST} and τ_{cur}
on the eddy dynamics?



northward propagation
of an anticyclonic eddy

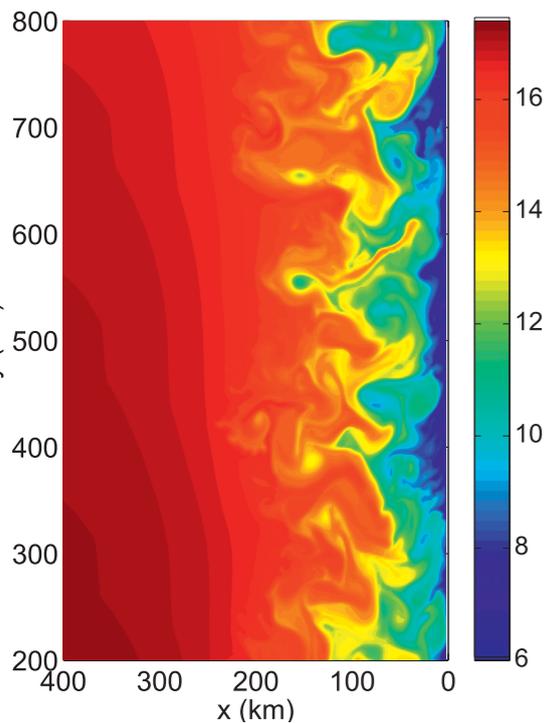
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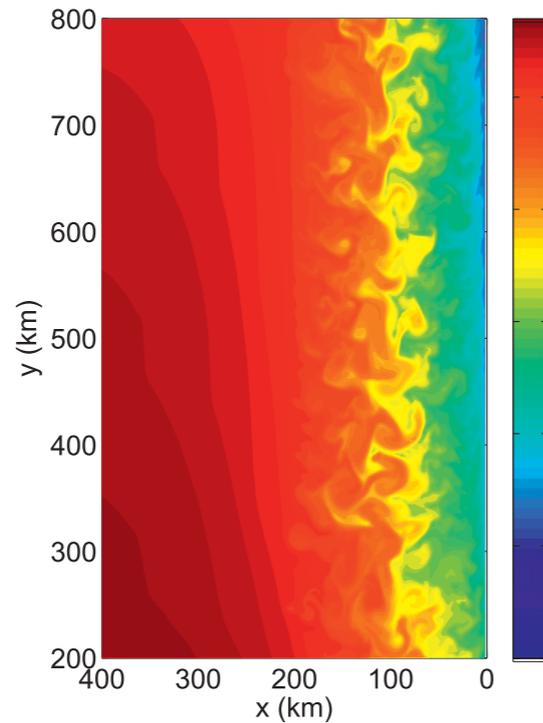
Previous studies on impacts of τ_{SST} and τ_{cur}

- Previous ocean-modeling studies show weakened eddy variability with inclusion of τ_{SST} and τ_{cur} .

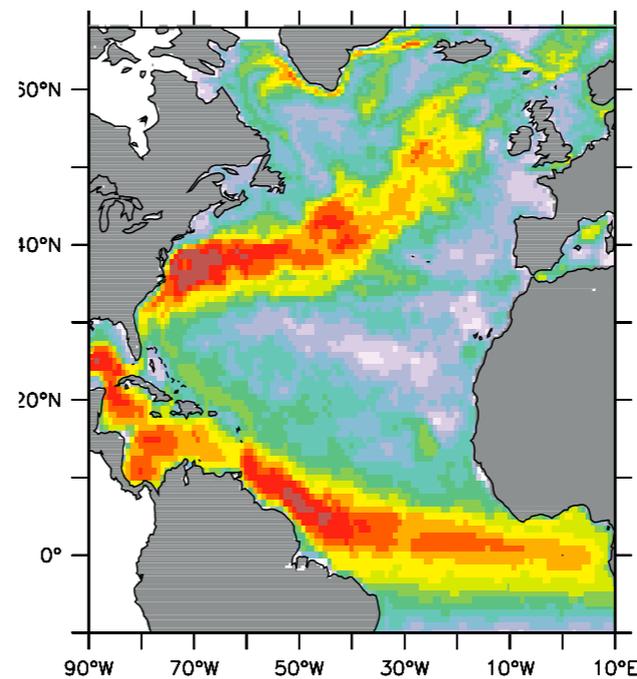
uncoupled SST



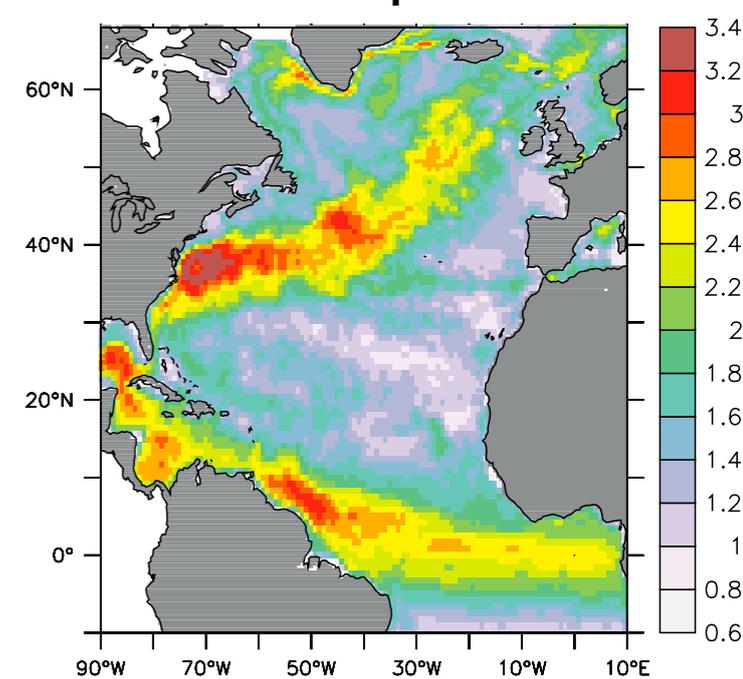
SST- τ coupled SST



uncoupled EKE



U_o - τ coupled EKE

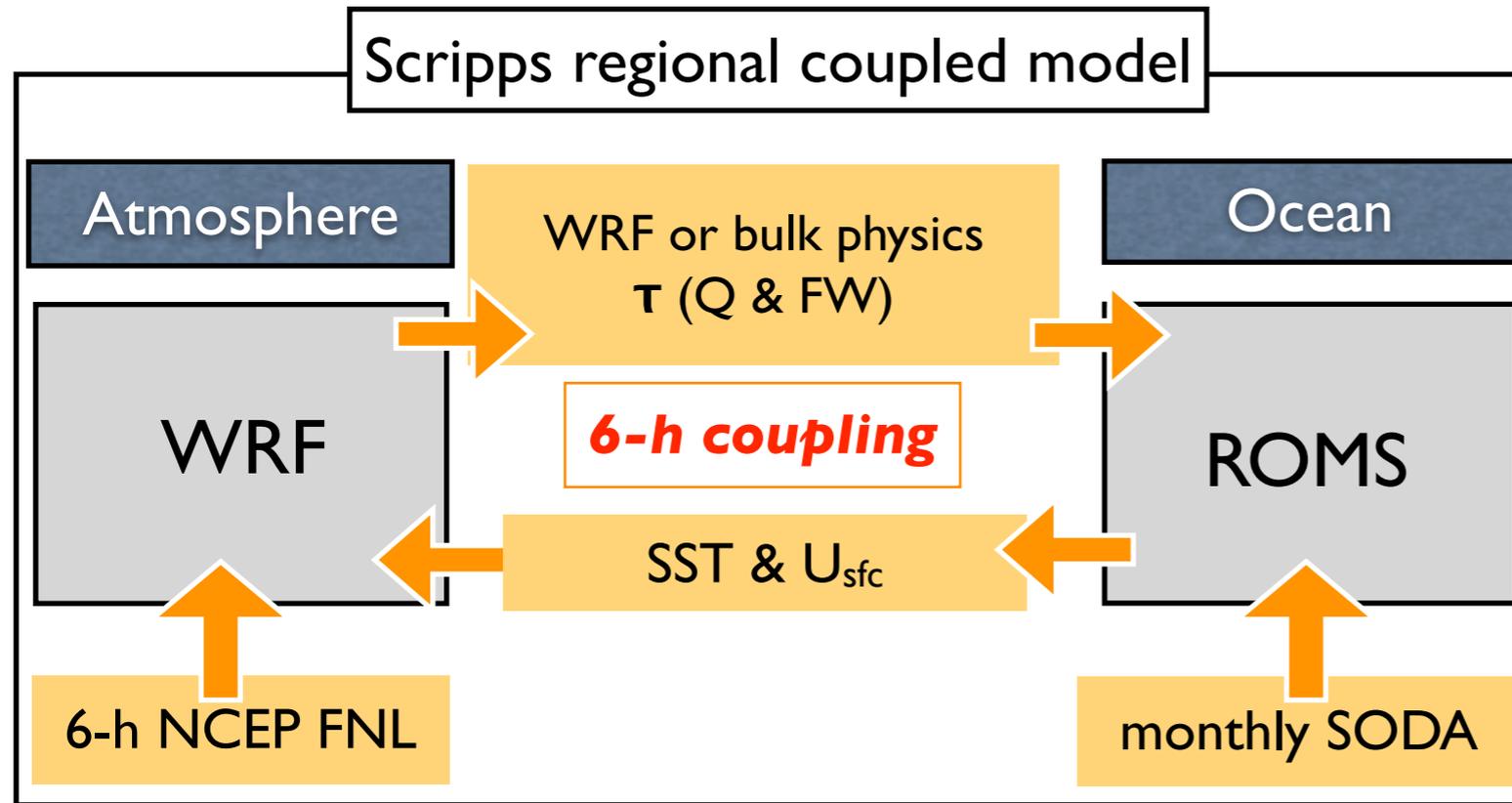


effect of τ_{SST} : *Jin et al. (2009)*

effect of τ_{cur} : *Eden and Dietze (2009)*

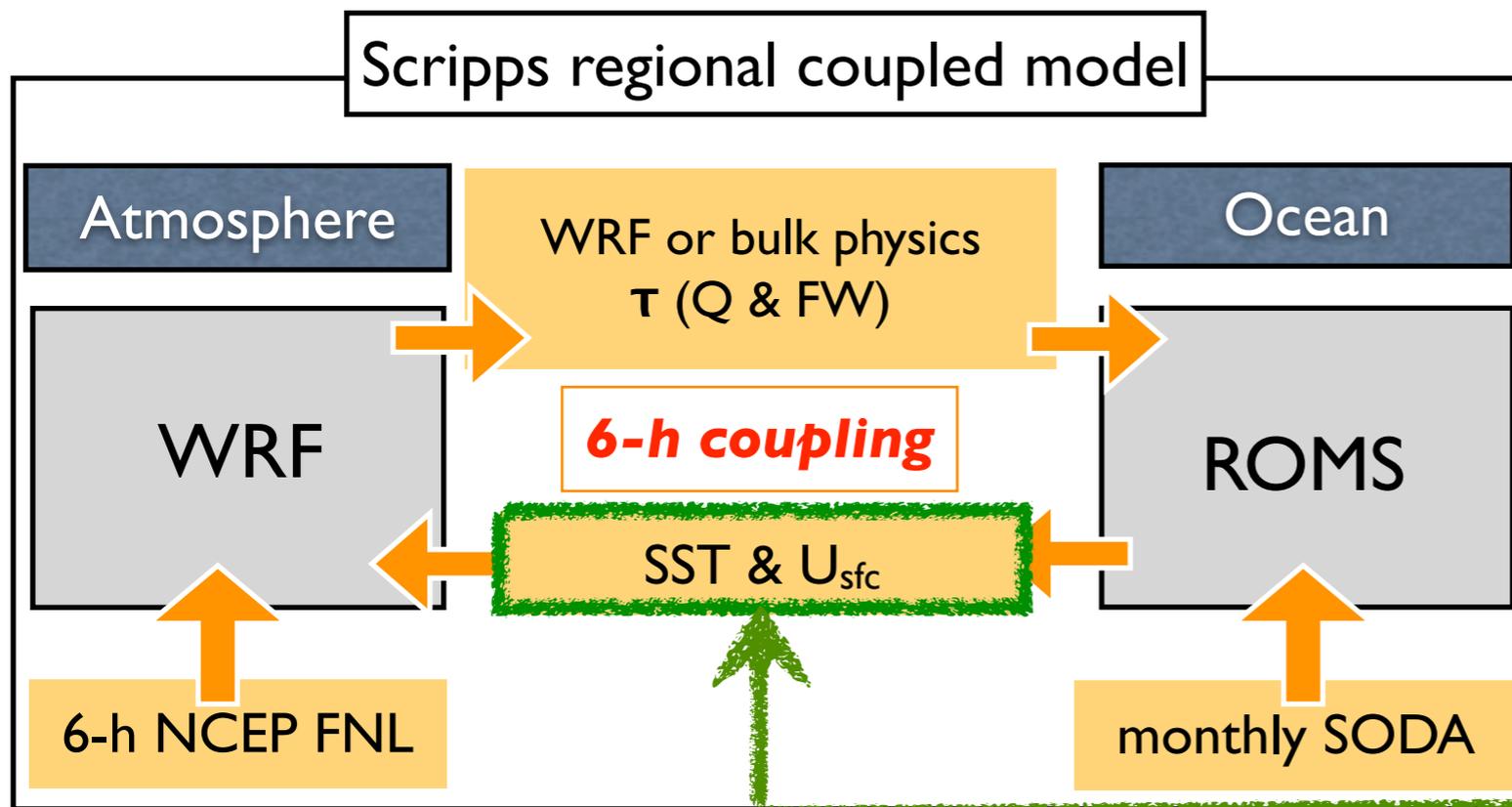
- This study examines the relative importance of τ_{SST} vs τ_{cur} using a fully coupled regional model.

Scripps regional coupled model and experiments



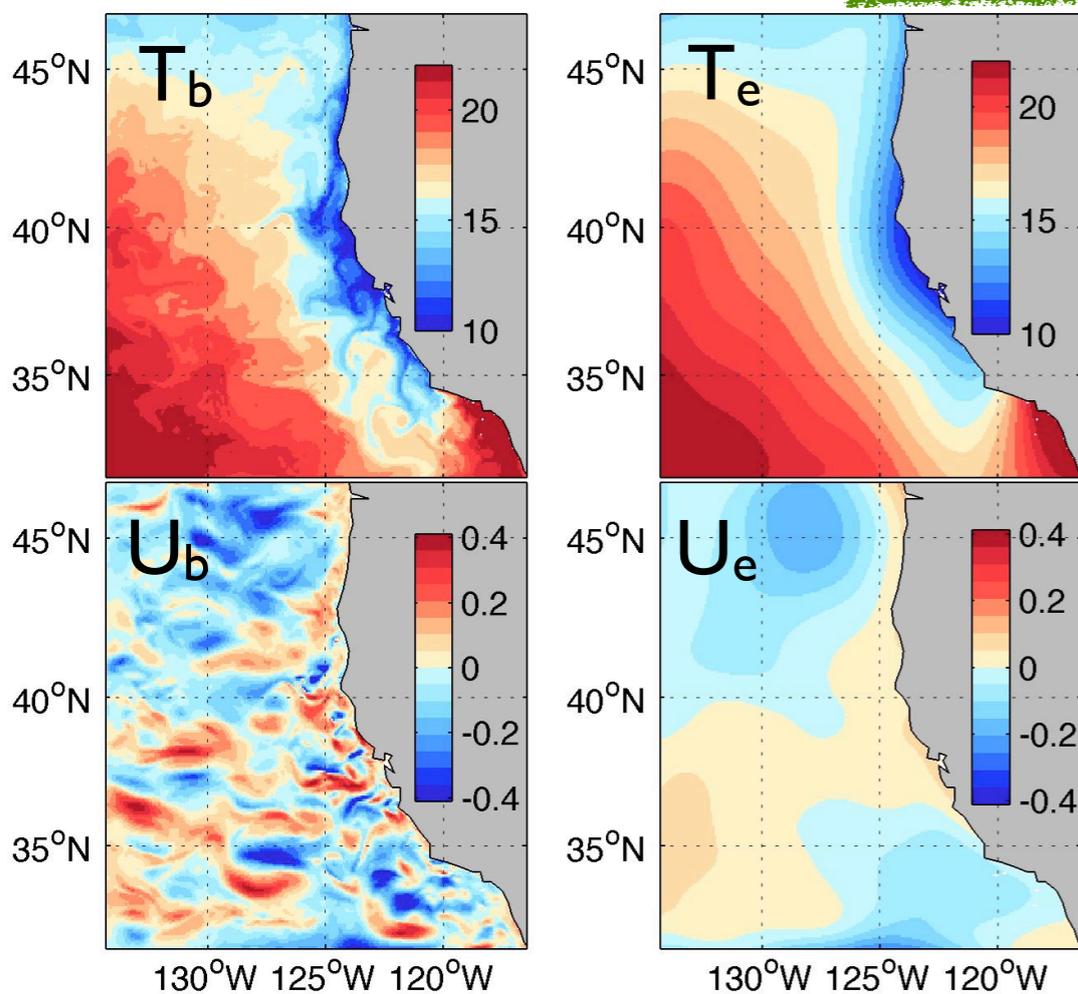
- Seo et al. 2007, 2014
- 7 km O-A resolutions
- 6-yr integration (2005-2010)

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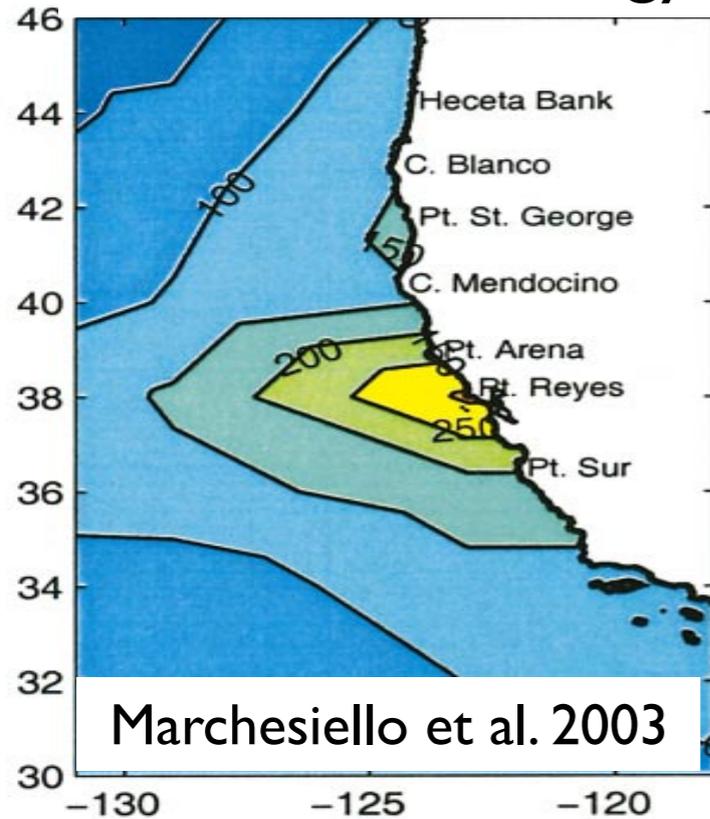
Removal of eddies with 5° loess filter
(Putrasahan et al. 2013)



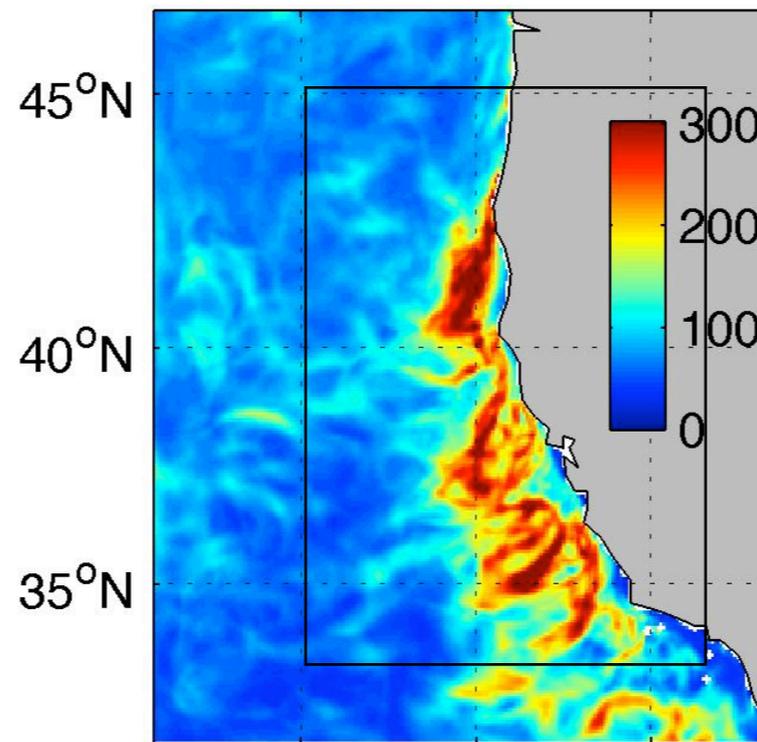
Exps	τ formulation includes			
CTL	T_b	T_e	U_b	U_e
no T_e	T_b		U_b	U_e
no U_e	T_b	T_e	U_b	

EKE significantly reduced by current effect on wind stress

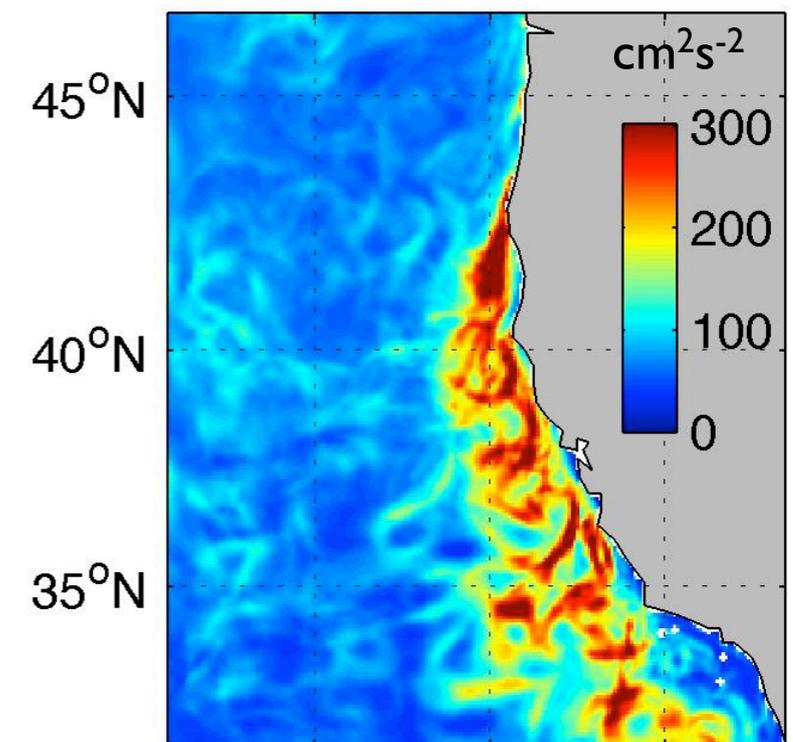
Drifter climatology



CTL

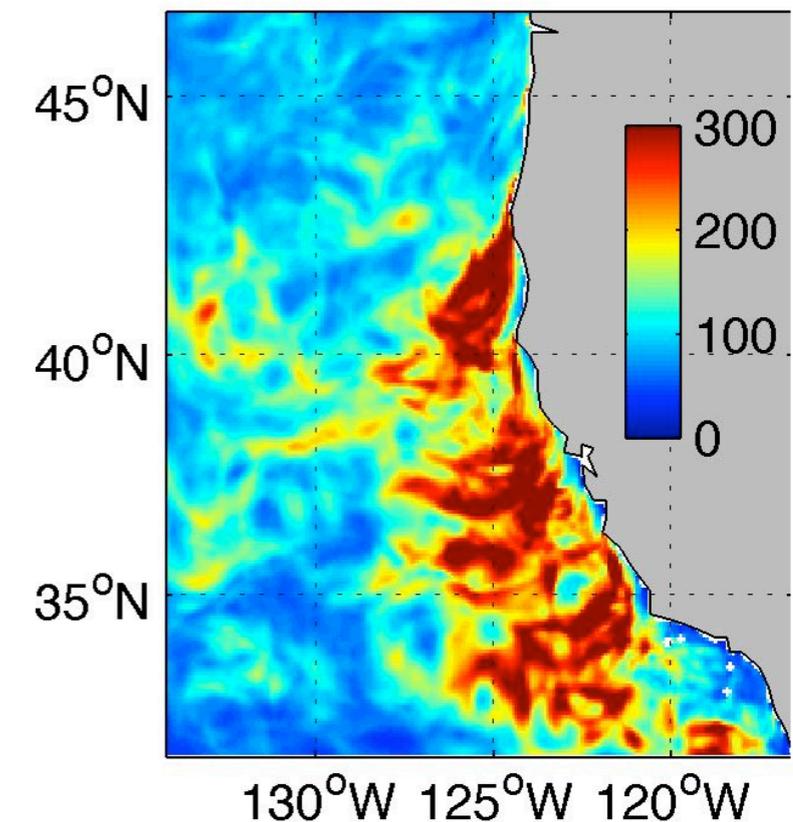


no T_e



- T_e no impact
- 25% weaker EKE with U_e
- Surface current dissipates the EKE

no U_e



JAS 2005-2010

Eddy kinetic energy budget

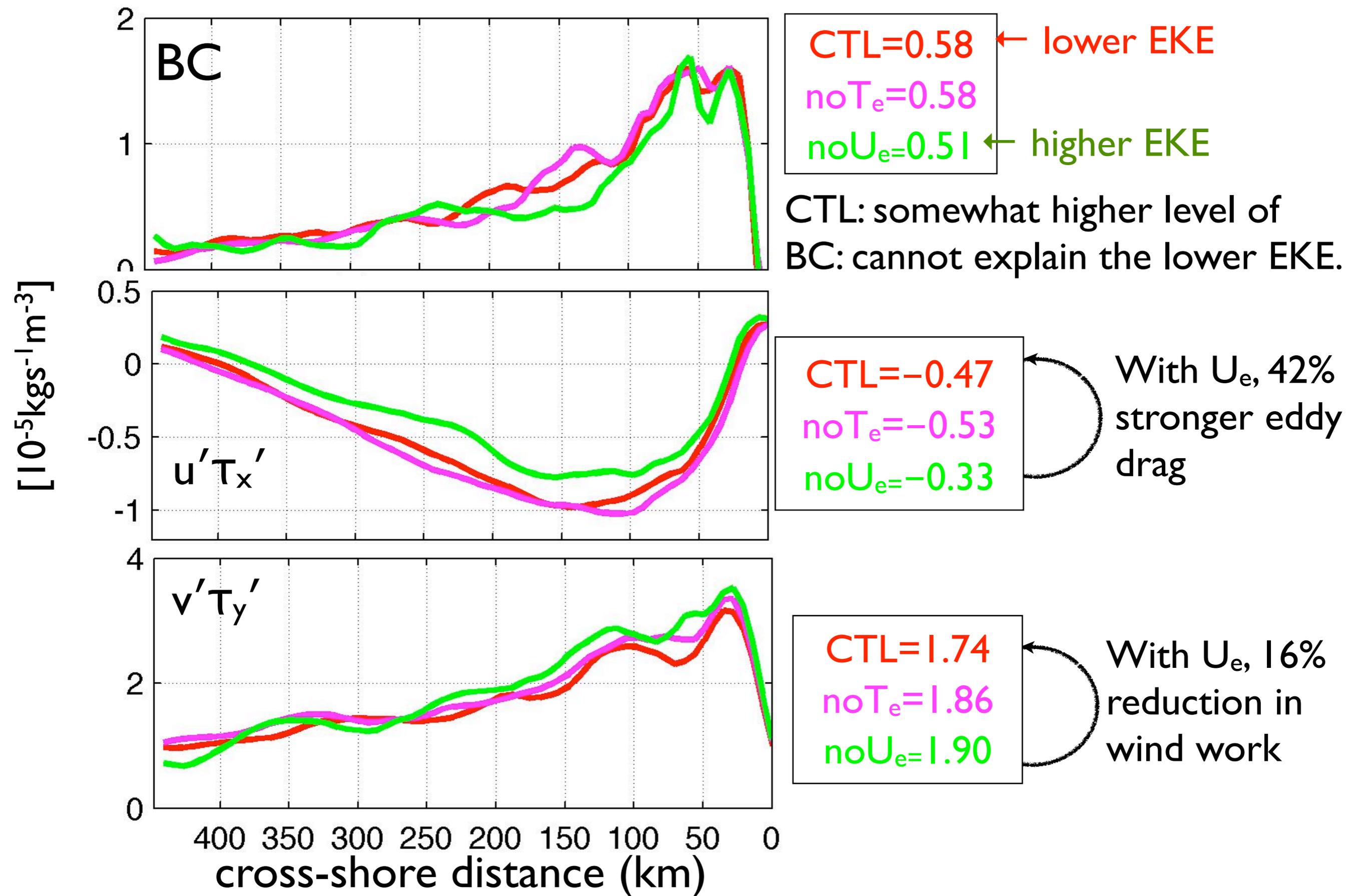
$$\begin{aligned}
 & Ke_t + \vec{U} \cdot \vec{\nabla} \vec{K}e + \vec{u}' \cdot \vec{\nabla} \vec{K}e + \vec{\nabla} \cdot (\vec{u}' p') = \\
 & + \rho_o (-\vec{u}' \cdot (\vec{u}' \cdot \vec{\nabla} \vec{U})) - g \rho' w' + \vec{u}' \cdot \vec{\tau}' + \varepsilon
 \end{aligned}$$

$\underline{K_m \rightarrow K_e}$
 barotropic
 conversion
 (BT)

$\underline{P_e \rightarrow K_e}$
 baroclinic
 conversion
 (BC)

Eddy-Wind terms:
 Wind work (P) if positive
 Eddy drag (ε) if negative

Reduced EKE is primarily due to enhanced eddy drag



Eddy-induced Ekman pumping velocity

Stern (1965) & Gaube et al. (2014)

$$W_{tot} = \frac{1}{\rho_o} \nabla \times \left(\frac{\tau}{(f + \zeta)} \right)$$

$$\tilde{W}_{tot} = W_{cur} + W_{SST}$$

background wind stress

$$= \underbrace{\frac{\nabla \times \tilde{\tau}}{\rho_o (f + \zeta)}}_{W_{lin}} - \underbrace{\frac{1}{\rho_o (f + \zeta)^2} \left(\tilde{\tau}^y \frac{\partial \zeta}{\partial x} - \tilde{\tau}^x \frac{\partial \zeta}{\partial y} \right)}_{W_{\zeta}} + \underbrace{\frac{\beta \tilde{\tau}^x}{\rho_o (f + \zeta)^2}}_{W_{\beta}} + \underbrace{\frac{\nabla \times \tau'_{SST}}{\rho_o (f + \zeta)}}_{W_{SST}}$$

W_{lin}

W_{ζ}

W_{β}

W_{SST}

Curl-induced
linear Ekman pumping

Vorticity gradient-induced
nonlinear Ekman pumping

β Ekman pumping
(negligible)

SST-induced Ekman
pumping
(Chelton et al. 2004)

Eddy-induced Ekman pumping velocity

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W_{lin}

W_{ζ}

W_{β}

W_{SST}

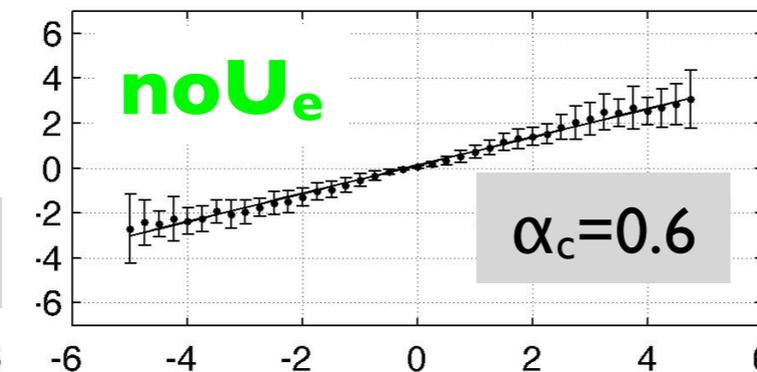
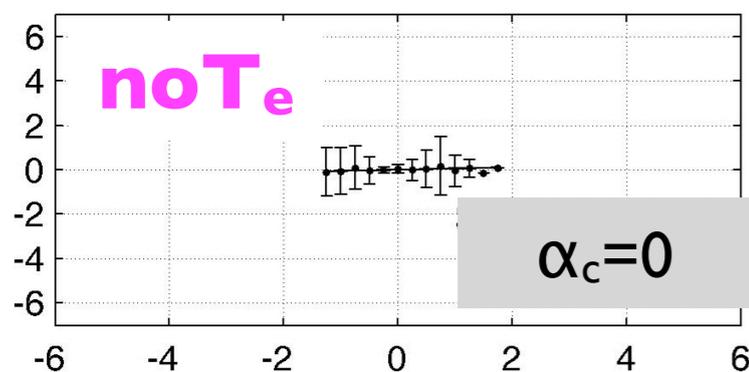
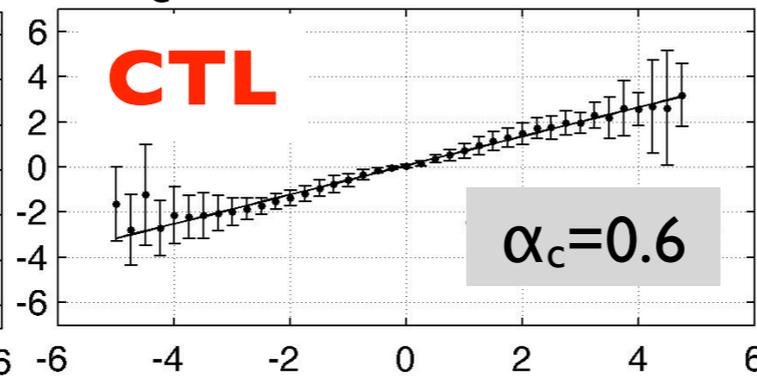
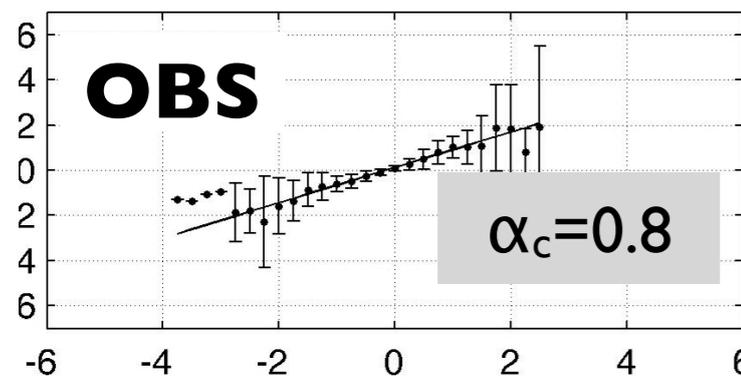
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Vorticity gradient-induced nonlinear Ekman pumping

β Ekman pumping (negligible)

SST-induced Ekman pumping (Chelton et al. 2004)

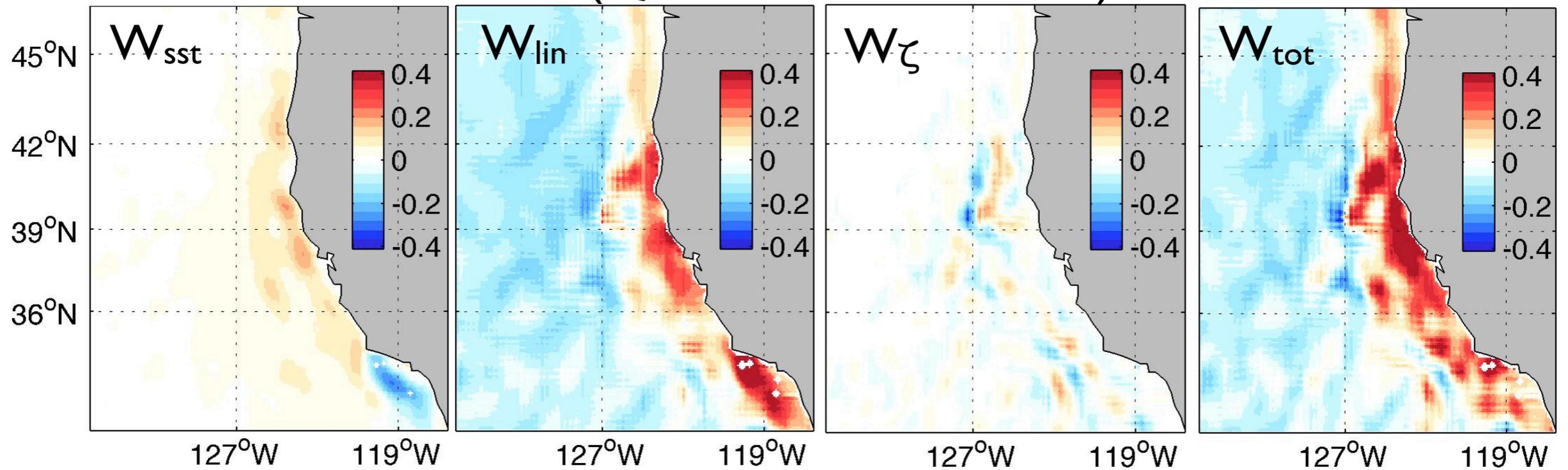
$\nabla \times \tau'$ vs $\nabla_c T'$



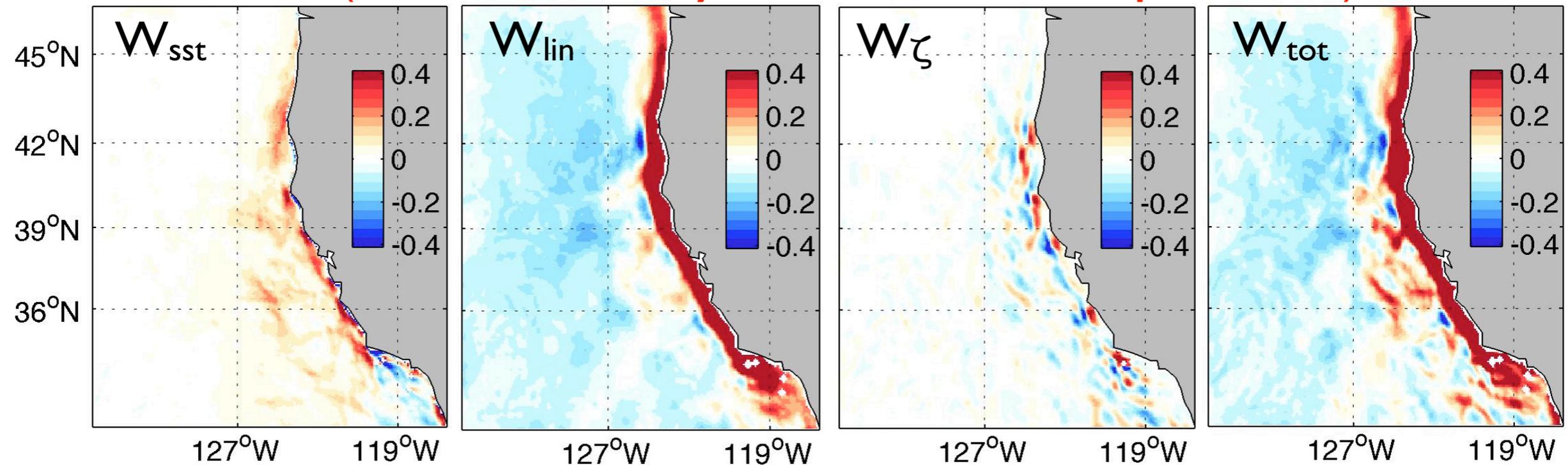
$$W_{SST} = \frac{\nabla \times \tau'_{SST}}{\rho_o (f + \zeta)} \approx \frac{\alpha_c \nabla_c SST}{\rho_o (f + \zeta)}$$

Ekman pumping velocity JAS climatology

OBS (QuikSCAT & AVISO)



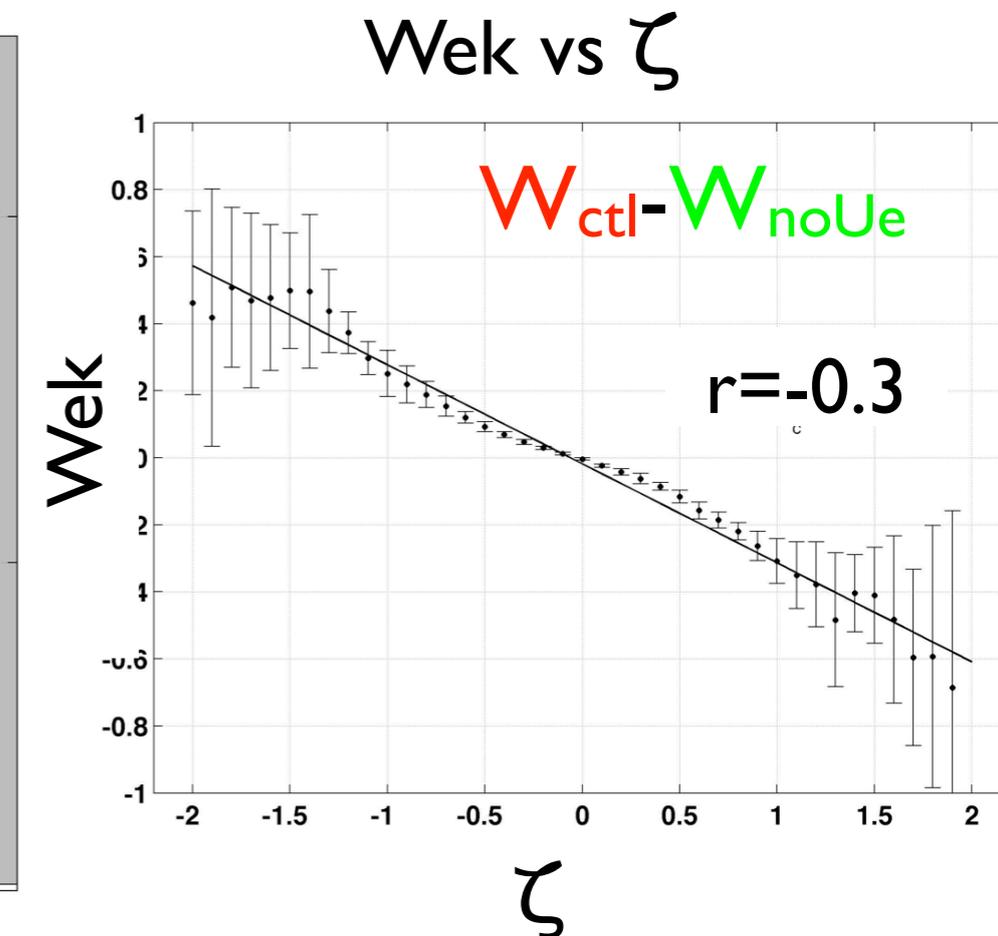
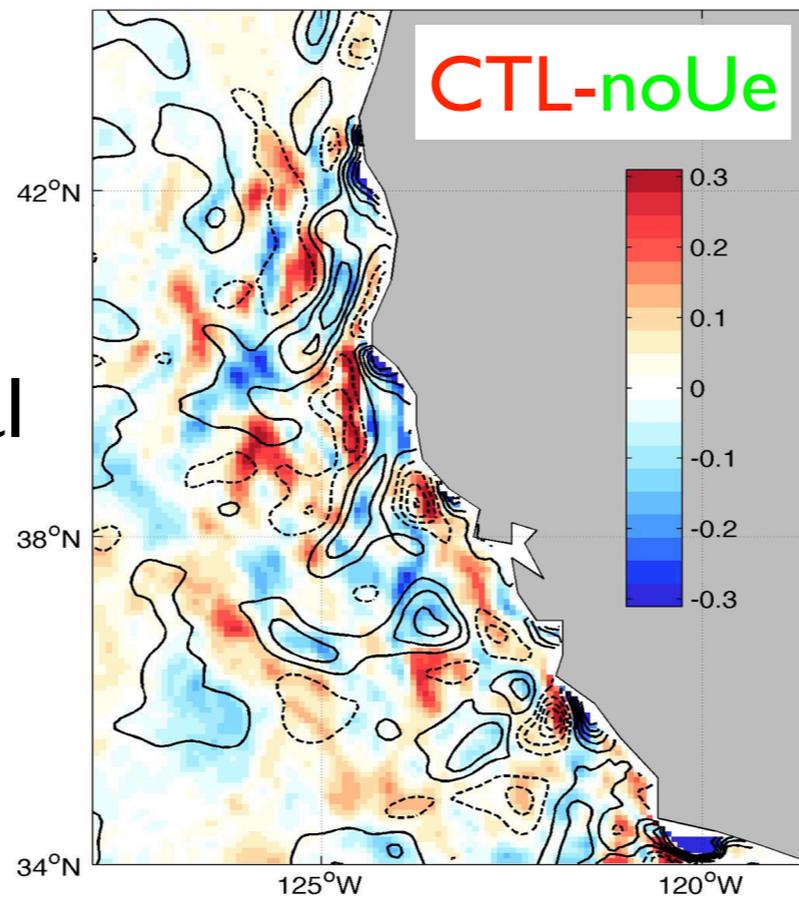
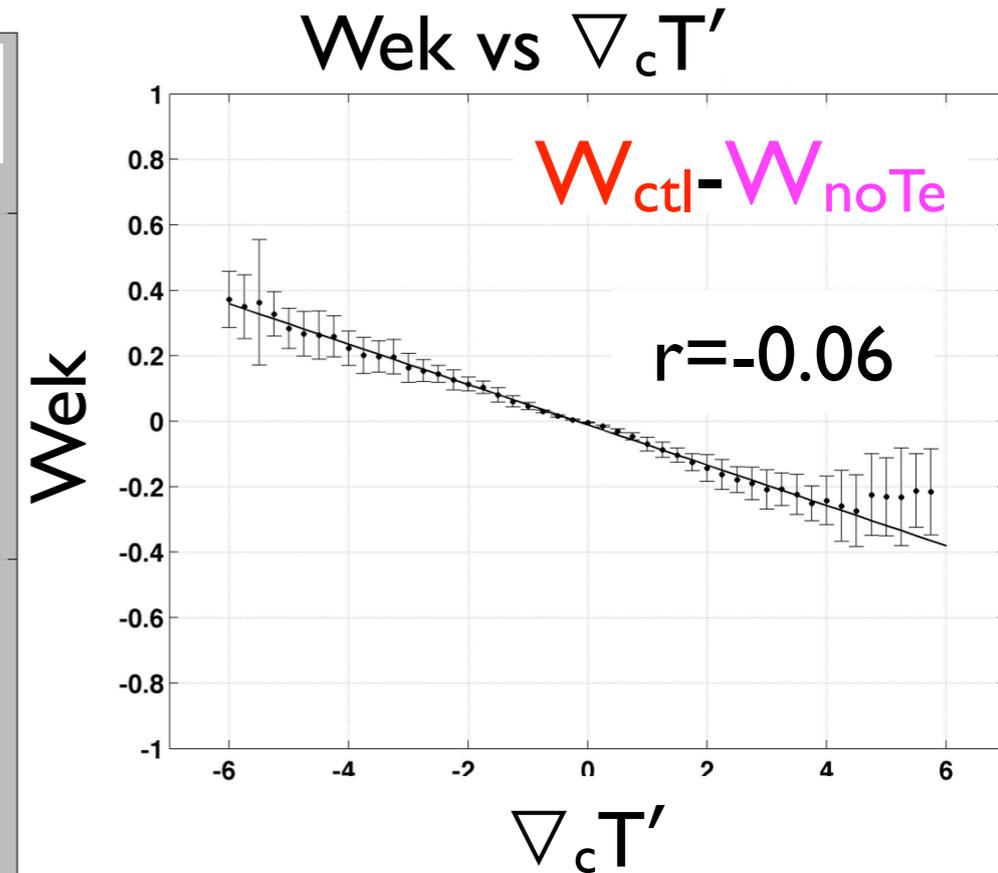
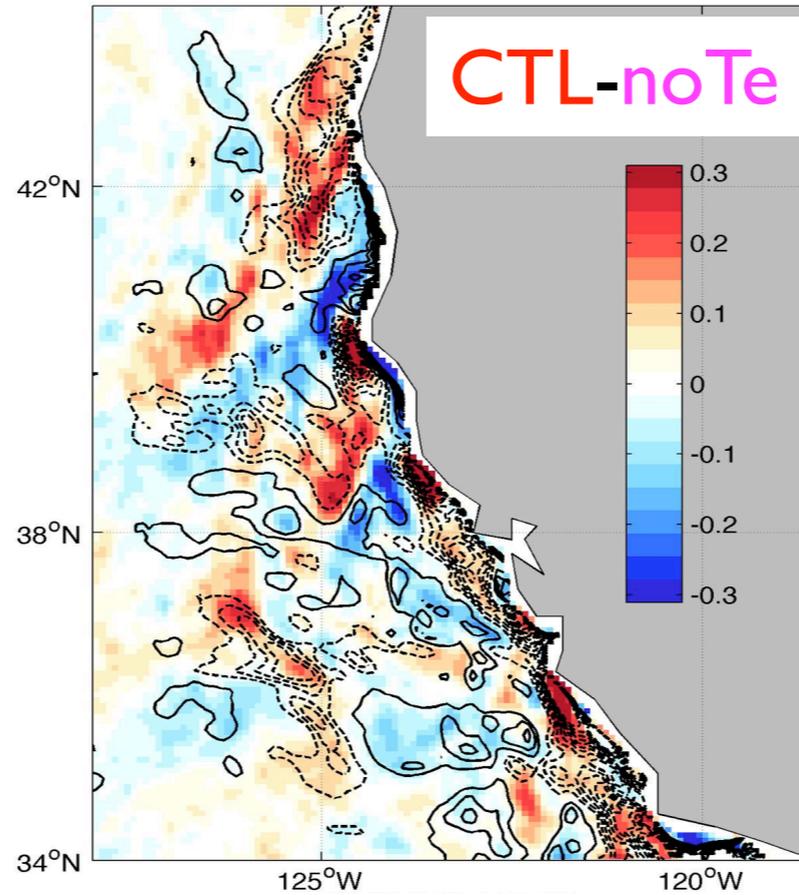
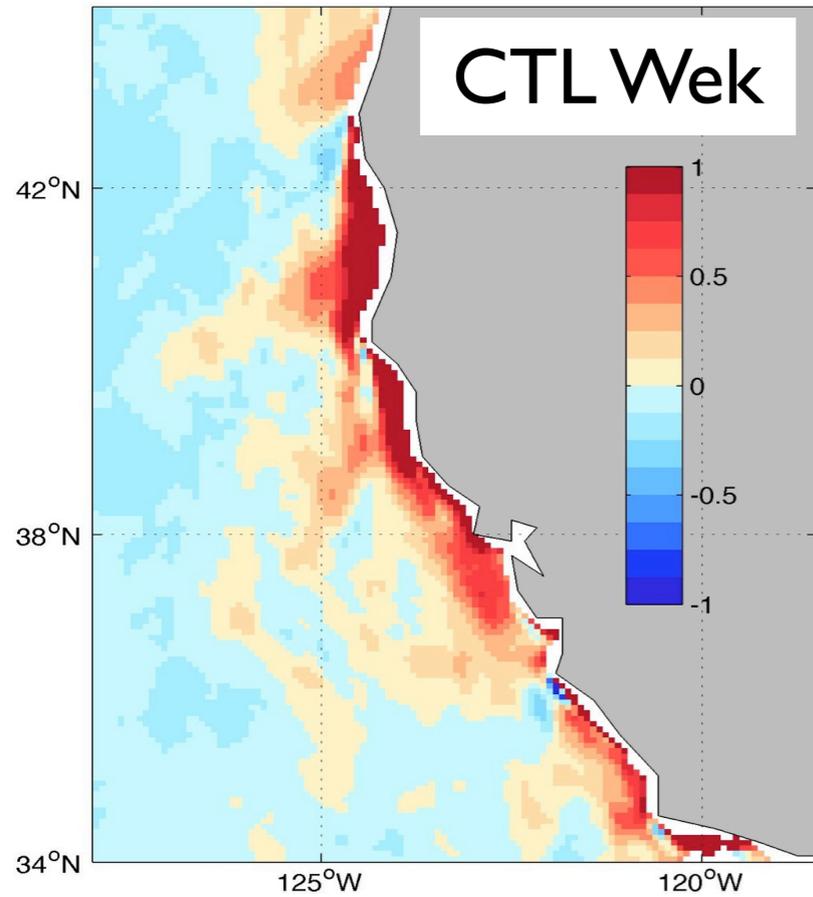
CTL (with both eddy current and temperature)



m/day

JAS 2005-2009

SST- and current-induced Ekman pumping velocity



- SST and current induce perturbation W_{ek} of comparable magnitudes but with distinctive spatial patterns.
- indicative of different feedback processes

Summary

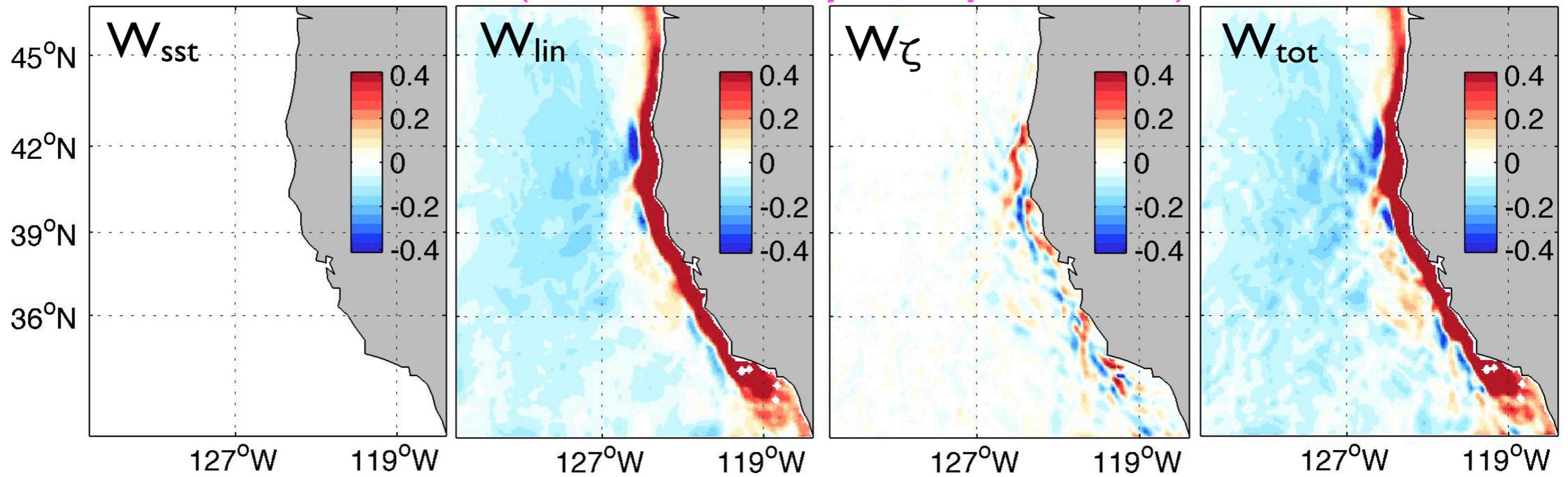
- Surface EKE is weakened almost entirely due to mesoscale current effect on wind stress.
 - SST has no impact (at odds with some previous studies)
 - EKE budget: eddies primarily enhance the eddy drag, and weaken the wind work of secondary importance.
- Change in eddy drag means changes in Ekman pumping velocities
 - Eddy-current and eddy-SST produce Ekman pumping velocity climatologies of comparable magnitudes and different distributions.
 - Implying different feedback processes, a subject of ongoing study.

Thanks!
hseo@whoi.edu

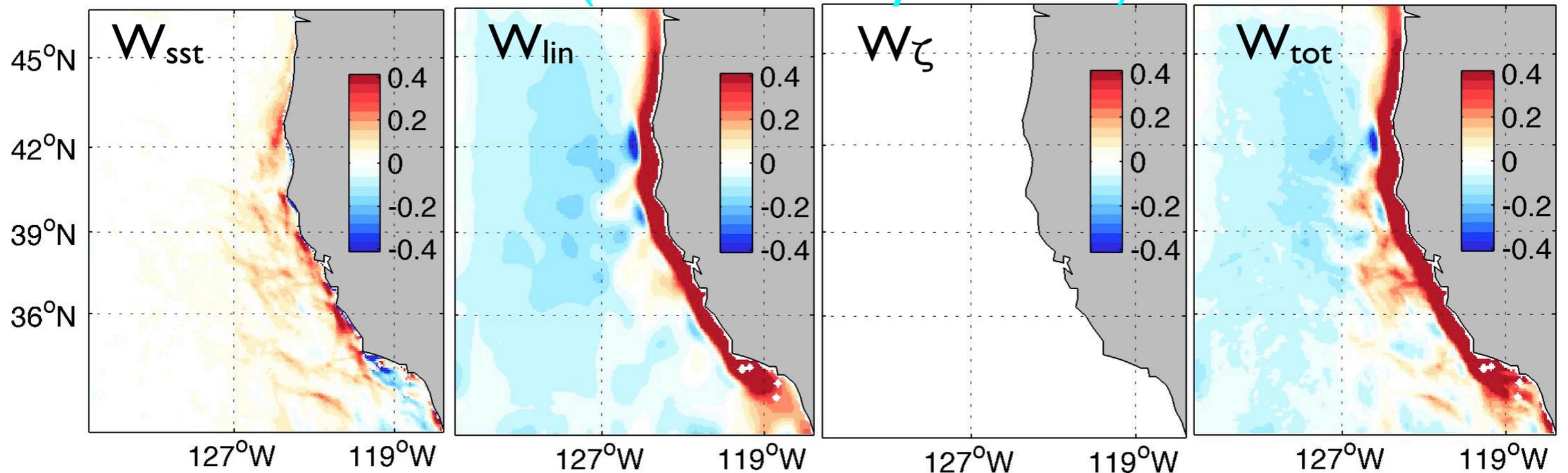
This study gratefully acknowledges NSF OCE-09060770.

Ekman pumping velocity JAS climatology

no T_e (without eddy temperature)



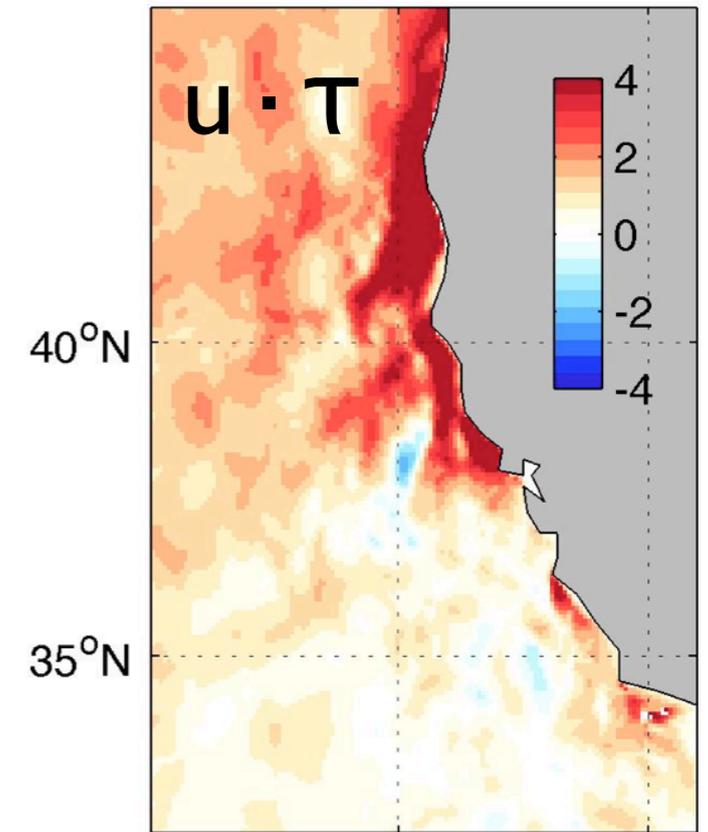
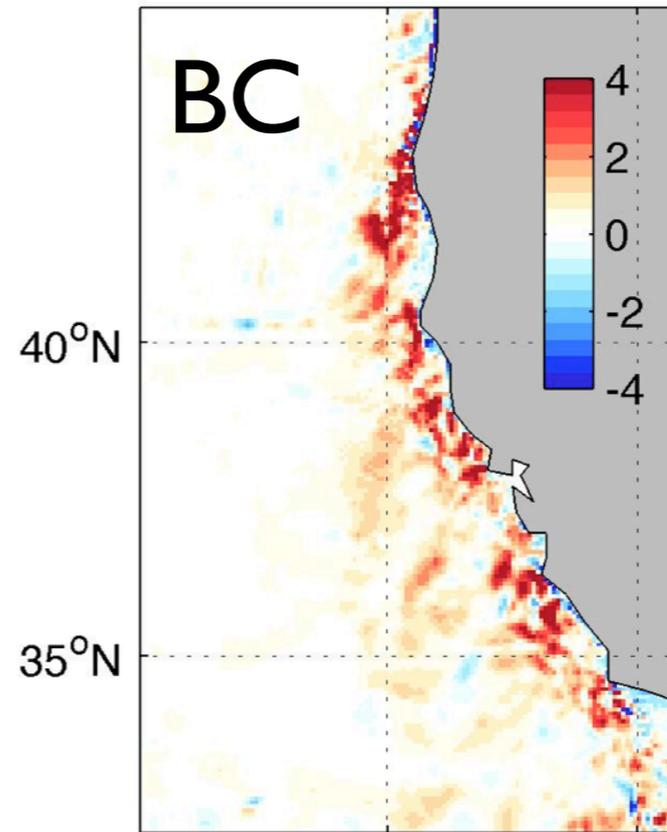
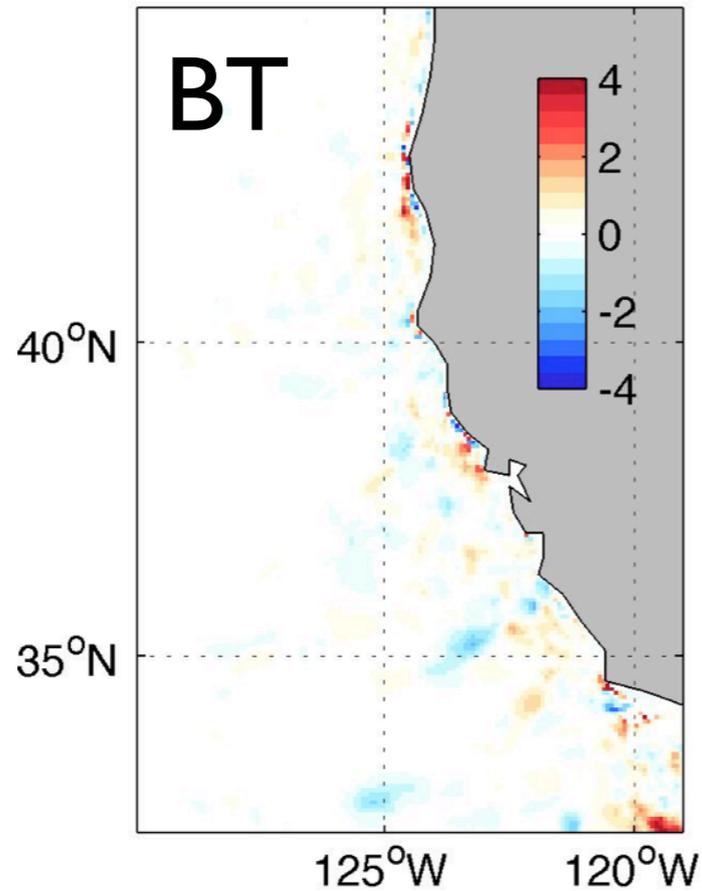
no U_e (without eddy current)



m/day

JAS 2005-2009

Summertime EKE budget in CTL

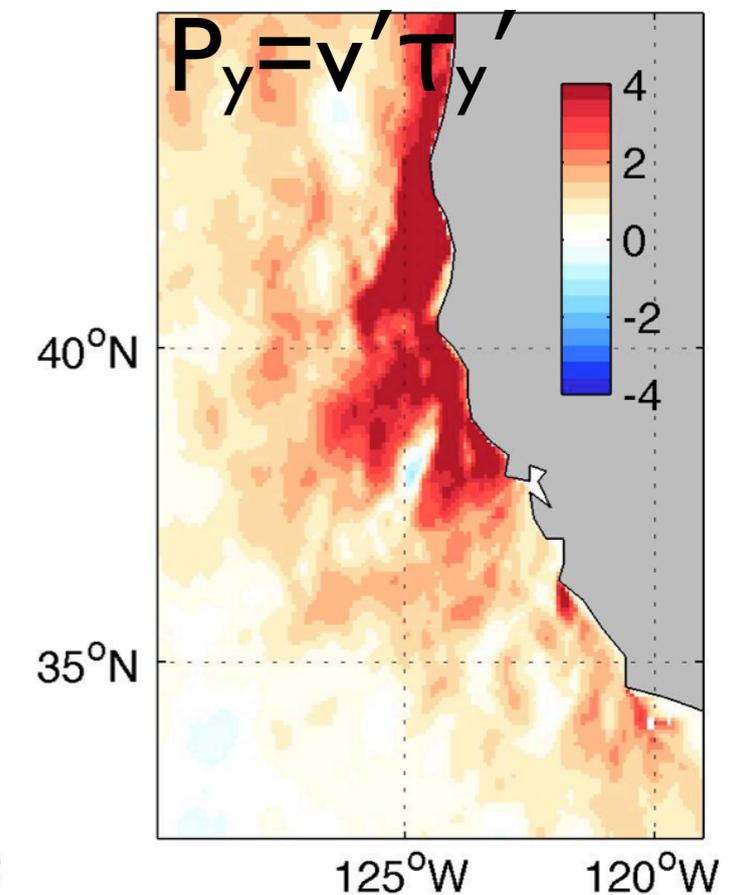
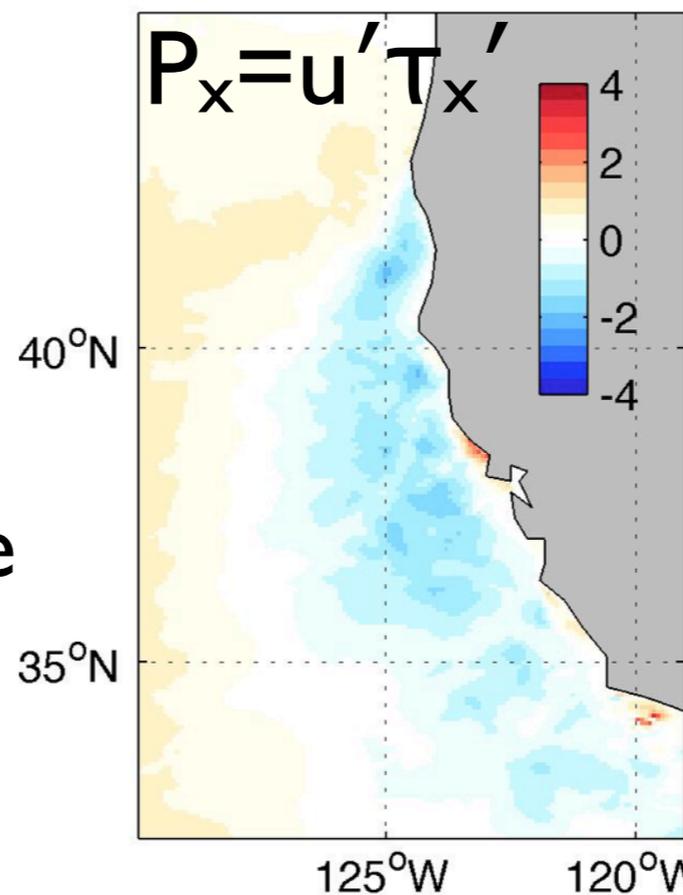


- Eddy wind work is a primary source

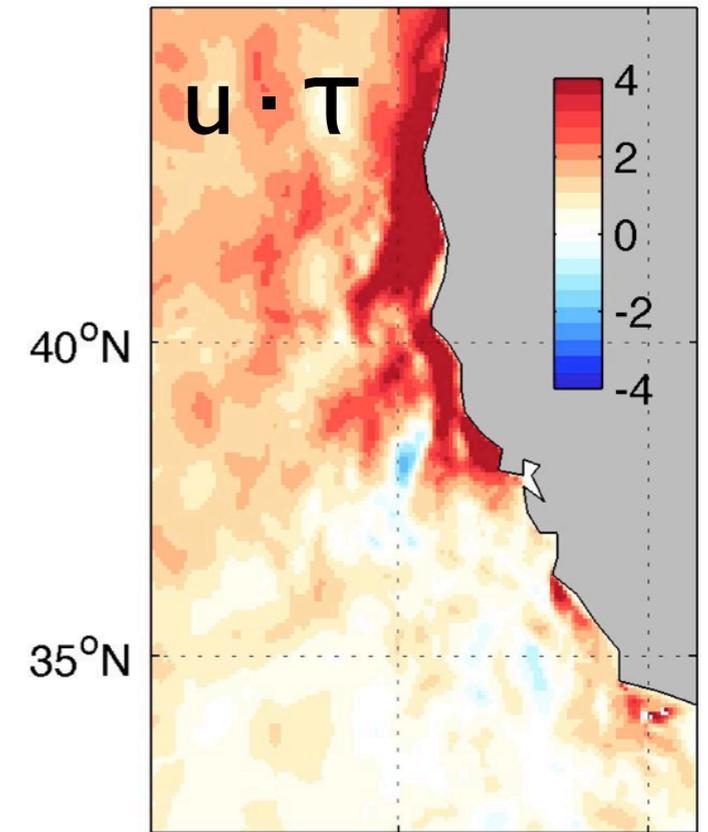
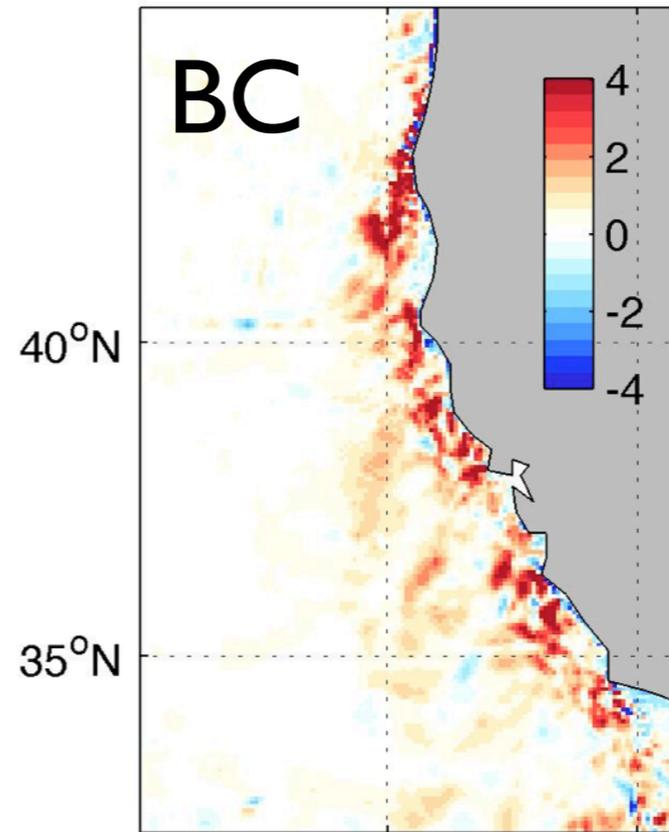
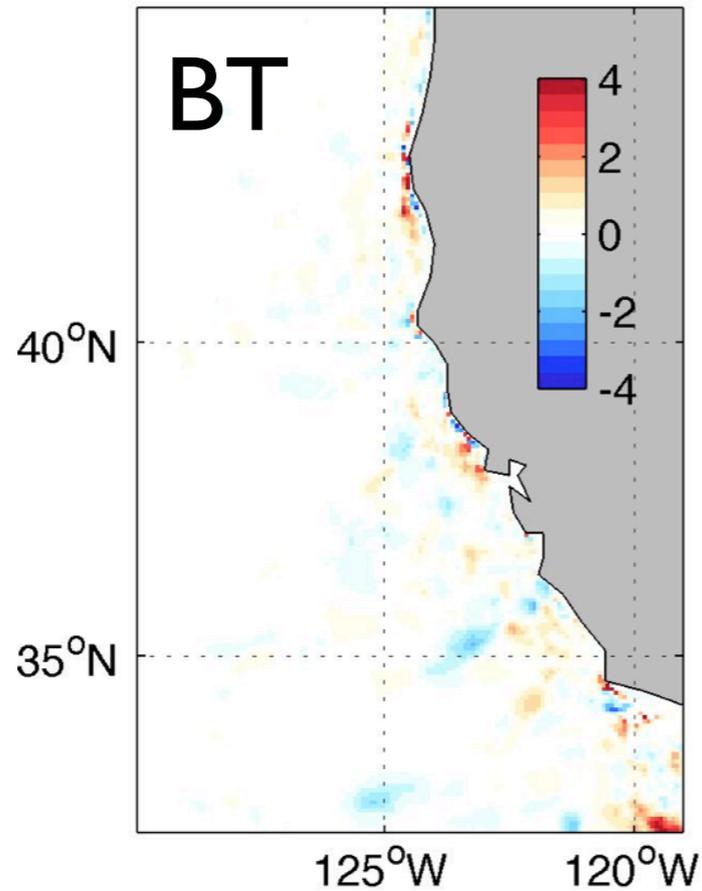
- BC secondary and BT negligible

- $v' \tau_y'$: Source of EKE nearshore

- $u' \tau_x'$: Dissipating EKE in the upwelling zone



Summertime EKE budget in CTL

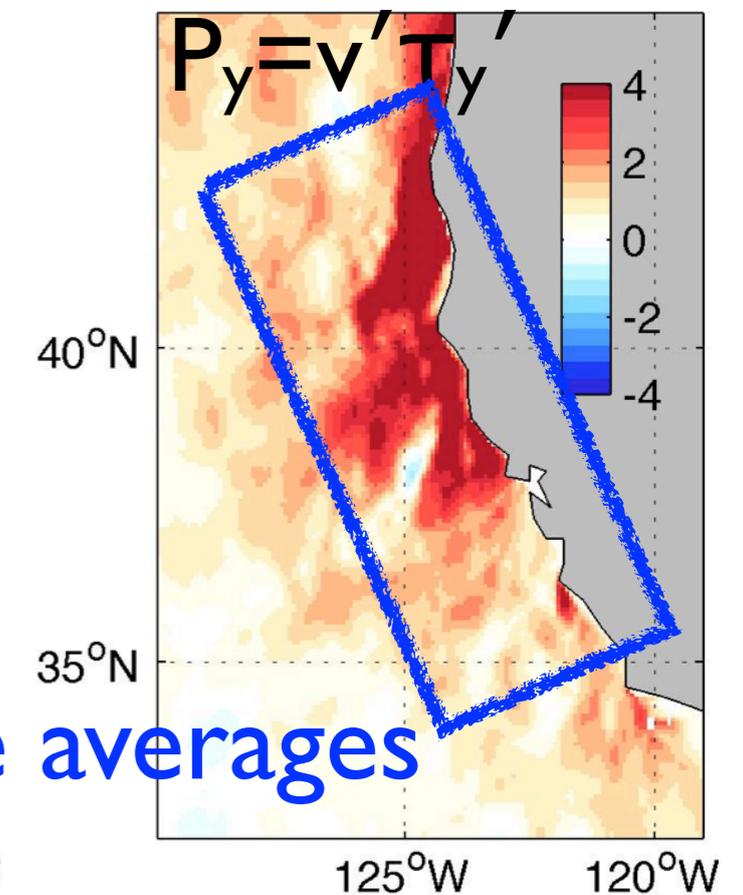
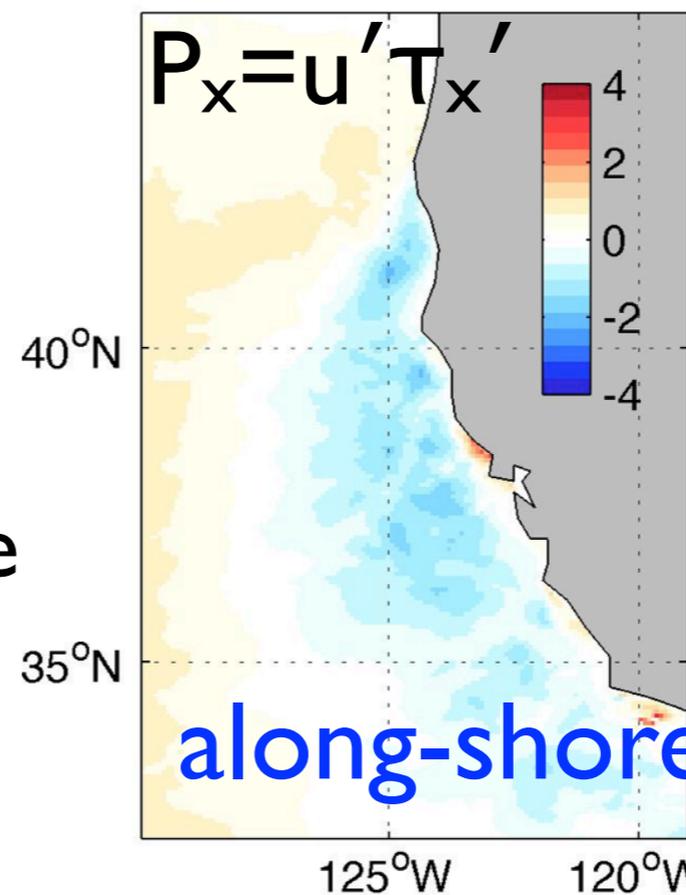


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along-shore averages