

The NAOSIM Data Assimilation System (NAOSIM-DAS)

Project aims and first obstacles

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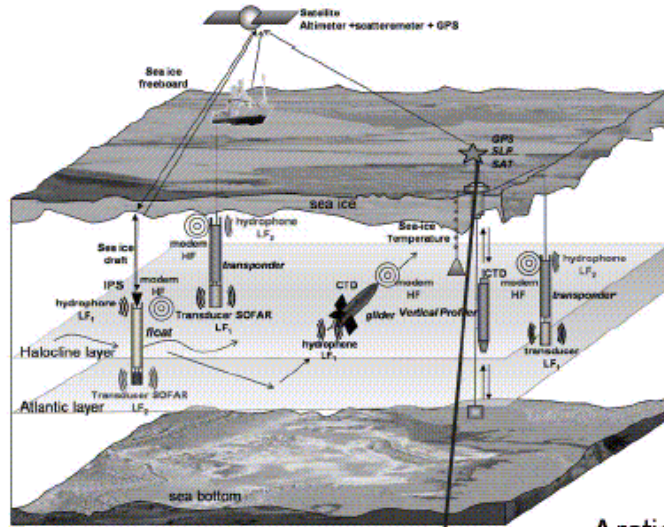
FastOpt, Hamburg

OUTLINE

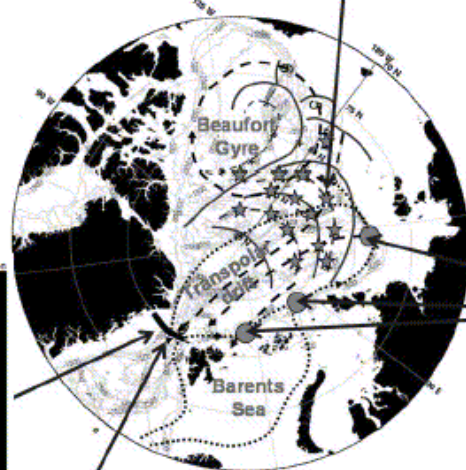
- Brief description of the EU-Project DAMOCLES
- The role of NAOSIM in the project
- Presentation of a finite-difference experiment with NAOSIM

Summary of the technological objective:

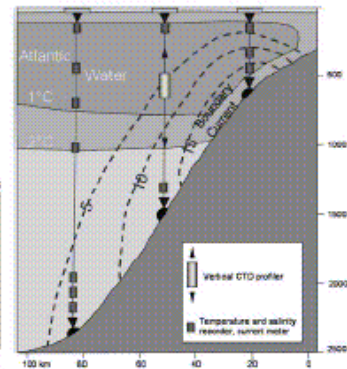
The main technological objective of DAMOCLES is to develop a prototype for an Arctic Ocean Observing System including major innovations and breakthrough in High Technology instrumentation ...



Arctic Ocean
Central Basin

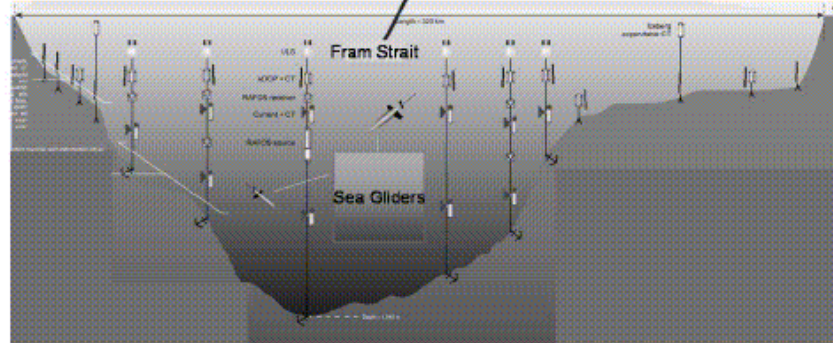
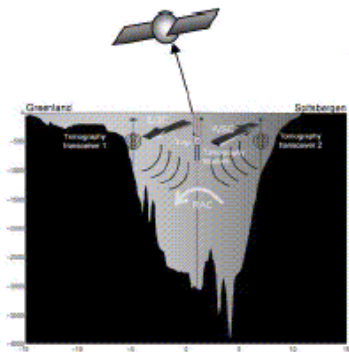


Eurasian
Continental Slope



(Selected topics)

- Satellite radar altimetry, Scatterometers, passive microwave radiometers, SAR imagery
- Sea-Gliders measuring 1000s of slanted profiles of T and S along transects between ITPs and moorings equipped with acoustic transponders
- Neutrally buoyant floats drifting at constant depth and equipped with Upward Looking Sonars to measure Sea-Ice draft
- Ice Tethered Platforms equipped with vertical CTD profilers for taking daily profiles of T and S versus depth
- Acoustic Doppler profilers measuring vertical profiles of horizontal currents
- Tomography for measuring T along vertical sections



Core theme 4: Integration and data assimilation in large-scale modelling and forecasting

Aims to integrate observations and modelling by combining output from the observation- and process-oriented themes 1-3 (and IPY and non-DAMOCLES observations) with dedicated regional and global scale numerical modelling.

A hierarchy of global coupled atmosphere-sea ice-ocean (AOGCM, BCM), regional AOCGM (RCAO, ORCM, HIRLAM/HIROMB), regional coupled sea ice-ocean models (TOPAZ, NAOSIM) to 1D vertical column models, will be used to

- Calculate model sensitivities (in case of NAOSIM with the adjoint ADNAOSIM), and to improve the models themselves
- Quantify predictability with the help of ensemble runs and data assimilation
 1. ...
 2. ...
 3. ...
 4. NAOSIM-DAS: Improved prediction skill is estimated by comparing the skill in a period with data assimilation with a the skill within the same period without data assimilation
- The overall goal is to produce a set of consistent atmospheric, oceanic and sea ice fields (analysed fields) for the DAMOCLES period

Variational Data Assimilation

Notation:

\mathbf{s} : state vector

(ocean: $u', v', s, \text{tpot}, \psi$; ice: $h, a, \text{age}, \text{hsn}$)

t : time

\mathbf{d} : vector of observations

σ : vector observational uncertainties

Principle:

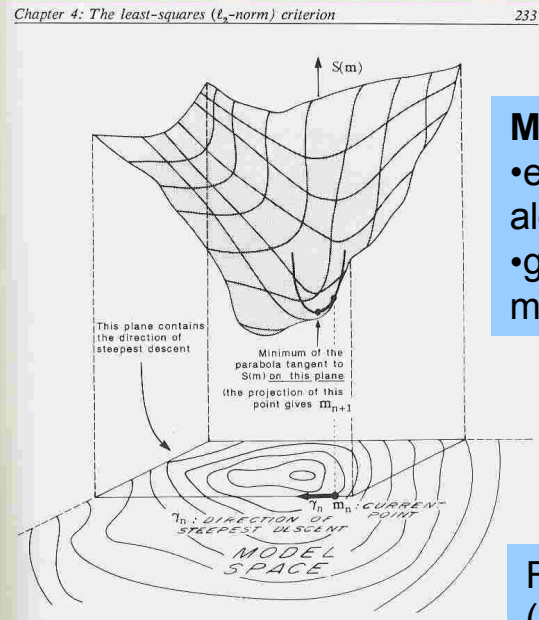
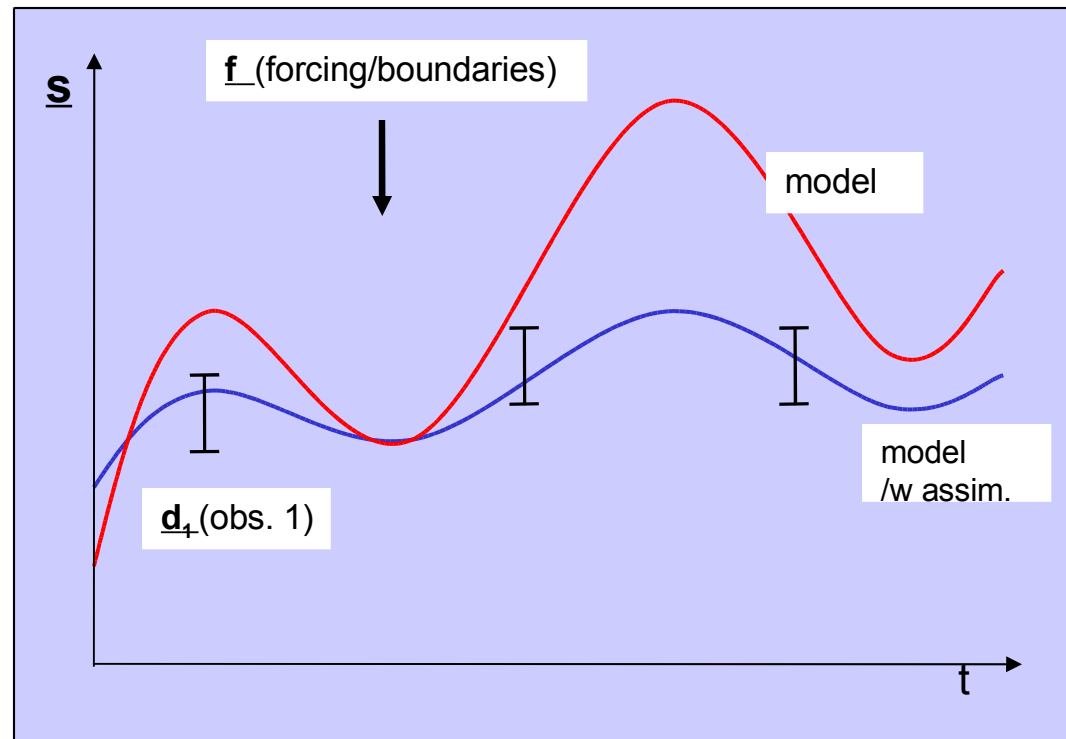
- define vector of control variables \mathbf{x} , e.g.,
 - initial state (\mathbf{s}_0)
 - forcing/boundary conditions (\mathbf{f})
 - internal model parameters (\mathbf{p})
- define quality of fit by cost function:

$$J(\mathbf{x}) = 1/2 \sum (\mathbf{d}_i - \mathbf{s}_i(\mathbf{x}))^2 / \sigma_i^2$$

- minimise $J(\mathbf{x})$ by variation of \mathbf{x}

Remarks:

- can handle **any** observation that can be computed from the model state
- in numerical weather prediction (NWP) 4DVar usually variation of initial state (\mathbf{s}_0) only
- can also include uncertainties from model error and correlated uncertainties
- can include more constraints in J , next slide...



Minimisation:

- efficient minimisation algorithms use gradient of J
- gradient provided by adjoint model

Figure: Tarantola (1997)

The ECCO costfunction – observational elements

$$\begin{aligned}
 \mathcal{J} = & (\bar{\eta} - \bar{\eta}_{TP})^t \mathbf{W}_{\text{geoid}} (\bar{\eta} - \bar{\eta}_{TP}) && \text{TOPEX absolute SSH} \\
 & + (\eta - \eta'_{TP})^t \mathbf{W}_{\text{TP}} (\eta - \eta'_{TP}) && \text{TOPEX SSH anomalies} \\
 & + (\eta - \eta'_{ERS})^t \mathbf{W}_{\text{ERS}} (\eta - \eta'_{ERS}) && \text{ERS SSH anomalies} \\
 & + (\bar{T}_{surf} - \bar{T}_{Reyn})^t \mathbf{W}_{\text{SST}} (\bar{T}_{surf} - \bar{T}_{Reyn}) && \text{Reynolds SST} \\
 & + (\bar{T} - \bar{T}_{Lev})^t \mathbf{W}_{\text{TLev}} (\bar{T} - \bar{T}_{Lev}) && \text{Levitus clim.} \\
 & + (\bar{S} - \bar{S}_{Lev})^t \mathbf{W}_{\text{SLev}} (\bar{S} - \bar{S}_{Lev}) && \text{Levitus clim.} \\
 & + (\tau_x - \tau_{x,NCEP})^t \mathbf{W}_{\tau_x} (\tau_x - \tau_{x,NCEP}) && \text{zonal wind stress} \\
 & + (\tau_y - \tau_{y,NCEP})^t \mathbf{W}_{\tau_y} (\tau_y - \tau_{y,NCEP}) && \text{merid. wind stress} \\
 & + (H_Q - H_{Q,NCEP})^t \mathbf{W}_{H_Q} (H_Q - H_{Q,NCEP}) && \text{NCEP heat flux} \\
 & + (H_F - H_{F,NCEP})^t \mathbf{W}_{H_F} (H_F - H_{F,NCEP}) && \text{NCEP freshwater flux}
 \end{aligned}$$

Currently added:

- Jason-1 altimetry (sea surface height)
- WOCE hydrography, XBT, TAO buoys
- PALACE/ARGO tracer profiles and drift velocities
- surface drifter velocities
- NSCAT/QuickScat surface wind stress fields
- TRMM/TMI tropical surface temperature fields

ECCO costfunction
courtesy: Patrick Heimbach (MIT)

setup for assimilation into MIT
ocean general circulation model

TAF

Transformation of Algorithms in Fortran

- Source-to-source translator for Fortran-77/90/95
- forward and reverse mode
- scalar and vector mode
- full and pure mode
- efficient Hessian code by applying TAF twice (e.g. forward over reverse)
- command line program with many options
- TAF-Directives are Fortran comments
- extensive and complex code analyses (similar to optimising compilers)
- generated code is structured and well readable

TAF

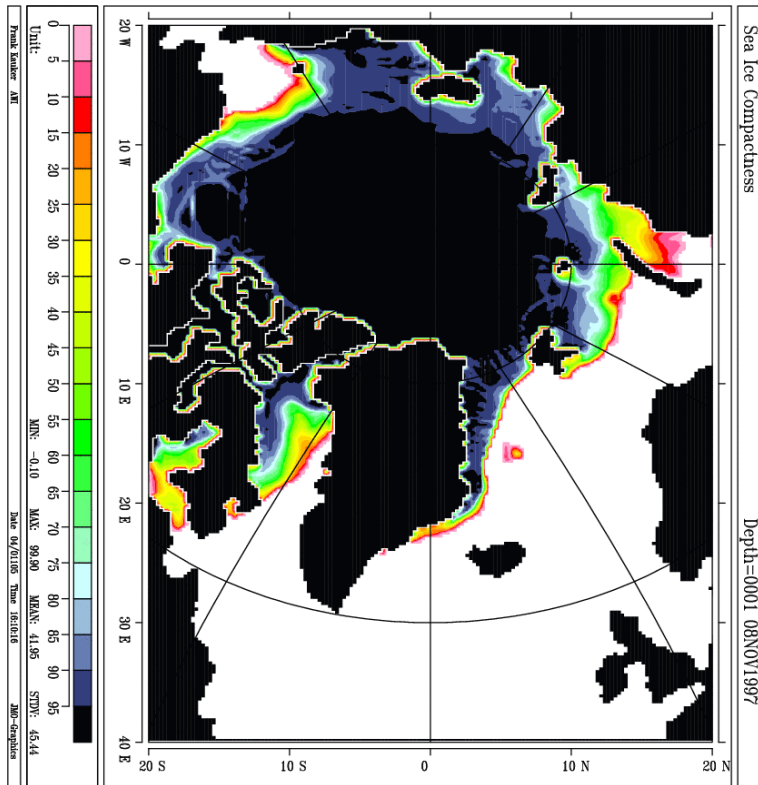
More features

- Generation of flexible store/read scheme for required values triggered by TAF init and store directives
- Generation of simple checkpointing scheme (Griewank, 1992) triggered by combination of TAF init and store directives
- Generation of efficient adjoint (Christianson, 1996, 1998) for converging iterations triggered by TAF loop directive
- TAF flow directives for black-box routines, or to include user provided derivative code (exploit linearity or self-adjointness, MPI wrappers, etc...)
- Automatic Sparsity Detection
- Basic support for MPI and OpenMP
- supports interrupting and restarting adjoint ('divided adjoint')

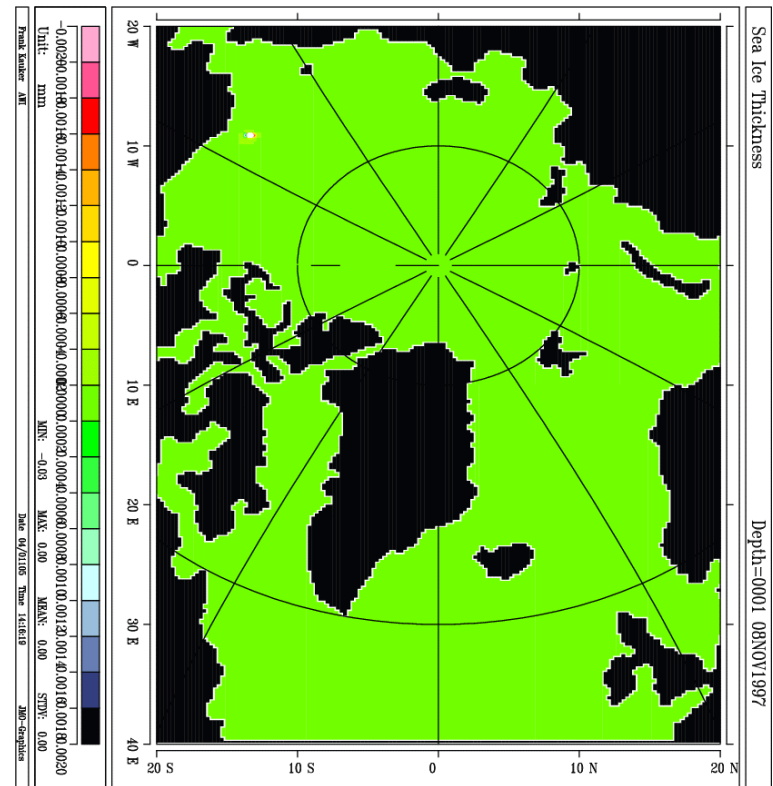
Finite-difference experiment (proposed by D. Menemenlis, JPL)

- Disturb the 2m-temperature at ONE grid point with marginal sea ice by +0.01K and -0.01K for 7days.
- Is the response weakly or strongly non-linear?

Sea ice concentration at the start (11/8/1997)

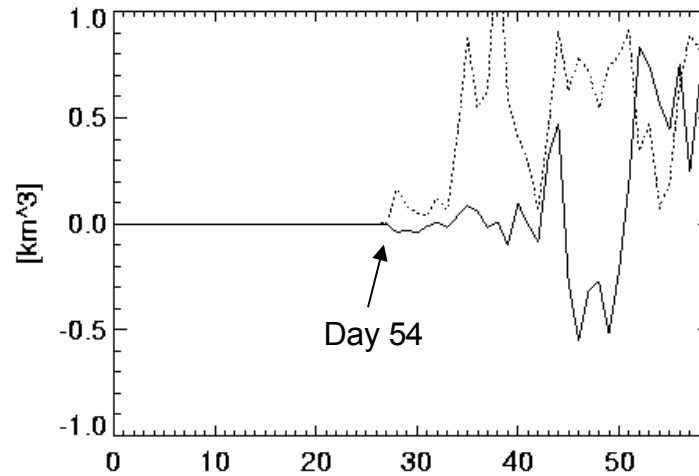
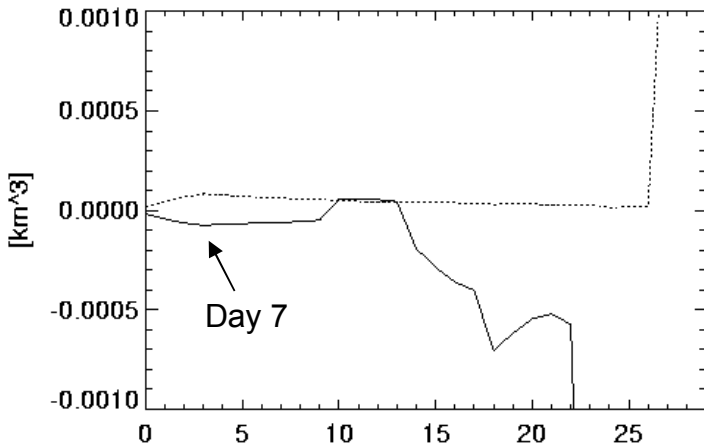


First response after 2 days in sea ice thickness



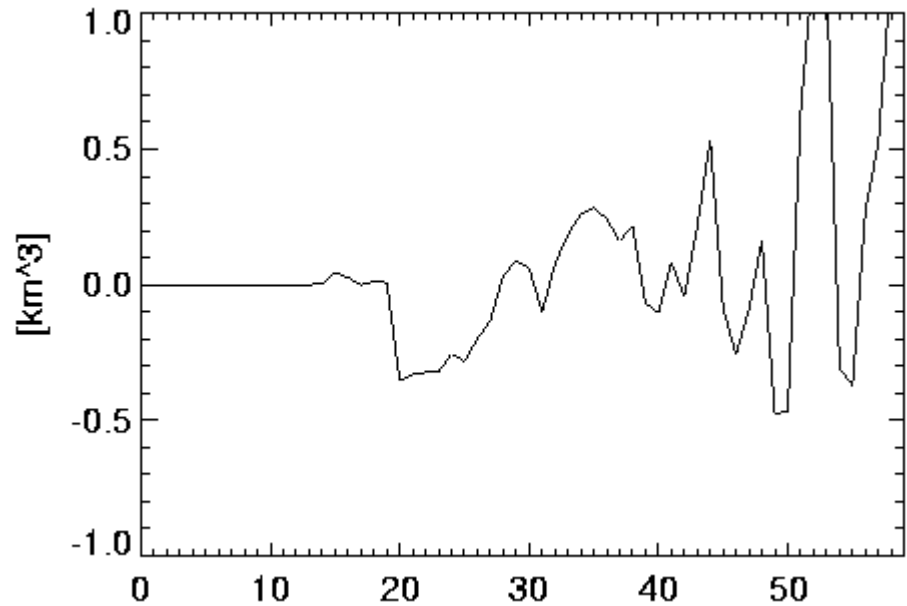
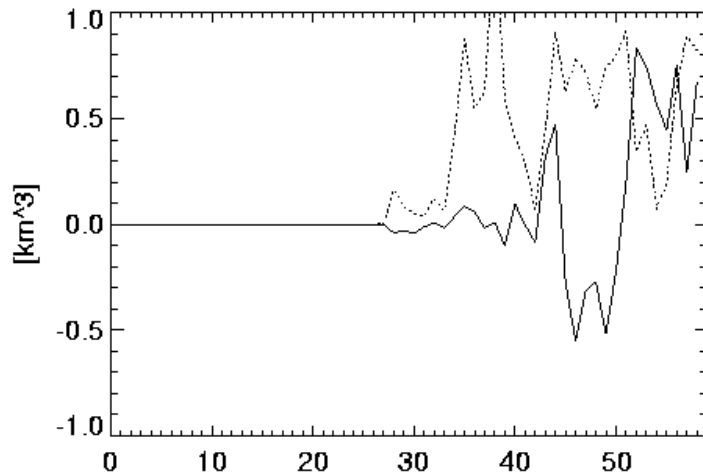
Finite-difference experiment:

- Look at some integral quantity: sea ice volume in the Arctic [km^3]
- Difference +0.01K-ctrl (straight) and -0.01K-ctrl (dashed)



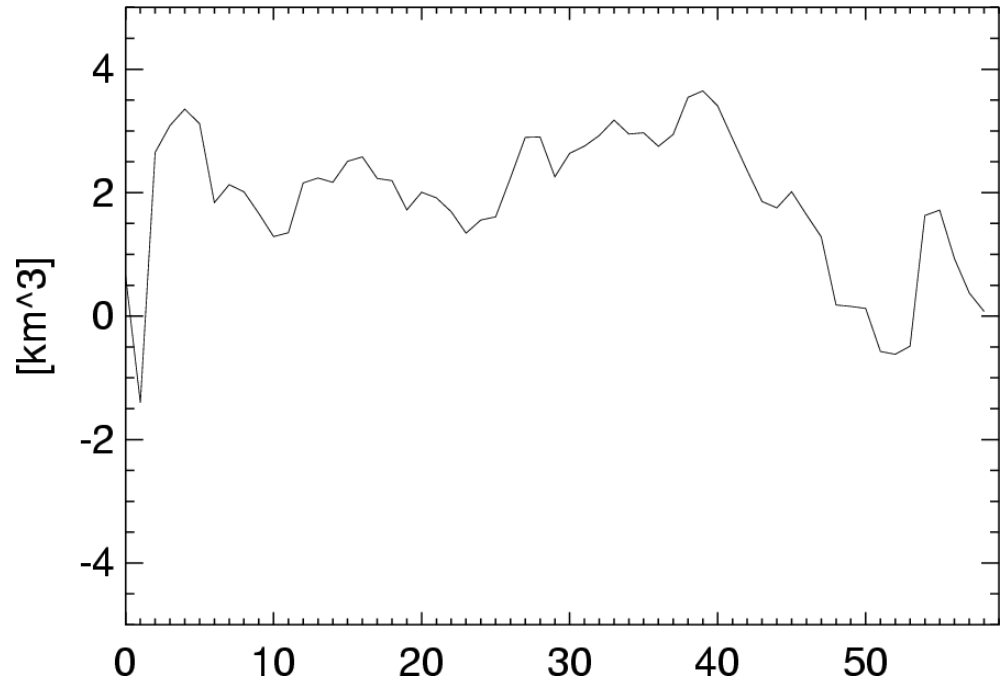
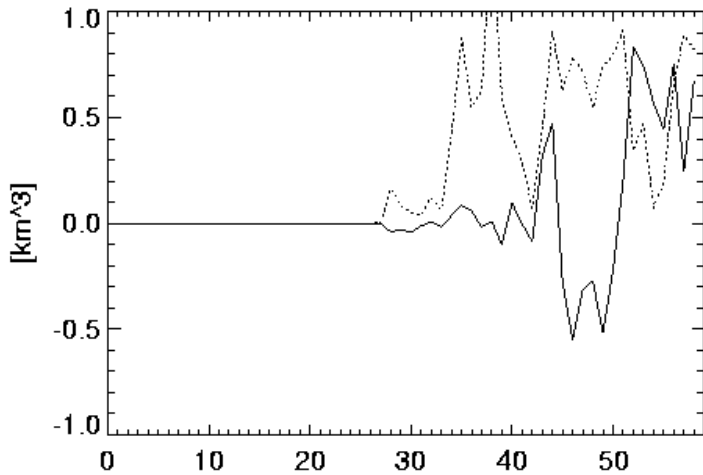
Finite-difference experiment:

- Look at some integral quantity: sea ice volume in the Arctic [km^3] (mean 20000 to 30000 km^3)
- $\text{amasmin}=1.e-6\text{m} \rightarrow 1.e-8\text{m}$
(if ice thickness lower amasmin free drift solution is applied)



Finite-difference experiment:

- Look at some integral quantity: sea ice volume in the Arctic [km^3]
- $v_{\text{max}}=1.e-4\text{m/s} \rightarrow 5.e-5\text{m/s}$
(if the global maximum of the drift change in the iteration is lower v_{max} stop iteration)



Questions finite-difference experiment:

- Strong non-linear behaviour due to the numerical formulation or the real physics?
- Hibler and Bryan (1986) stress coupling?
- Is EVP more well behaved?