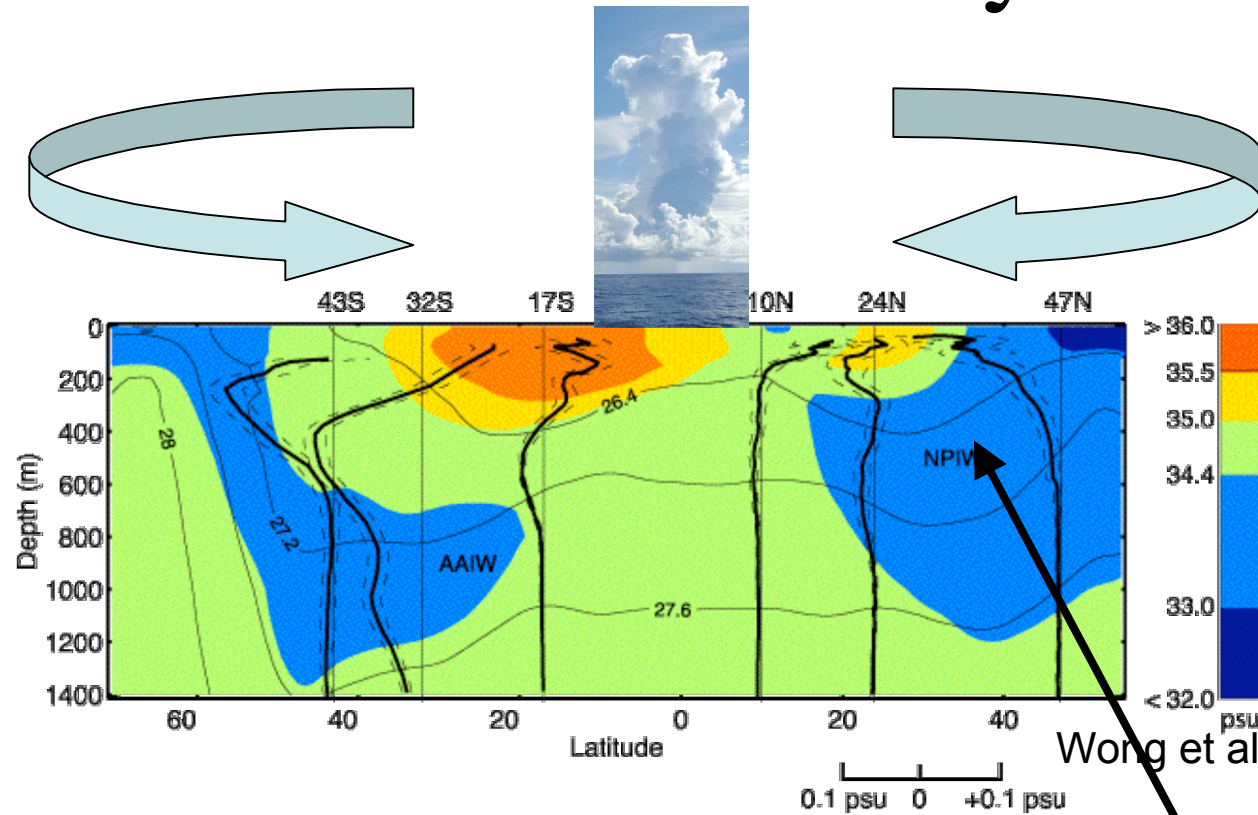


# The Upper Cell

1. Climate change illuminates processes operating in the basic state
2. Link between interior and mixed layer key
3. Role of eddies complex - much vertical structure hidden by net results

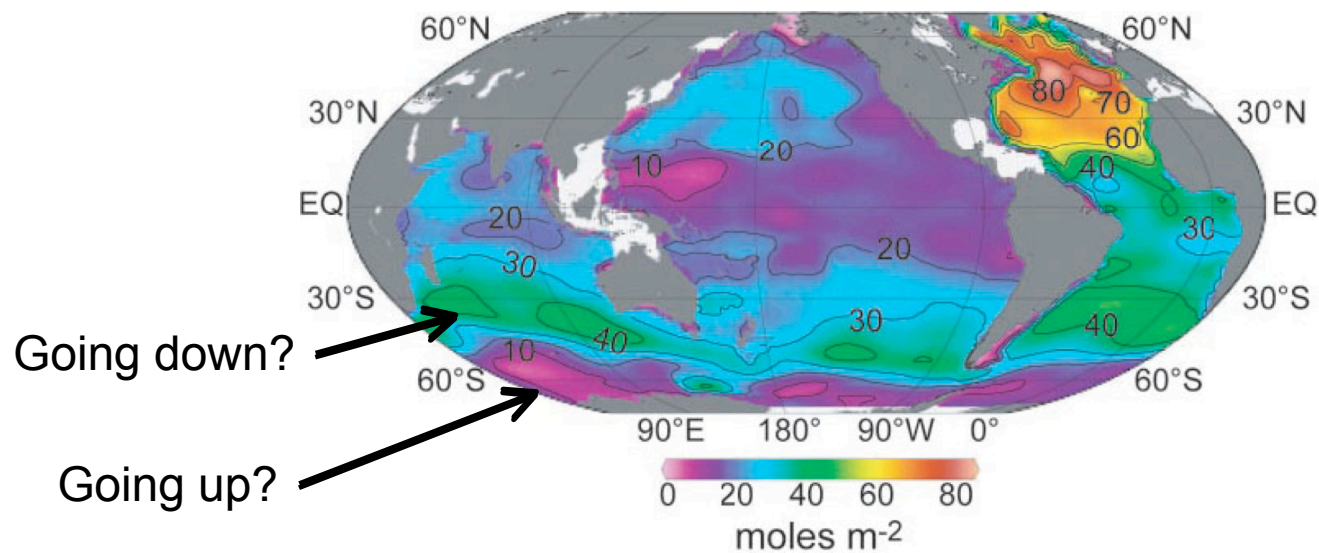
# Enhanced Water Cycle



Salinity suggests intermediate freshwater “cell” but interior climate signal is strongest in mode waters

mean advection versus diffusion from source region

# Subduction and the Upper Cell



Ekman transport  
Mean flow  
Eddy transport

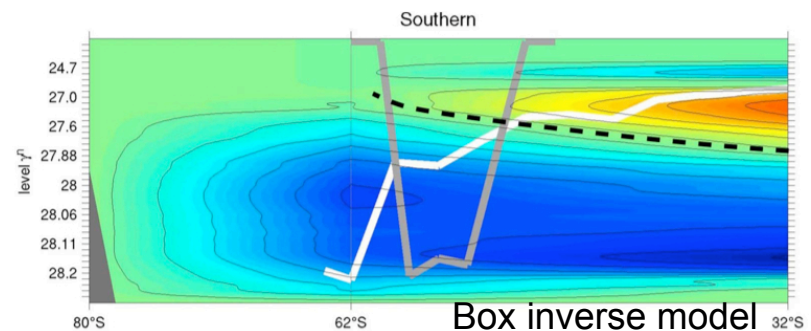
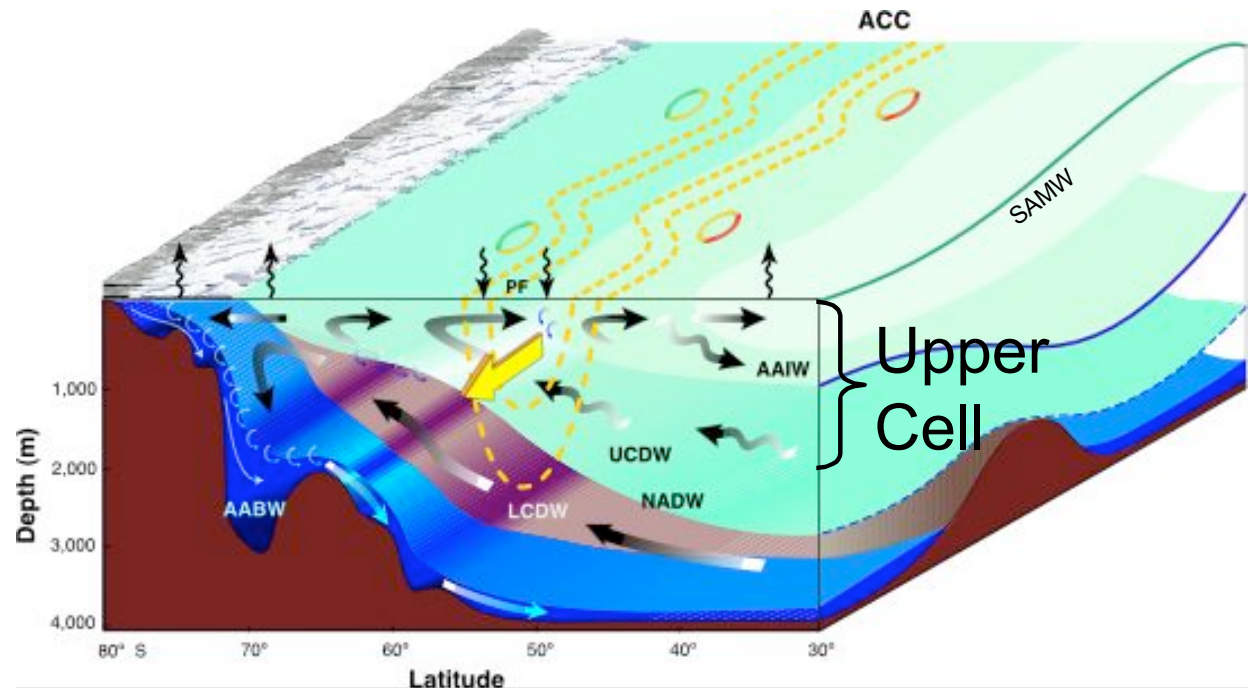
# Southern Ocean

## *The Upper Cell*

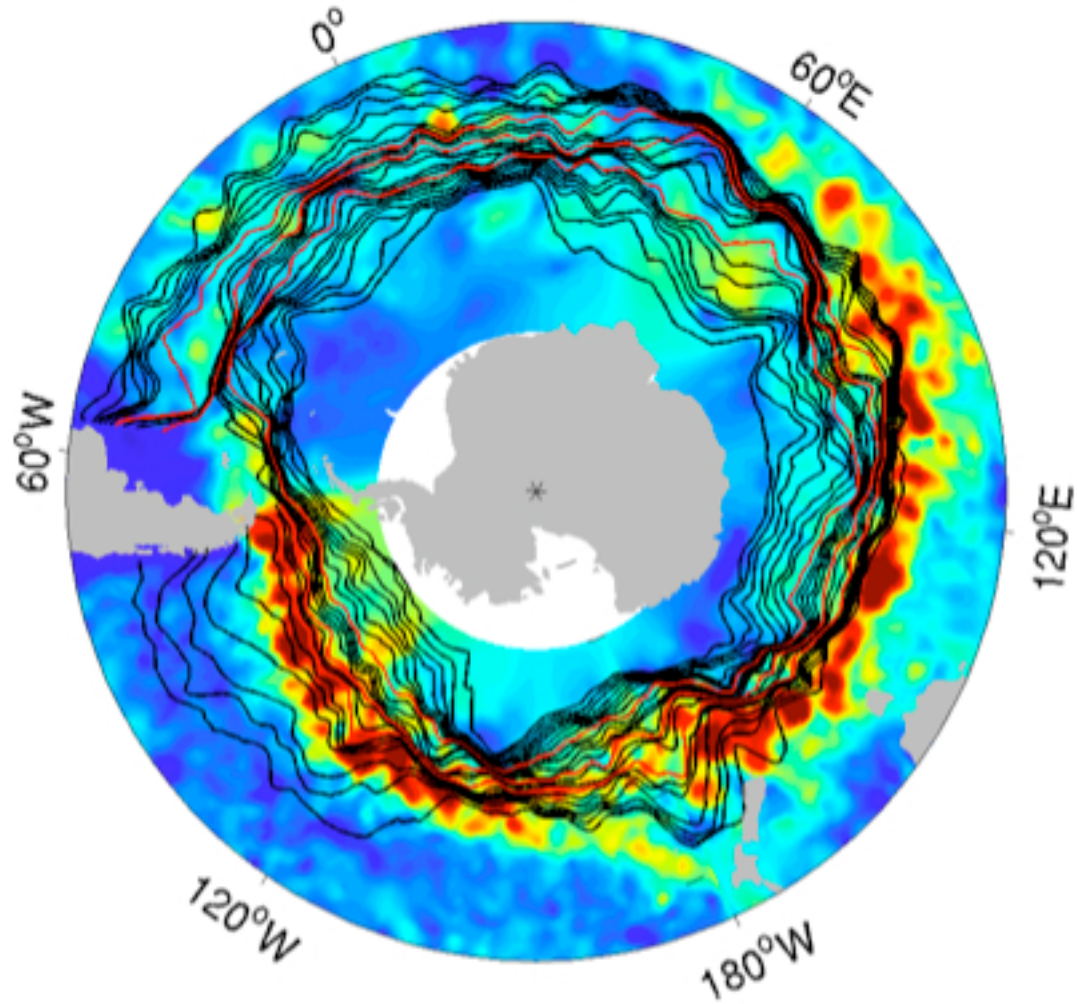
Brings warm UCDW up  
 Brings natural carbon up  
 Brings low oxygen up

Takes surface carbon down  
 Takes surface freshening down

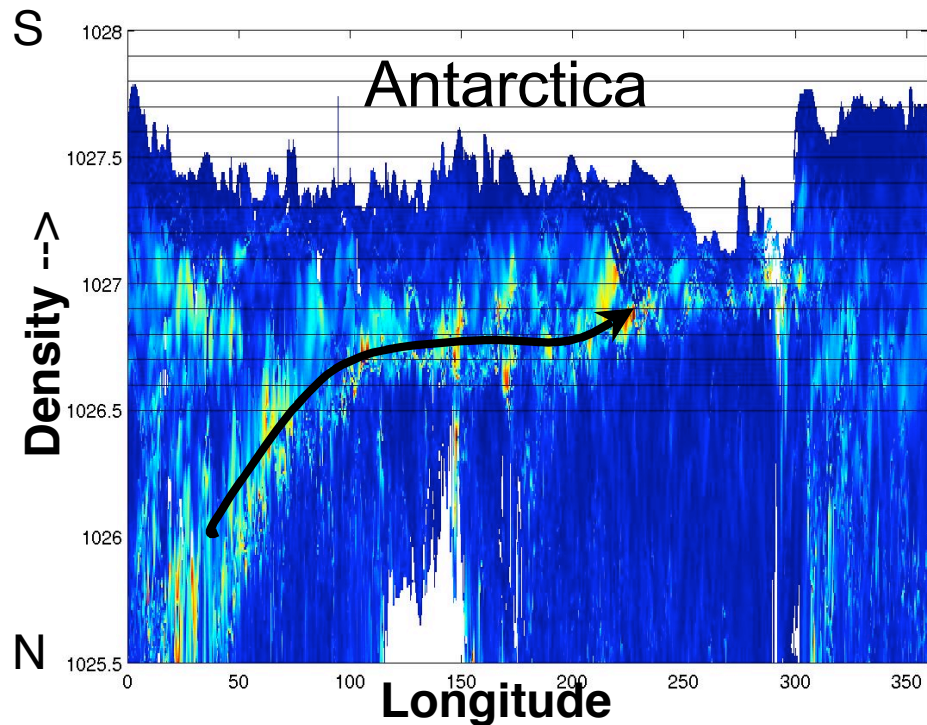
Has mixed layer processes imbedded



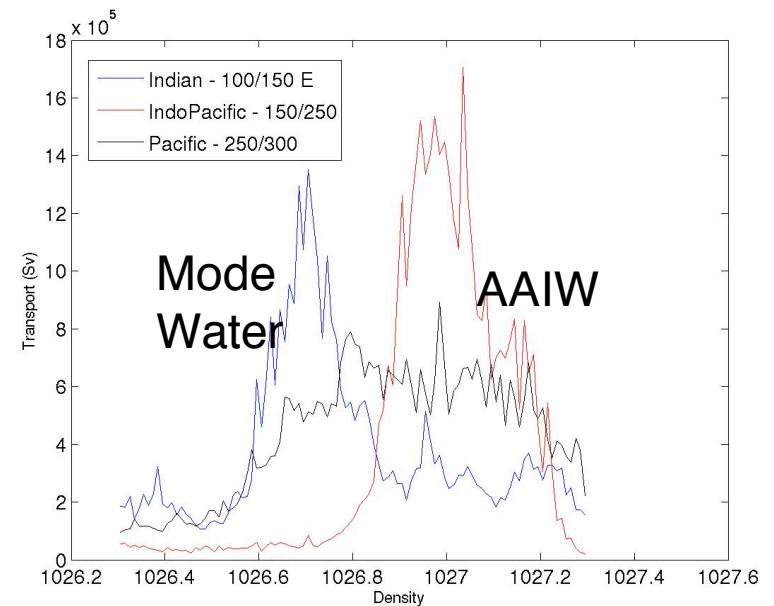
# Mean Flow and MLD



# ACC and mixed layer mass transport

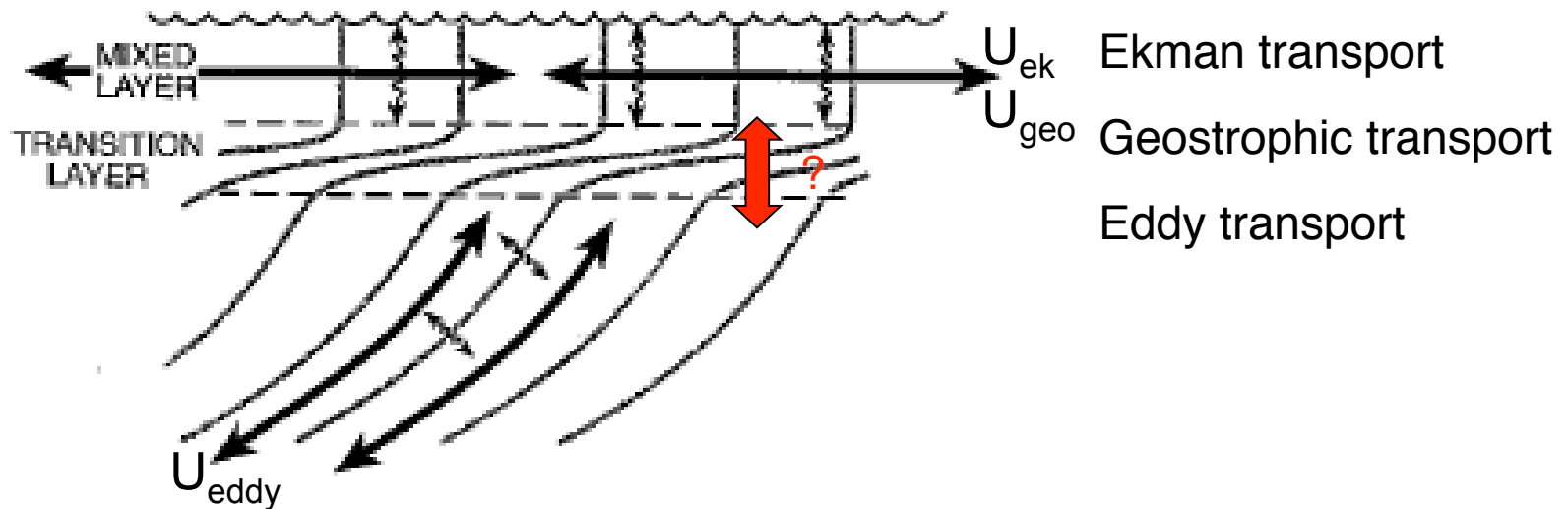


Geostrophic transport/length *in mixed layer*  
density classes -> accumulation in denser  
deeper mixed layers



Net ML transport by density:  
20 Sv transferred from mode  
water to AAIW *in ML*

# Mixed Layer Transports

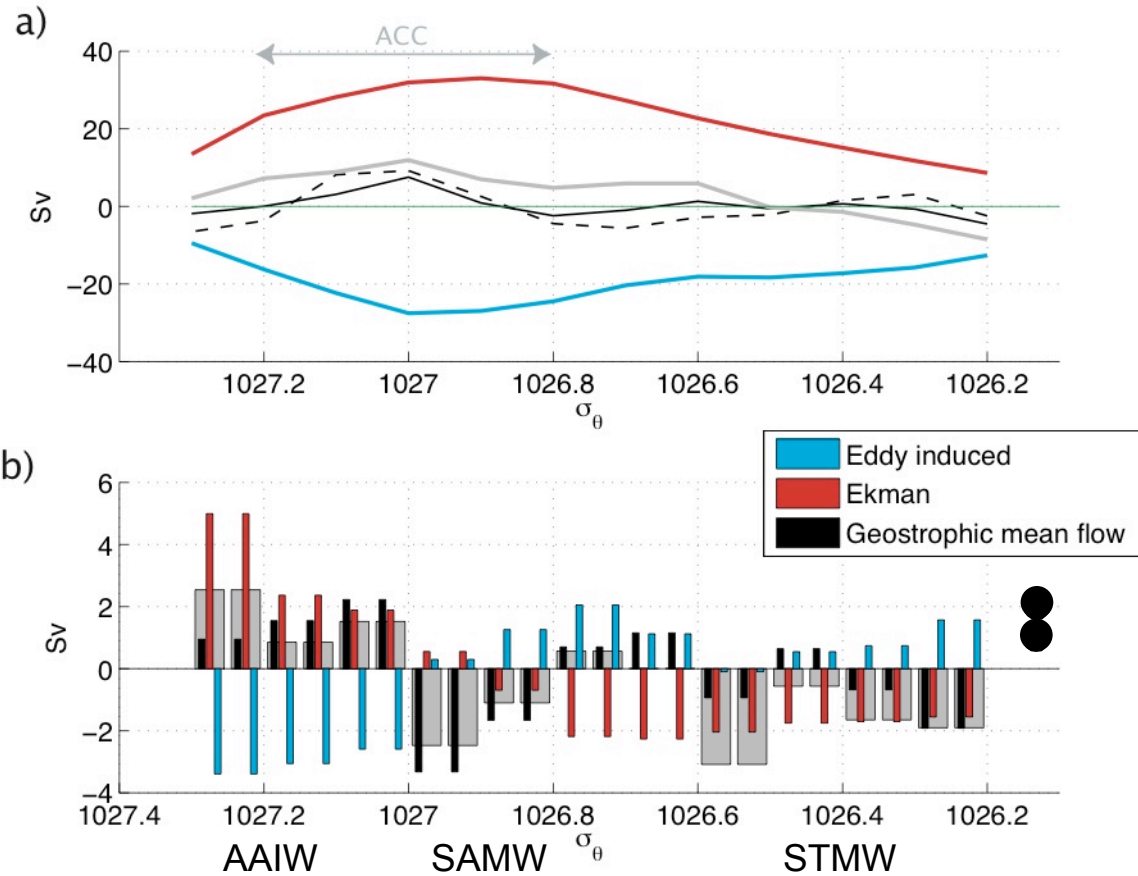


Gent McWilliams interior (sub-mixed layer) eddy parameterization

Treguier et al. (97) everything spreads into the ML.

Ferrari & McWilliams (08): only a portion of the transition layer transport spreads into the ML.

# Zonal Average View of Subduction



$$\text{Subduction} = \nabla \cdot (U_{\text{eddy}} + U_{\text{ek}} + U_{\text{geo}})$$

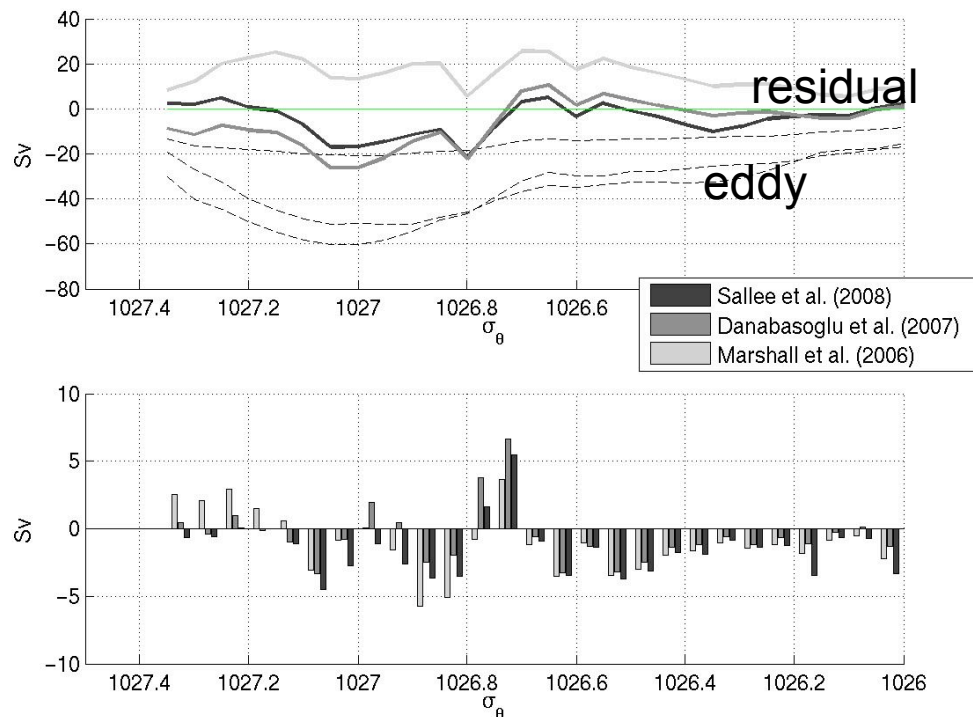


# Sensitivity to lateral mixing coefficient K

$$U_{\text{eddy}} = K \times \text{isopycnal slope}$$

Use K from Sallee et al (2008), compare to others

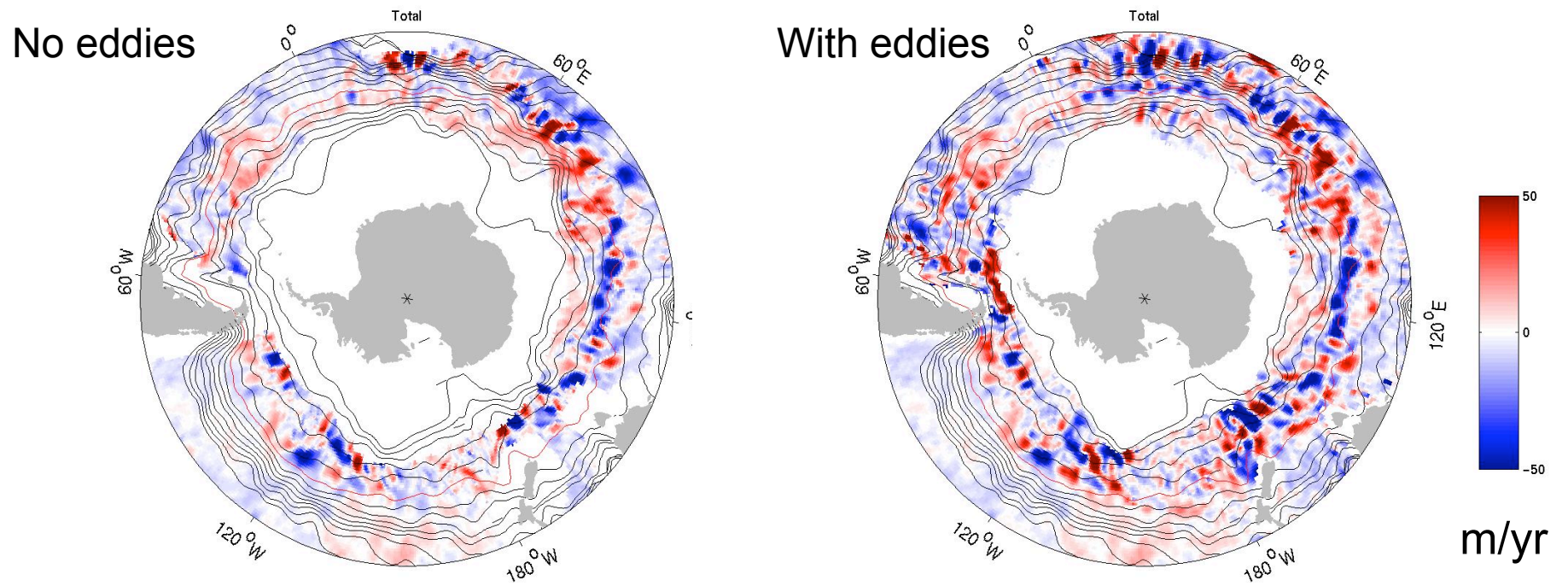
- surface estimate, projected throughout seasonal ML



Net cross-isopycnal flow very different

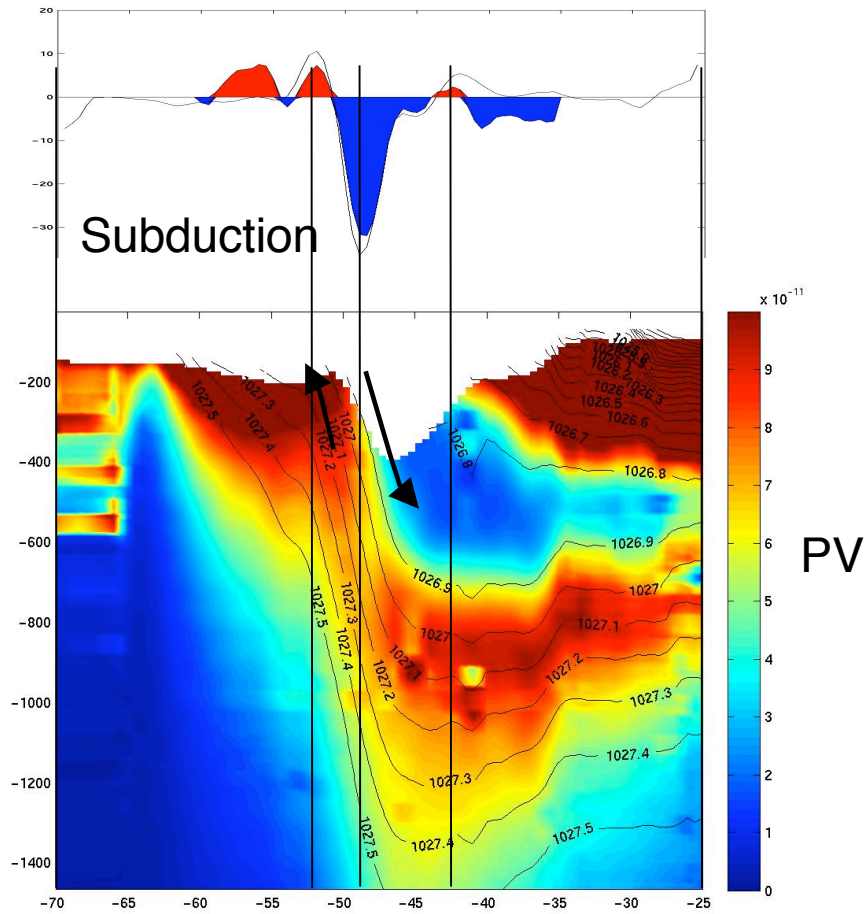
Net subduction not very different (shape is dominated by slope of isopycnal)

# Maps of Subduction

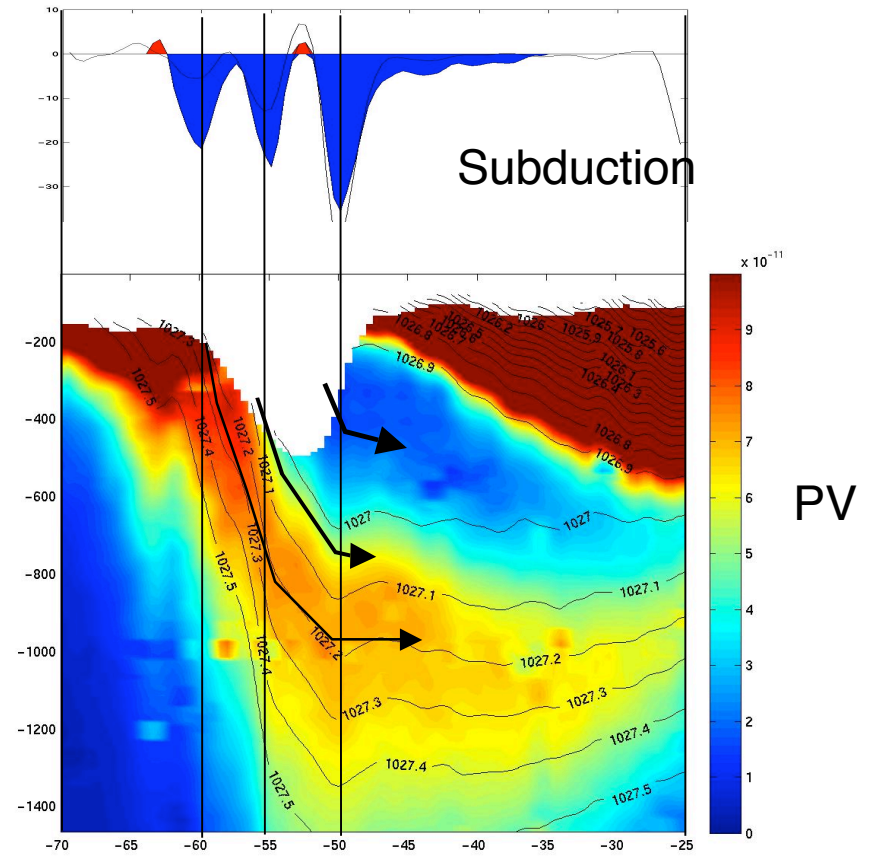


# Sections of Subduction

## Indian (110-130 E)

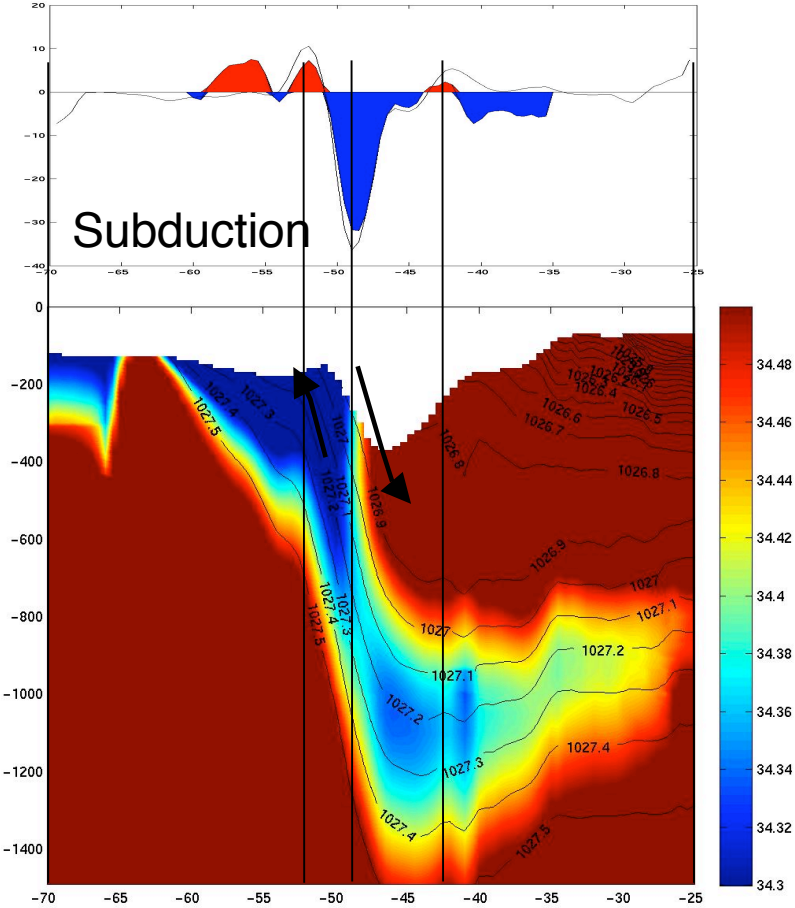


## Pacific (120 130 W)

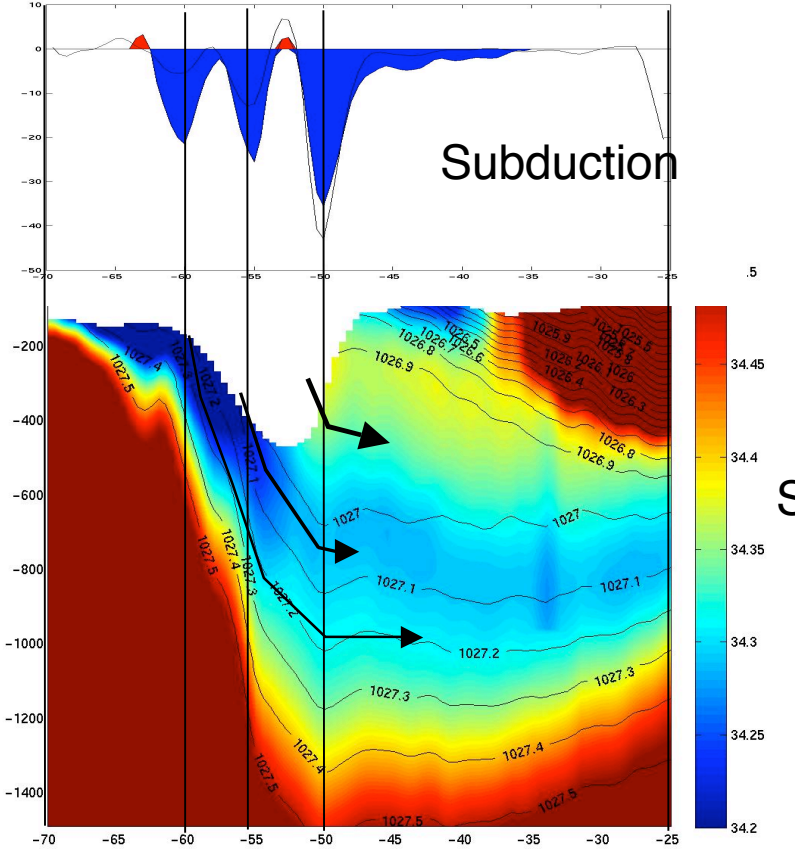


# Sections of Subduction

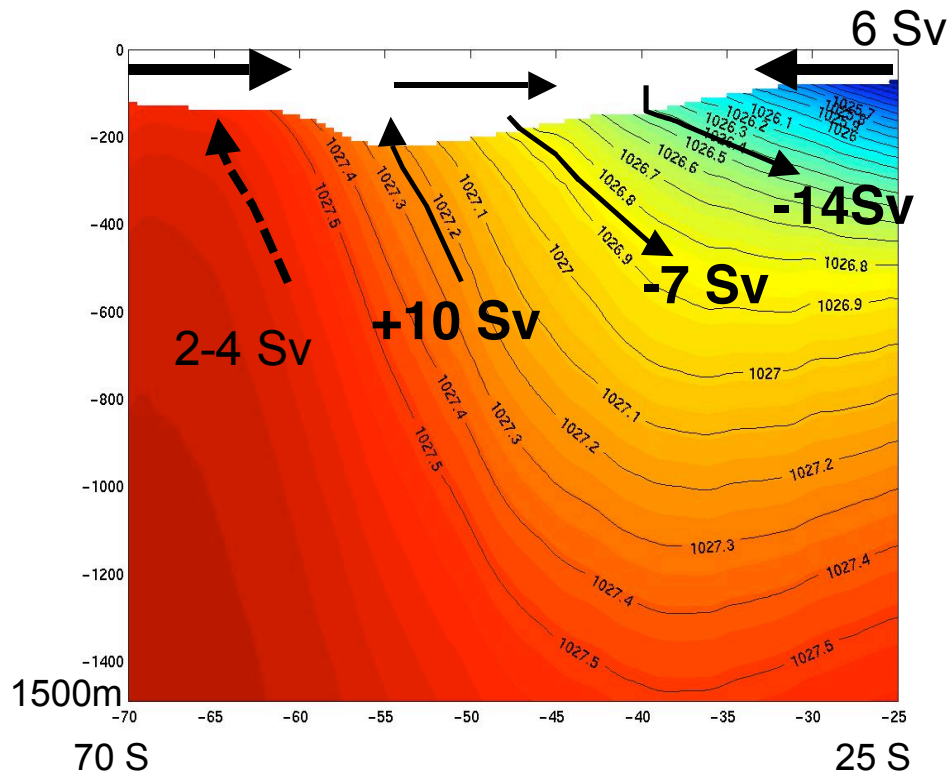
Indian (110-130 E)



Pacific (120 130 W)



# Summary



The upper Upper Cell

20 Sv coming up + south, 20 Sv going down, creating mode waters

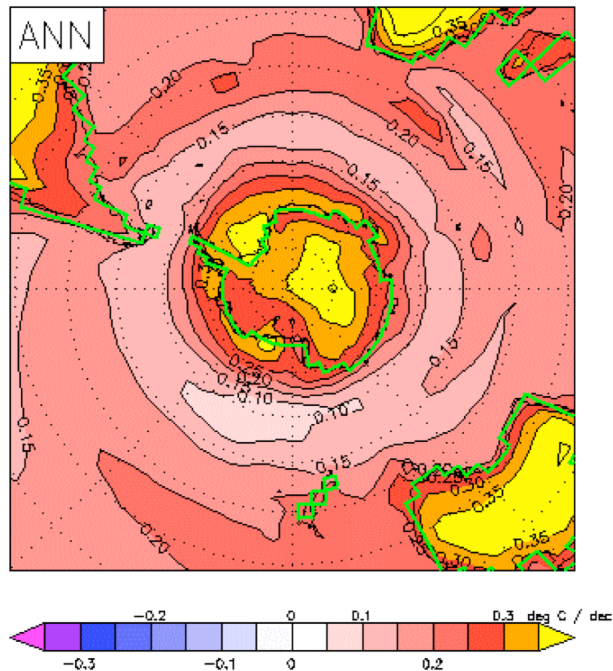
Eddies tend to cancel mean flow at high and low densities

Eddies tend to concentrate subduction at fronts

Sallee, Rintoul, Matear (2009) find carbon fluxes associated with these subduction estimates.

# ML Evolution

## Predicted trends in ST



*Predicted trends in surface temperatures over the next 100 years from a weighted average of the 20 coupled models used in IPCC AR4 (Bracegirdle et al., 2008).*

# Mixed Layer Response: SST

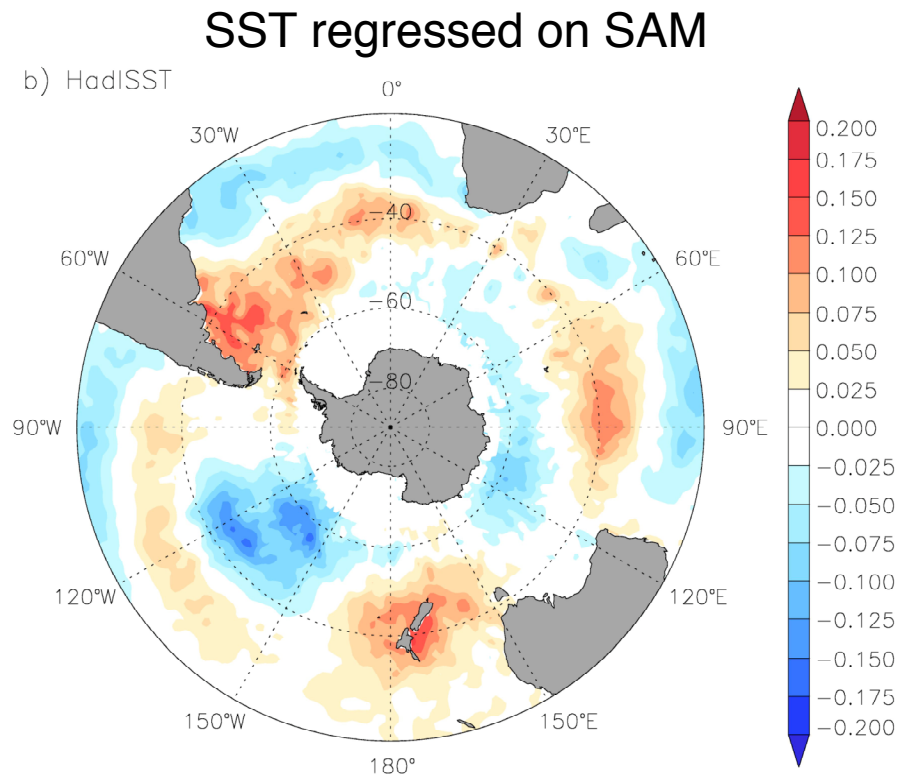
**Driven by modified atmosphere-ocean heat fluxes and Ekman flow**

## **FAST**

- SST responds to the SAM within 1 month
- Between 50-65°S cooling due to enhanced northward Ekman transport
- Between 30-45°S warming due to anomalous southward Ekman transport

## **SLOW.**

- 2-3 yr lagged warming in eddy-resolving models in south, *outweighs* initial cooling

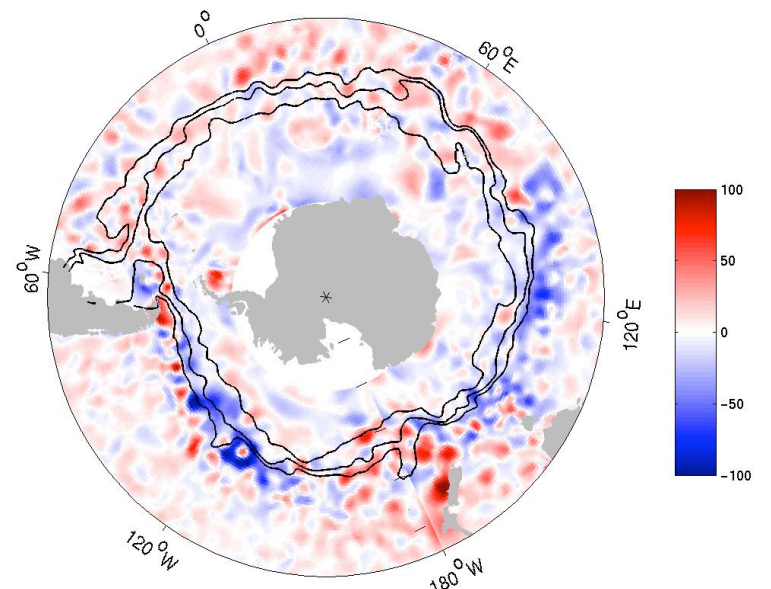
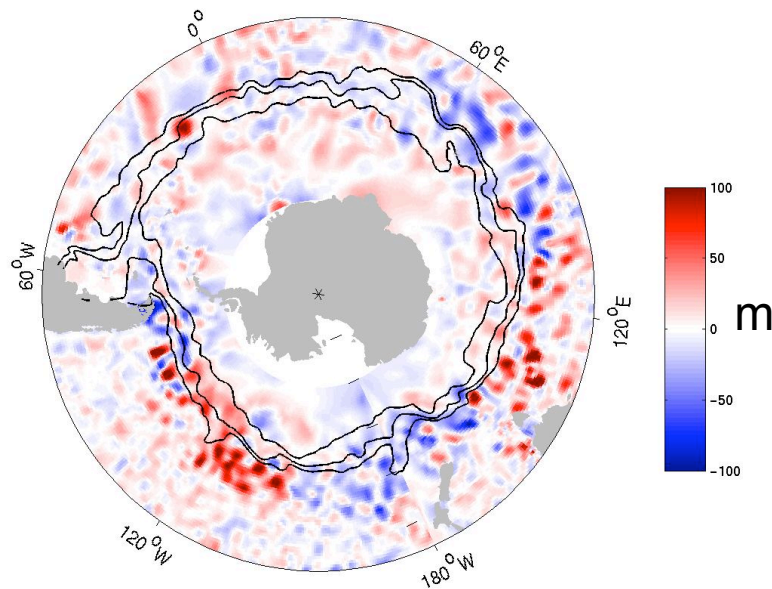


Screen et al J. Climate 2009

# Mixed Layer Response: MLD

SAM+

SAM-



MLD Trends may be important for subduction, productivity

Sallee et al. 2009



# Questions

## **Is there an “oceanic SAM”?**

- Most variability is associated with meandering fronts
- Most meandering is a response to SAM or ENSO
- Internal EKE variations are  $\sim$  EKE variations forced by wind

## **Is overturning adiabatic or diabatic below the ML?**

- Nonlocal fluxes
- Nonlinear EOS
- Highly intermittent processes

## **How are natural carbon, warm deep water brought to the surface?**

- Ekman upwelling
- Recirculation in the TL and ML
- Upper Cell eddy-driven upwelling

## **What do we need a SOOS for?**

- Freshwater fluxes
- Acidification