

Figure 1. Redrawn from Bard *et al.* (1990), with time reversed to increase from left to right. U curve is estimated sea level with respect to today using Barbados terraces and U/Th dates. L curve is identical except that ^{14}C dates are used. U/Th dates are believed more accurate, but no other variables are specified relative to ^{14}C . A 7 meter adjustment for the greater height of the interglacial was made by Bard *et al.* (1990).

Sea level has been rising globally on average since the glaciers started melting about 20,000 years ago.

From Douglas et al. 2001

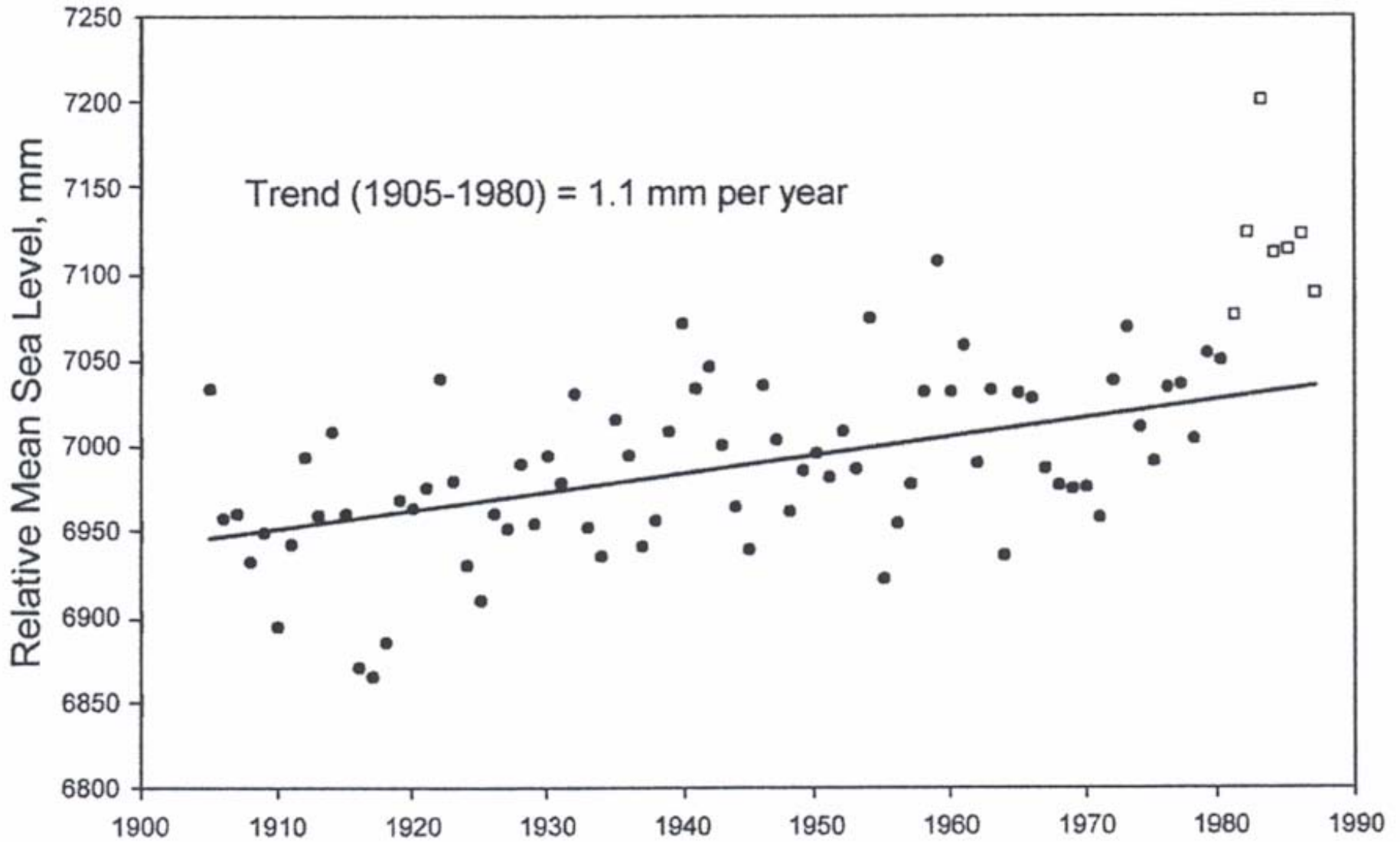
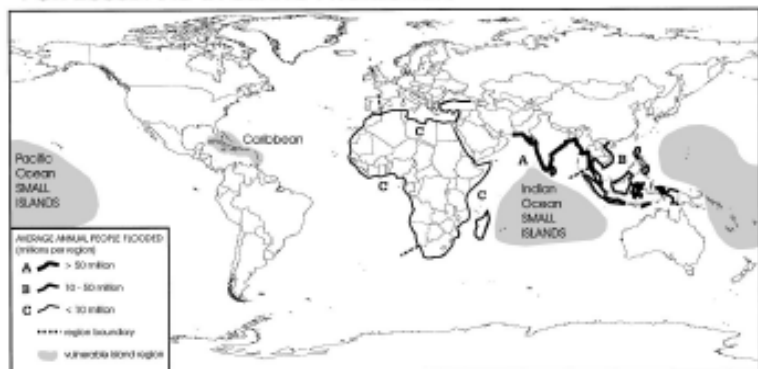


Figure 3.5 Annual mean relative sea level at Buenos Aires.

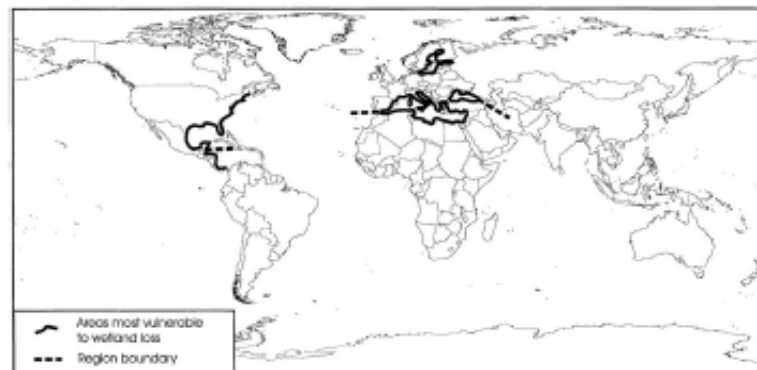
AREAS VULNERABLE TO COASTAL FLOODING FOR 2080s AND EVOLVING PROTECTION



(a) Flood impacts

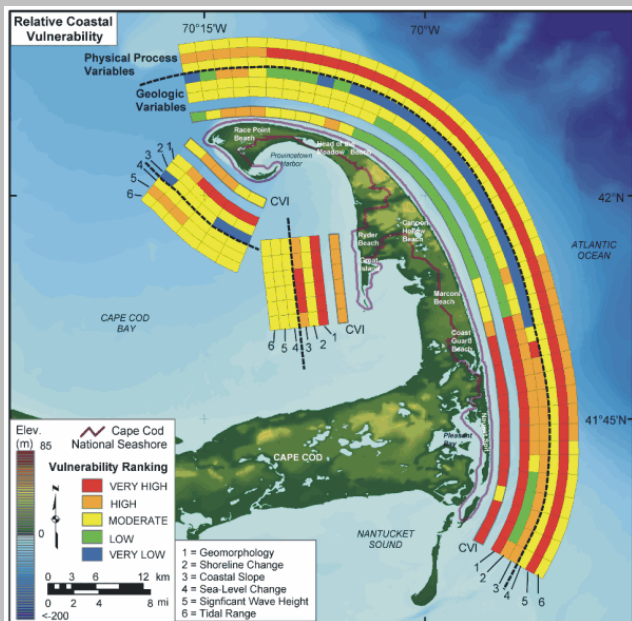
Fig. 14. Regional implications of sea-level rise — the regions most affected by flood impacts given the HadCM2 (mean) scenario for the 2080s.

AREAS MOST VULNERABLE TO COASTAL WETLAND LOSS



(b) Wetland loss

Fig. 17. Regional implications of sea-level rise — the regions where wetland losses may exceed 65% due to the HadCM2 (mean) scenario by the 2080s.



USGS



USGS, WHOI

From Douglas et al., 2001



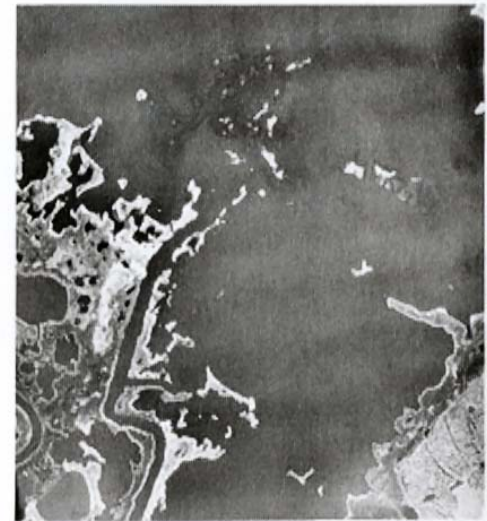
1938



1957

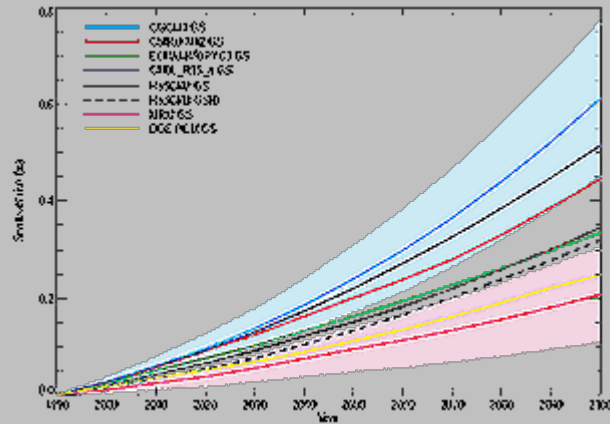


1972



1988

Figure 8.21 Progressive drowning and loss of coastal marshes in Blackwater National Wildlife Refuge near Cambridge, Maryland, in response to a high relative sea level rise.



From Church et al. (2001)

Figure 11.1: Global average sea level changes from thermal expansion simulated in AOGCM experiments with historical concentrations of greenhouse gases in the 20th century, then following the IS92a scenario for the 21st century, including the direct effect of sulphate aerosols.

See [Tables 8.1](#) and [9.1](#) for further details of models and experiments

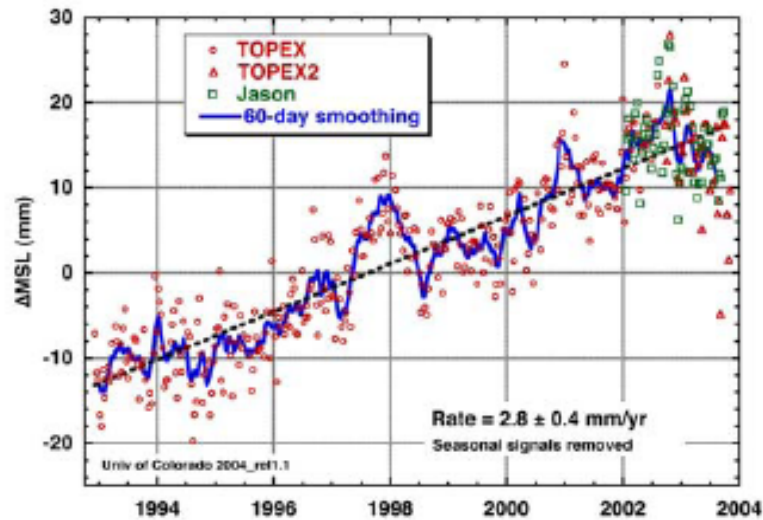


Figure 5. Global mean sea level variations every 10 days from T/P (red circles, red triangles after T/P was moved to new ground track) and Jason (green squares) and after smoothing with a 60-day boxcar filter (blue line) [Leuliette *et al.*, 2004]. No inverted barometer correction was applied to the altimeter data, and seasonal variations have been removed.

No accurate altimetry prior to 1992 .

There exist a number of published attempts at determining the global average prior to the altimeter measurements based on (1) tide gauge records; (2) temperature and salinity changes.

These have been the subject of considerable dispute, as the calculations prove very difficult.

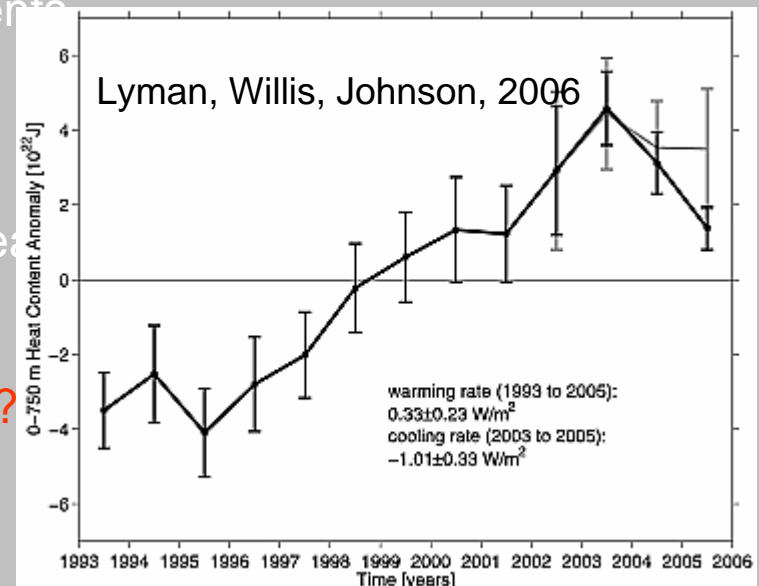
Issues pertain to:

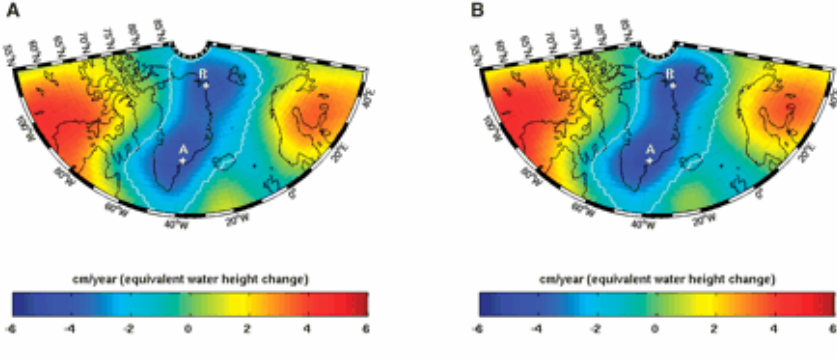
The spatial distribution of the measurements

The interpretation of density changes

Calibration (or lack of it) in the various measurements

Is sea level rising? Is the ocean getting warmer?
where, and by how much, and why?





Chen, Wilson, Tapley, Science, 2006
 Satellite Gravity Measurements
 Confirm Accelerated
 Melting of Greenland Ice Sheet
 (note 3 years of data)

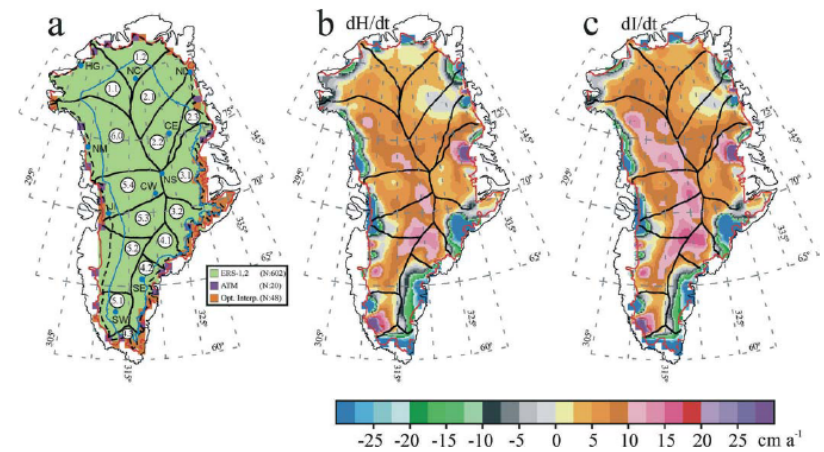


Fig. 1. Greenland. (a) Distribution of surface elevation change data by source, derived from ERS-1 and -2 radar altimetry, ATM (closest-neighbor interpolation from airborne surveys), and obtained by optimal interpolation: ice terminus of coterminous ice sheet (red), equilibrium line (black dashes), 2000 m elevation contour (blue), drainage divides (black), drainage system designation (number in circles), and location of $H(t)$ series depicted in Figure 3a (labeled blue full circles). (b) Distribution of elevation change (dH/dt). (c) Distribution of ice-thickness change (dI/dt).

Zwally et al., J. Glaciology 2005. “The Greenland ice sheet ...[has] a small overall mass gain, $+11 \pm 3 \text{ Gt/a}$.”

Are the ice sheets growing or shrinking?

From temperature

From freshwater (salinity change)

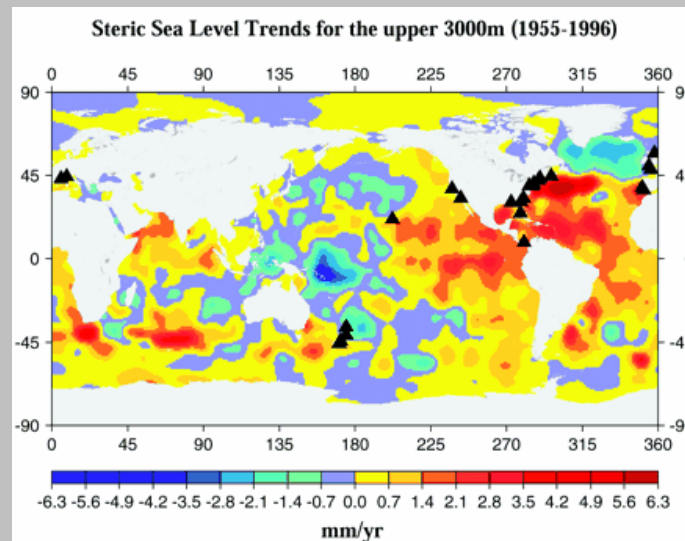
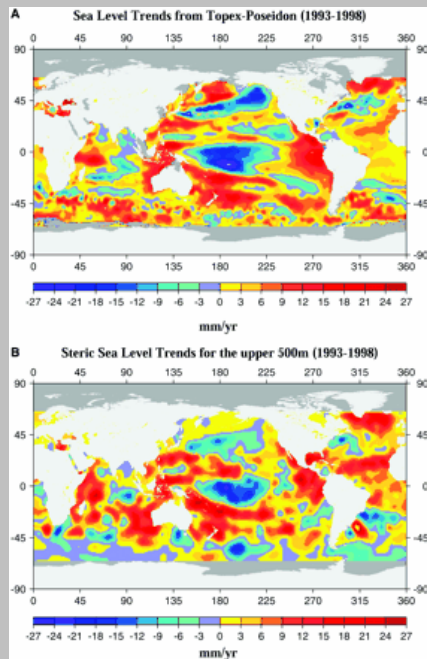


Thermosteric Change [mm/y]	Integr. Depth [m]	Freshwater Change [mm/y]	Total Change [mm/y]	Interval	Comment
Antonov et al. (2005)					
≈ 0.33	700			1955–2003	hydrography
Antonov et al. (2002)					
	3000	1.3 ± 0.5		1957–1994	hydrography
Willis et al. (2004)					
1.6 ± 0.3	750			1993–2003	XBT & altimetry
Carton et al. (2005)					
2.3 ± 0.8	1000			1958–2001	altimetry & hydrography
Plag (2006)					
		0.49 ± 0.12	1.06 ± 0.75	1950–1998	tide gauges & reworked hydrography*
Ishii et al. (2006)					
0.31 ± 0.071	700	1.44 ± 0.36	1.75 ± 0.36	1995–2003	hydrography ⁺
Miller and Douglas (2004)					
			≈ 2	20th century	tide gauges
Hansen et al. (2005)					
$0.85 \pm 0.15 \text{ W/m}^2 =$ 1.1 ± 0.2				1993–2003	coupled model alone

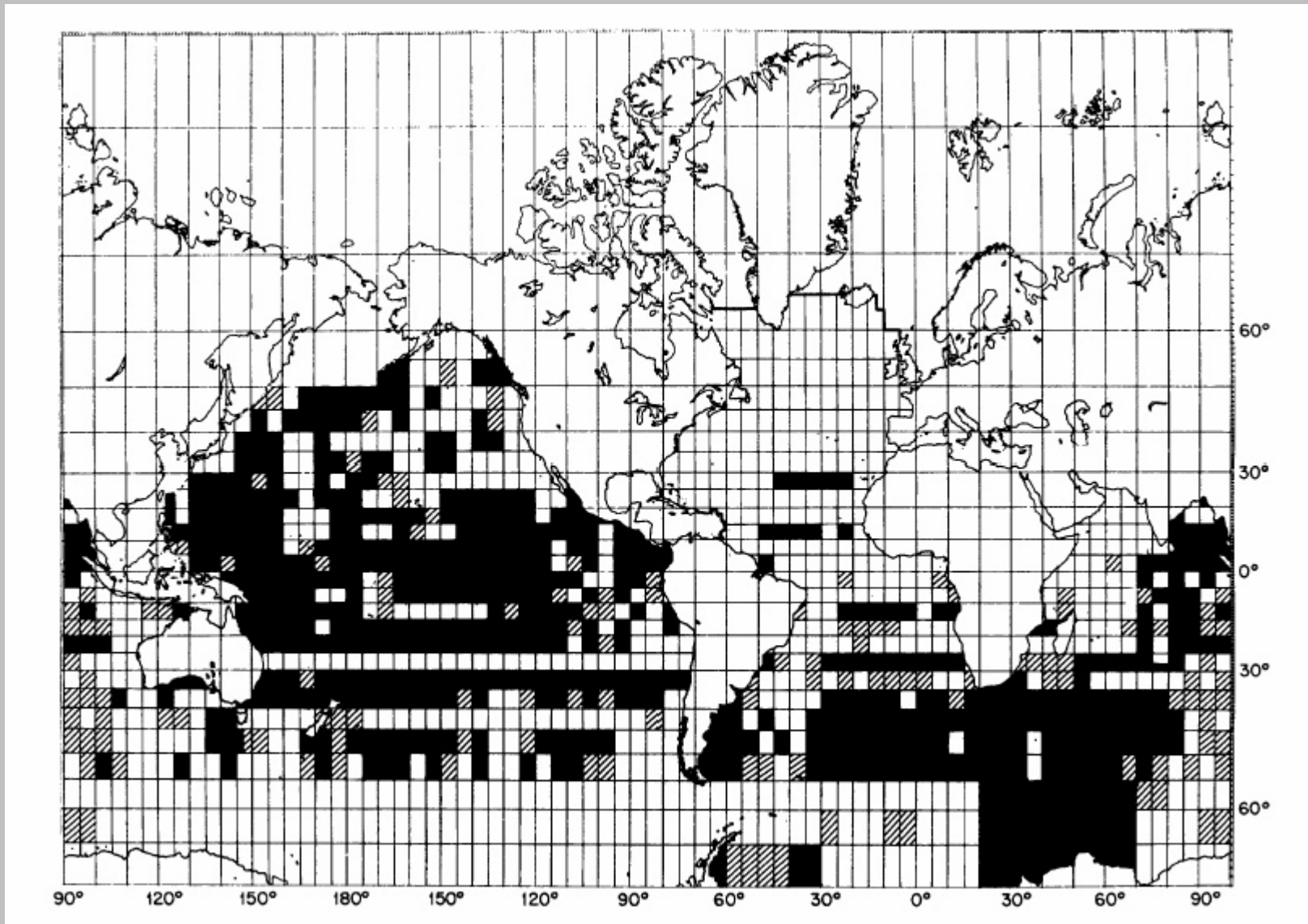
Sea Level Rise During Past 40 Years Determined from Satellite and in Situ Observations

Cecile Cabanes, Anny Cazenave, Christian Le Provost (2001)

“The 3.2 ± 0.2 millimeter per year global mean sea level rise observed by the Topex/Poseidon satellite over 1993-98 is fully explained by thermal expansion of the oceans. For the period 1955-96, sea level rise derived from tide gauge data agrees well with thermal expansion computed at the same locations. However, we find that subsampling the thermosteric sea level at usual tide gauge positions leads to a thermosteric sea level rise twice as large as the "true" global mean. As a possible consequence, the 20th century sea level rise estimated from tide gauge records may have been overestimated.”

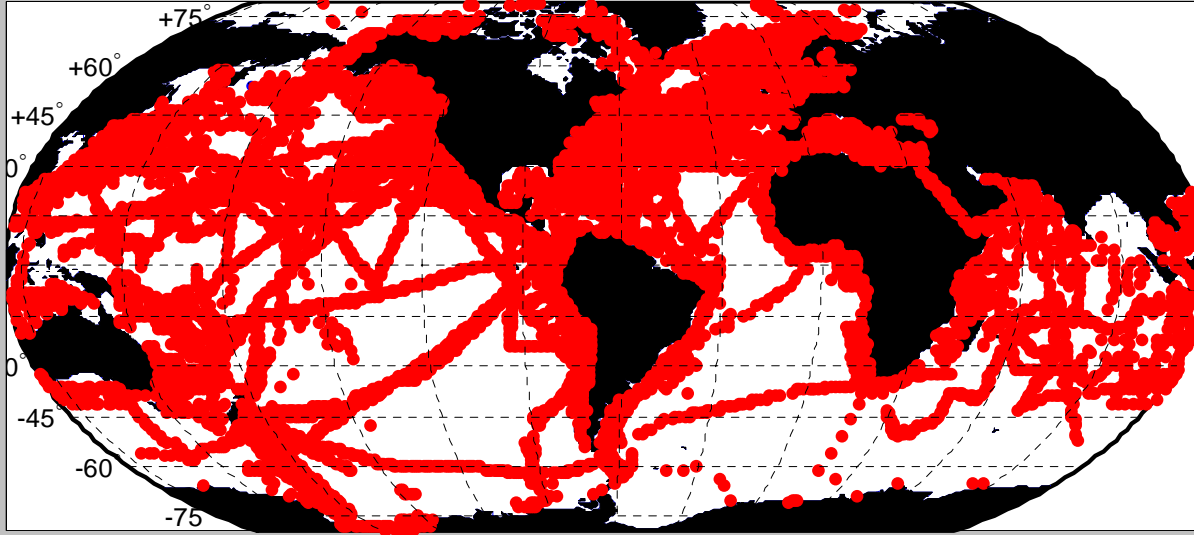


But, Miller and Douglas, Nature, 2004, show the apparent bias is due to smoothing in the Levitus et al., hydrography



As of 1977, white squares had at least one acceptable deep station, Hatched had at least one intermediate depth station, and black had no acceptable station.

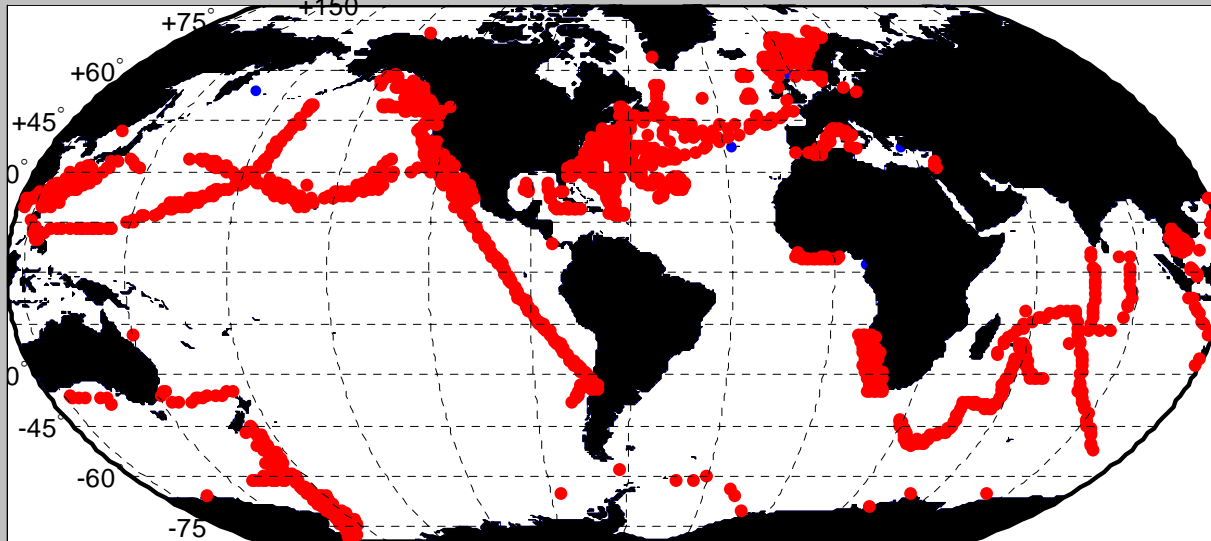
sampling in 1960 WOA
+120° +150° -180° -150° -120° - 90° - 60° - 30° 0° + 30° + 60° + 90°



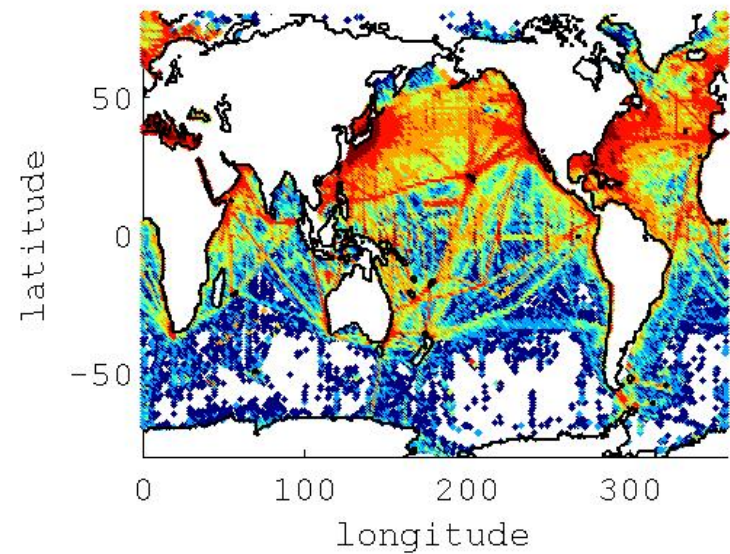
mainly MBTs

sampling in 2002

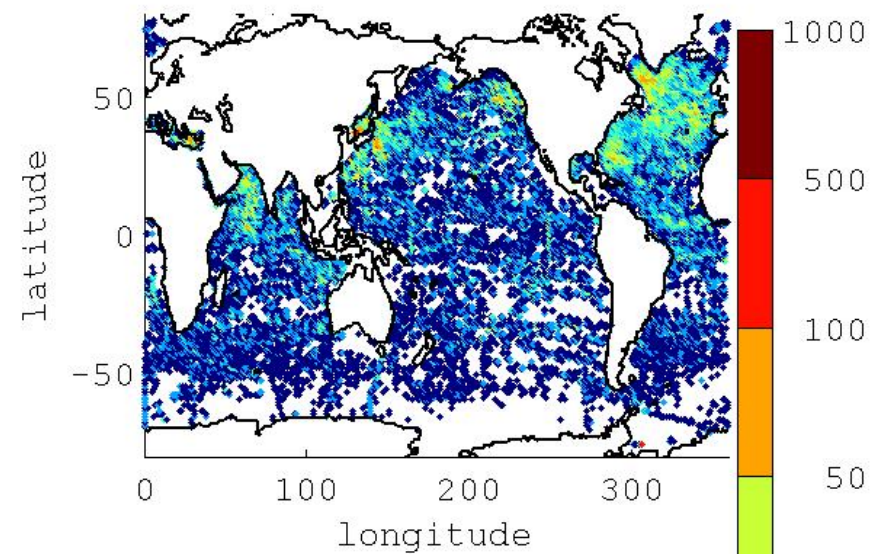
+120° +150° -180° -150° -120° - 90° - 60° - 30° 0° + 30° + 60° + 90°



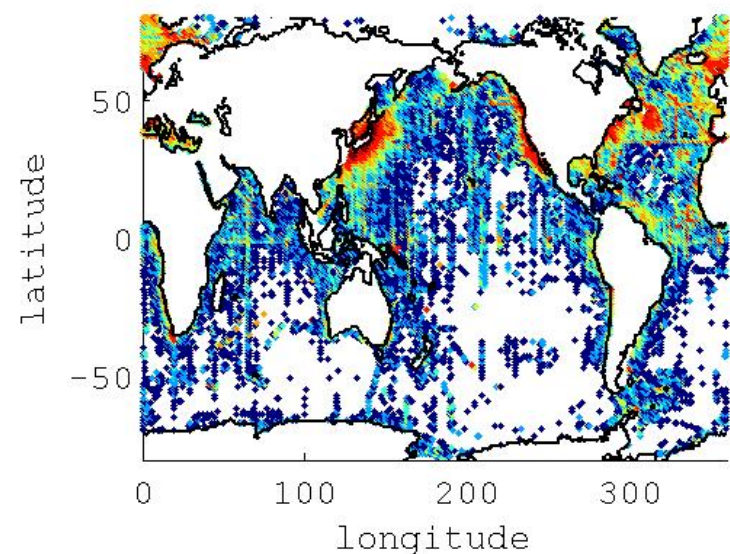
Levitus data, for T at 300m



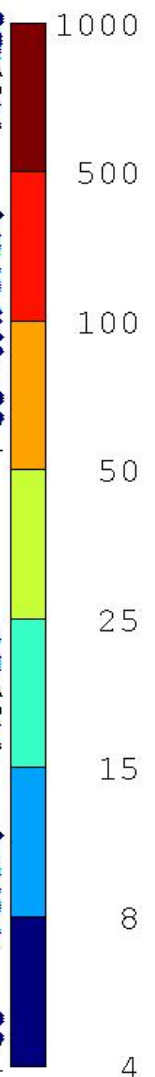
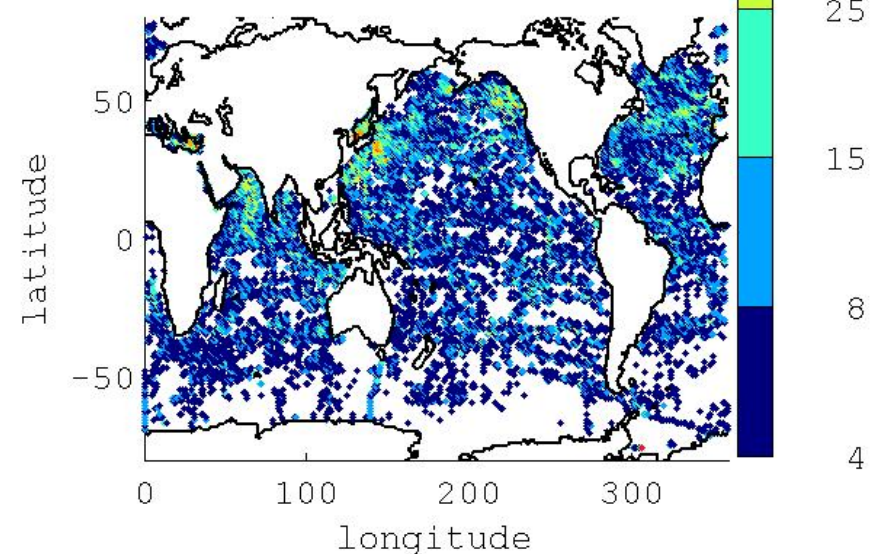
recent data, for T at 300m



Levitus data, for S at 300m



recent data, for S at 300m



More than 4 measurements in a 1 degree square. "Recent" means WOCE & later.

Munk (2004) pointed out that there was some misinterpretation of salinity changes as they influence sea level.

The equation of state is approximately linear:

$$\rho = \bar{\rho}(1 - \alpha T + \beta S)$$

Suppose the fluid is well-mixed over a depth $h(0)$ with temperature $T(0)$ and salinity $S(0)$. Then keeping the salinity fixed, and changing the temperature to $T(1) = T(0) + \Delta T$, the fluid mass does not change, and sea level shifts by the "thermosteric" amount, h_T so that

$$\begin{aligned} \bar{\rho}(1 - \alpha T(0) + \beta S(0))h(0) &= \\ \bar{\rho}(1 - \alpha T(0) - \alpha \Delta T + \beta S(0))(h(0) + h_T) & \end{aligned}$$

or,

$$h_T \approx \alpha \Delta T h(0),$$

which is the excess volume of fluid required to compensate for a uniform temperature increase of ΔT . (In a Boussinesq model, $h_T = 0$, and the mass does change.)

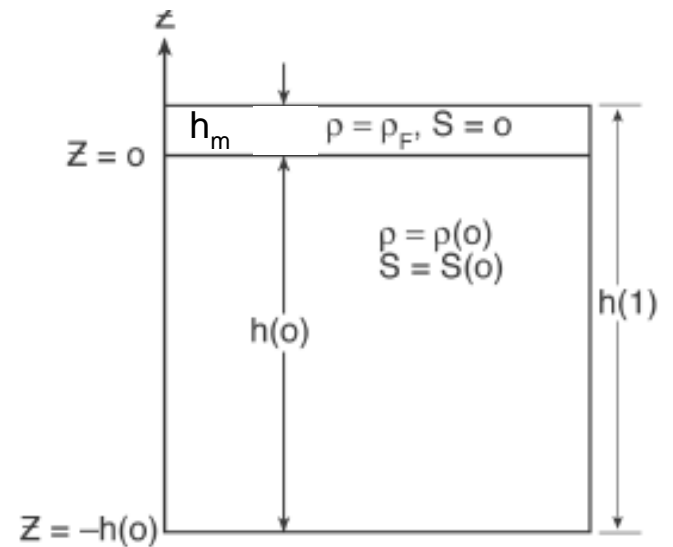
If a layer of fresh water of thickness h_m at temperature $T(0)$ is added to the fluid and then well-mixed, the salinity change is,

$$\Delta S = -S(0) \frac{h_m}{h(0)}.$$

The salt content does not change, and the new column depth is $h(0) + h_s$, where h_s is the halosteric change, determinable from

$$\begin{aligned} \bar{\rho}(1 - \alpha T(0) + \beta S(0))S(0)h(0) &= \\ \bar{\rho}(1 - \alpha T(0) + \beta S(0) + \beta \Delta S)(S(0) + \Delta S)(h(0) + h_s), & \end{aligned}$$

or



Definition sketch of a fluid of initial depth h

or

$$h_s = \beta S(0) h_m$$

$$h_m = \frac{h_s}{\beta S(0)}$$

where the "Munk multiplier" $1/\beta S(0) \approx 36$.

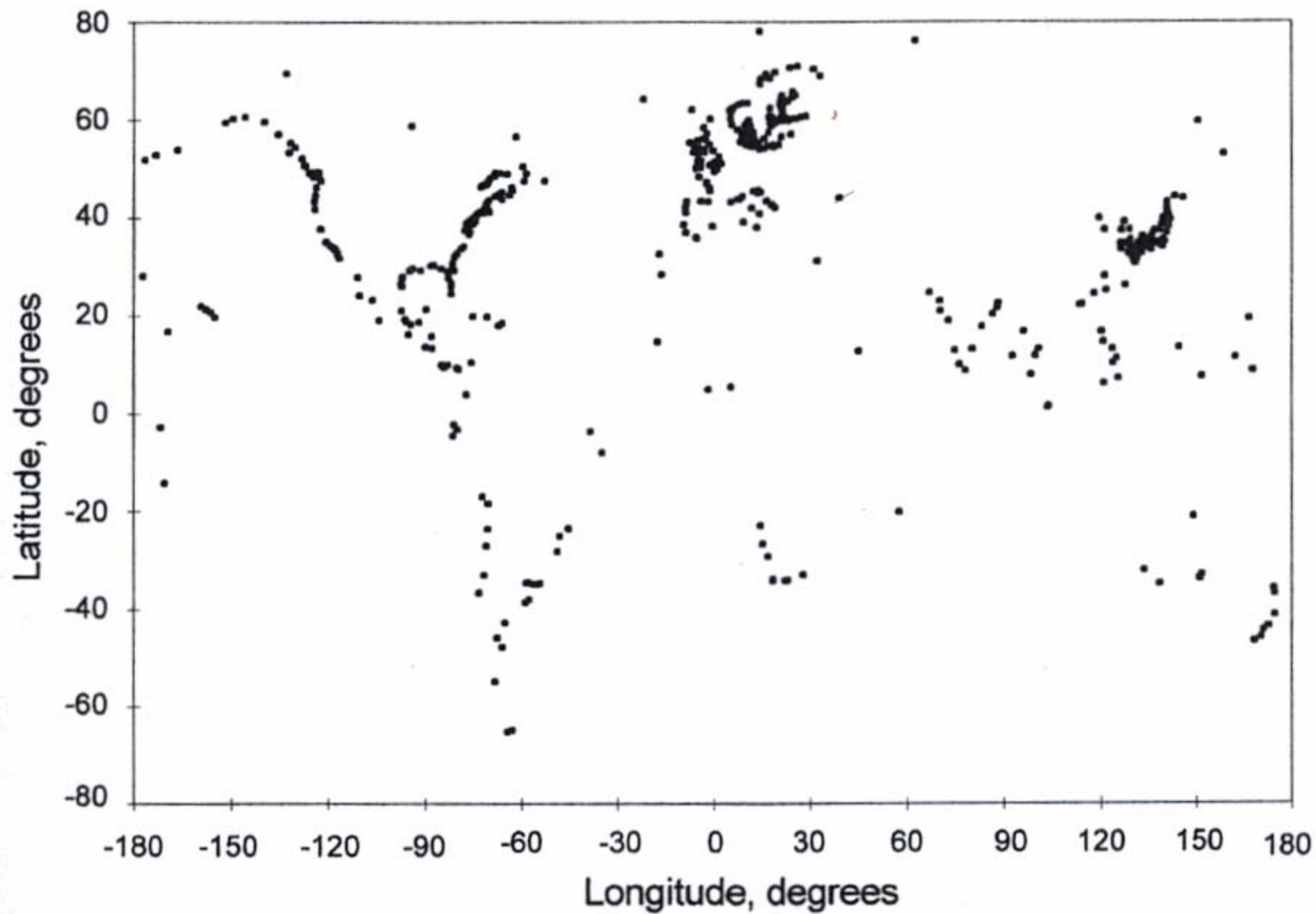
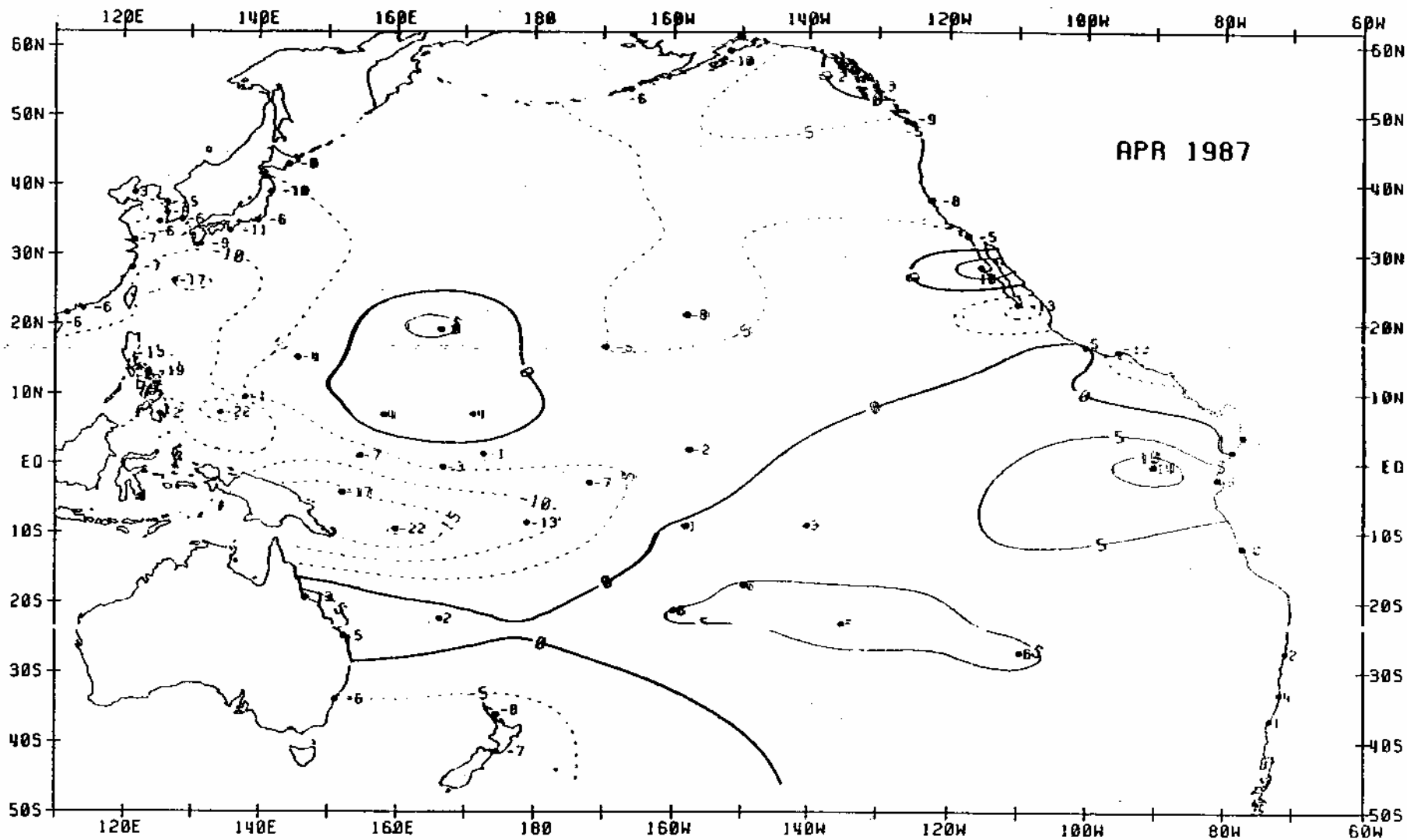


Figure 3.12 RLR tide gauge locations with records longer than 20 years.



DEVIATION OF SEA LEVEL IN APRIL 1987 FROM THE 1975 TO 1981 MEAN SEA LEVEL IN CENTIMETER.

From Douglas et. al.
2001.

Note that sealevel
appears to be falling
around Scandinavia.

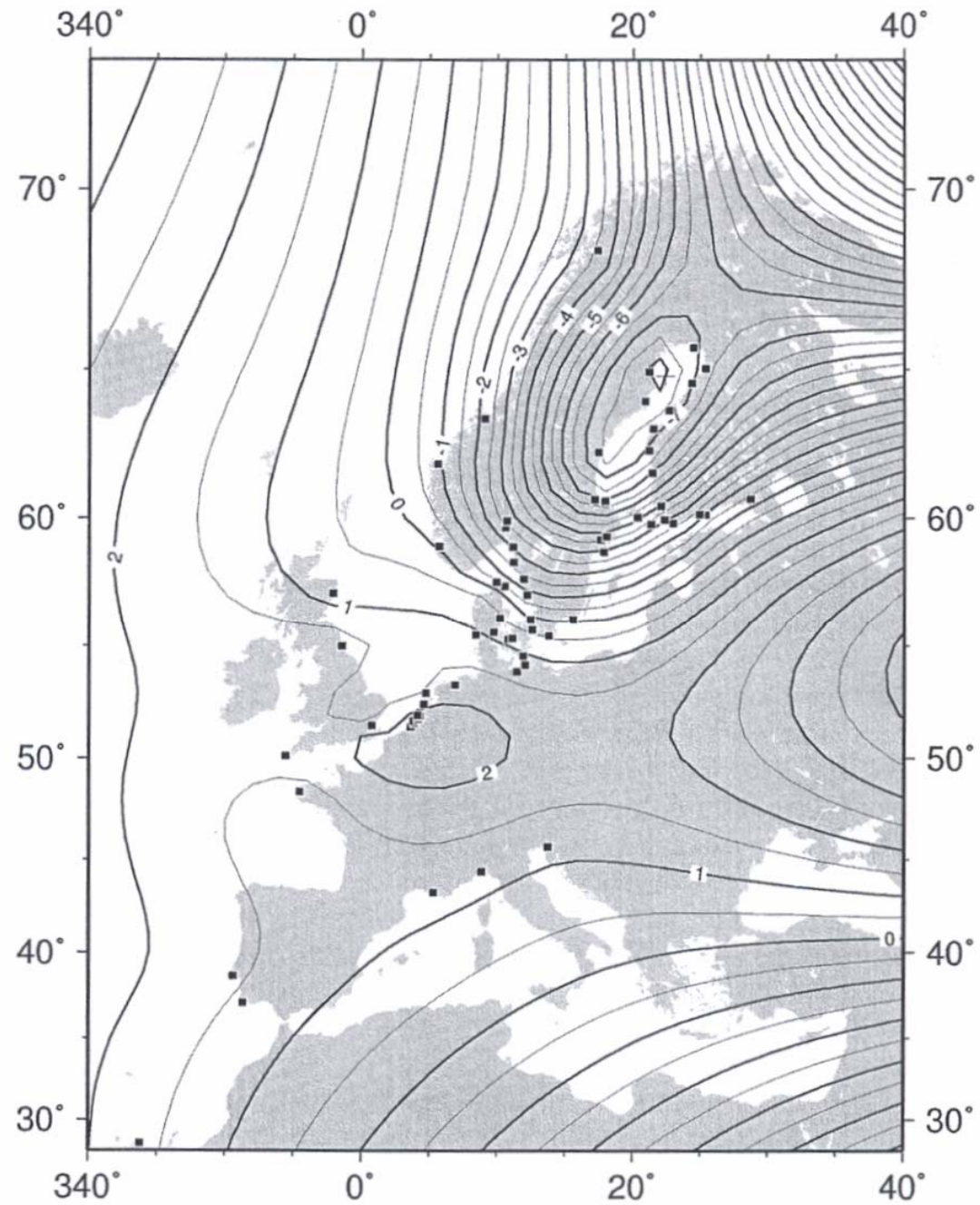


Figure 1.2 Contour plot of 20th-century European sea level trends (mm/yr).

Example of possible systematic errors in data leading to trends:

Salinity:

the technology changed from water samples run by titration to water samples run on conductivity machines which evolved (to Schleicher/Bradshaw). The definition of salinity changed (not in itself an issue). Samples were often drawn into poorly rinsed bottles; they were often stored for later measurement ashore, weeks or even months later (evaporation was found to be a problem). Seasonal and latitudinal sampling biases.

Temperature:

Main issue is probably the depth inference, which shifted from reversing thermometers to pressure gauges. Gouretski and Jahncke (2001) found in their climatology an inordinate number of samples at the nominal bottle depth. Seasonal and latitudinal sampling biases.

Tide Gauges:

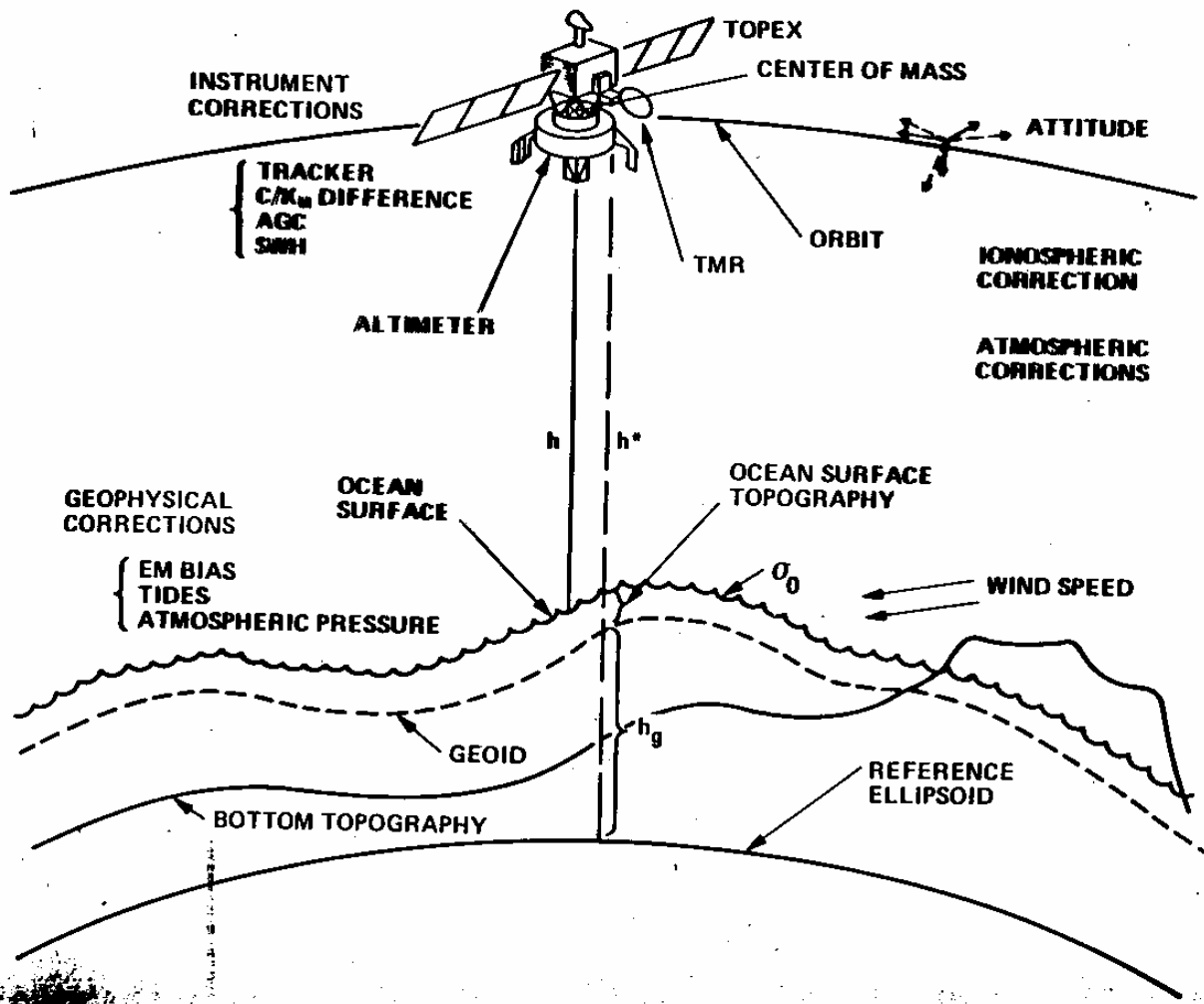
Tectonic corrections. Location shifts. Harbor construction.

Are these errors random, or systematic?

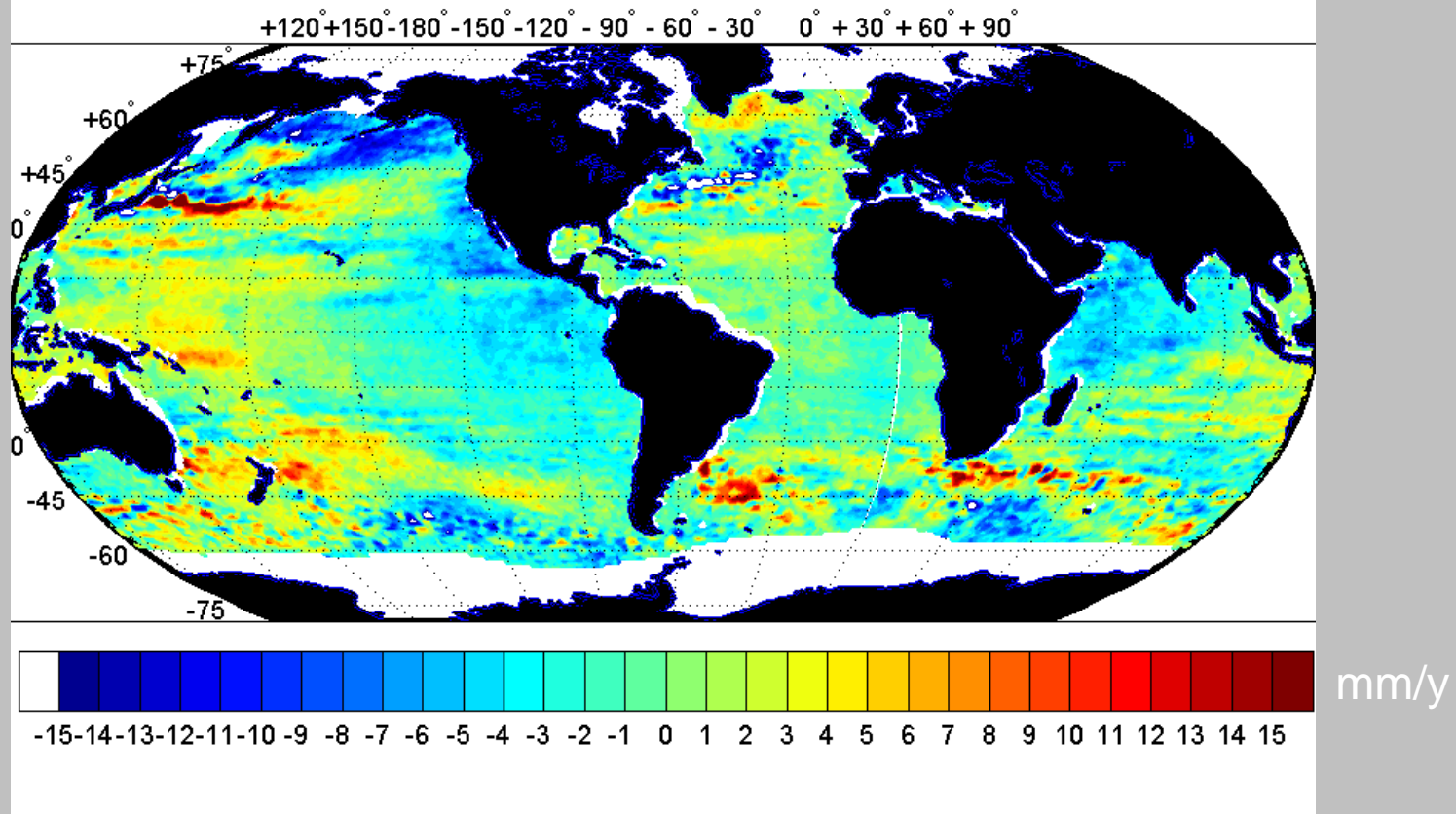
How large are they?

JPL

MEASUREMENT EFFECTS SCHEMATIC



d, S. Nerem UC , IB crctd, min: -30.655 max: 64.9542.8 removed 16-May-2006 21:32:38 CV



Directly measured by a satellite. Note how complicated the pattern is. The global mean value is estimated as about 2.8mm/y +/-0.3mm/y

According to Peltier (1991) should add another 0.33mm/y for post-glacial rebound

Seek to use the ECCO-GODAE synthesis (MIT/AER) to estimate global patterns of sea level rise, and partition it amongst heating/cooling, evaporation/precipitation/runoff, and general circulation shifts as a function of depth.

As much data as can gather and understand in the interval 1992-2004 and a reasonably complete GCM.

There exist serious modelling questions in this context:

Boussinesq approximation---conserves volume, not mass; conserves salinity, not salt.

Surface boundary conditions for salt---at least 3 in use, including virtual salt flux (Huang, 1993).

Approximations in the equation of state.

Errors in the meteorological forcing including large-scale imbalances.

Incomplete sea ice models.

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Not the present focus and here will ignore for the most part.

Meteorological Variables	Number
NCEP/NCAR (6-hrly. windstress, buoyancy flux, short/long radiation)	2.1×10^8
Oceanographic Variables	
Altimetry (TOPEX, Jason, GFO, ERS-1/2, ENVISAT)	3.3×10^7
XBT	1.0×10^7
Argo profile temperature and salinity	2.1×10^7
CTD temperature and salinity	2×10^6
Hydrographic climatologies	1.6×10^7
Sea surface temperature	5.2×10^6
TMI temperatures	1.5×10^6
GRACE geoid	57,600
Bottom topography	57,600
QSCAT winds	1.0×10^7
Approximate Number Oceanographic Observations	1.1×10^8
Approximate Number Total Weighted Values	2.11×10^8

We are solving the least-squares problem of minimizing model/data misfit, while keeping the model rigorously correct.

$$J' = [\mathbf{x}(0) - \mathbf{x}_0]^T \mathbf{P}(0)^{-1} [\mathbf{x}(0) - \mathbf{x}_0]$$

$$+ \sum_{t=1}^{t_f} [\mathbf{E}(t)\mathbf{x}(t) - \mathbf{y}(t)]^T \mathbf{R}(t)^{-1} [\mathbf{E}(t)\mathbf{x}(t) - \mathbf{y}(t)]$$

$$+ \sum_{t=0}^{t_f-1} \mathbf{u}(t)^T \mathbf{Q}(t)^{-1} \mathbf{u}(t)$$

initial condition
controls

model/data misfit

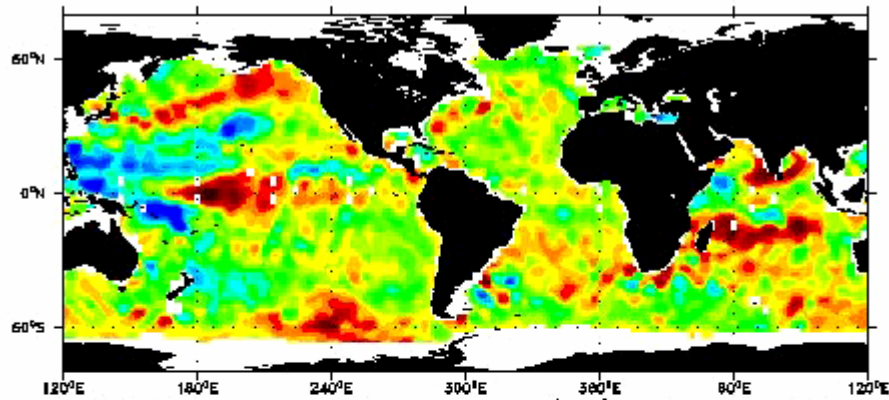
control parameters

Here the control parameters are primarily adjustments made to the surface meteorological forcing (wind stress, buoyancy exchange) and the initial conditions.

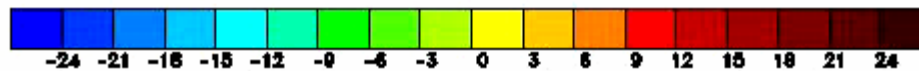
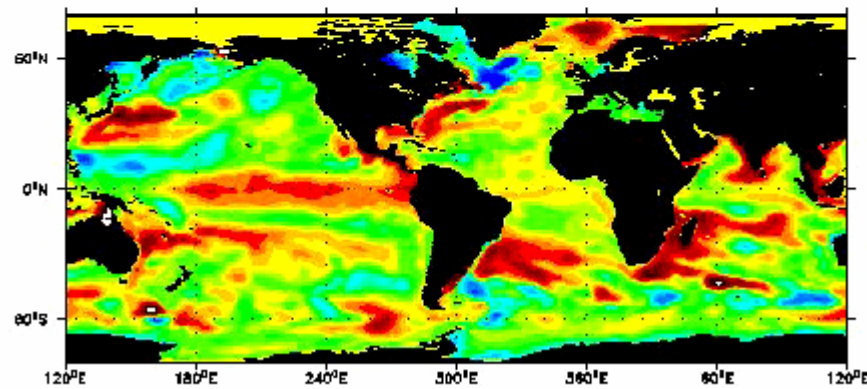
It's a problem of *constrained least-squares* and is using the model to *interpolate* the data --- a much easier task than forecasting.

But, results are obtained from an unconstrained, forward, calculation using the adjusted parameters. That is no non-physical processes are present.

TOPEX POSEIDON SSH anom 9yr. (cm) Dec 31, 1992

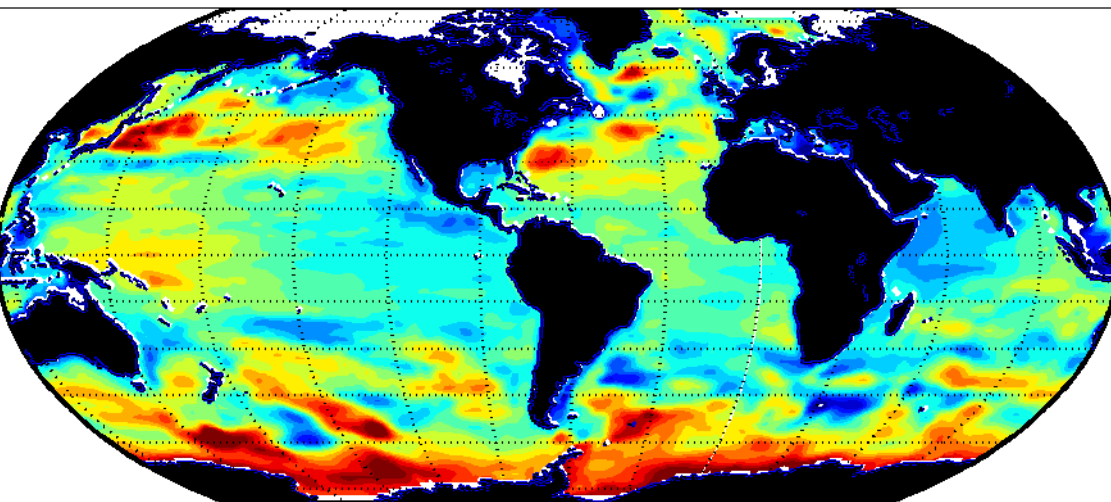


ECCO-GODAE 12yr SSH anom. (cm) Dec 31, 1992

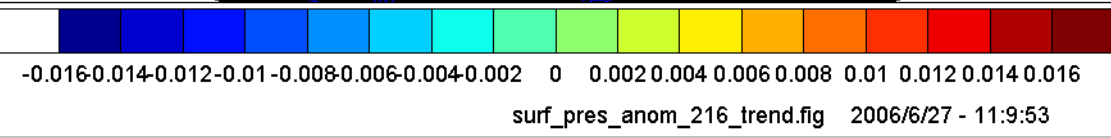


cms

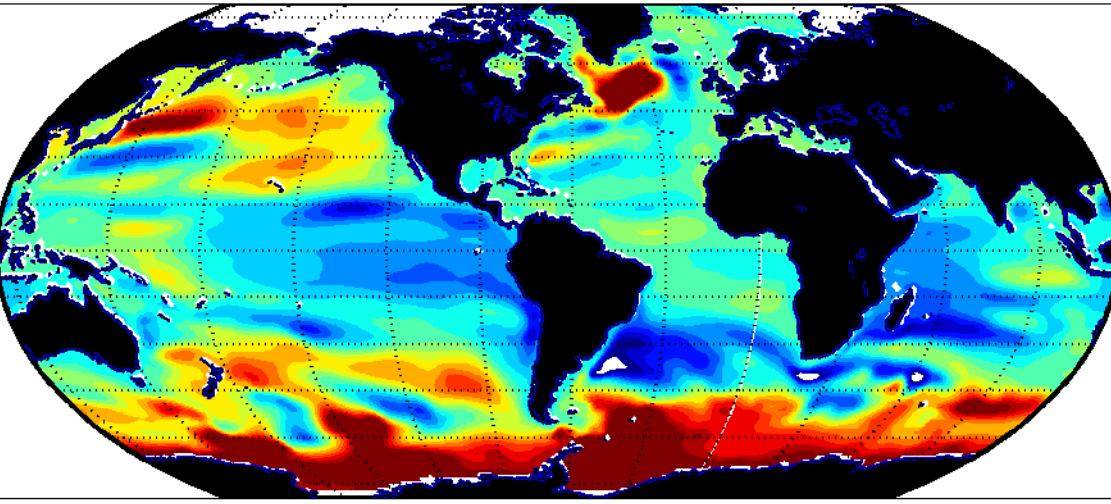
surface pres anom yearly trend (m/year) min: -0.084194 max: 0.027809 awm: 0.00044769



ECCO-GODAE estimate
v2.216 sea level trend in mm/y



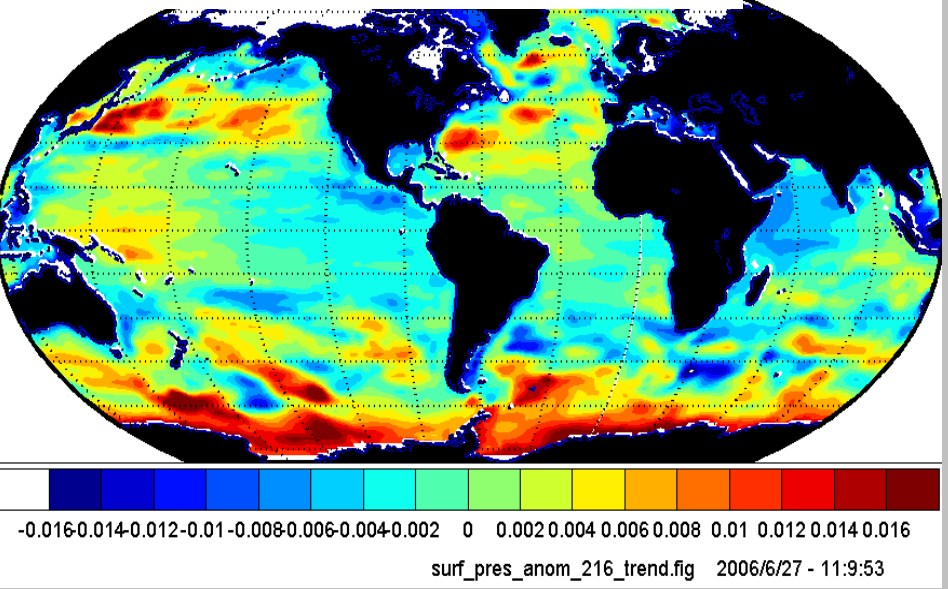
ES balanceNO iter0 yearly trend(m/year) min: -0.019225 max: 0.043917 awm: -8.9387e-09



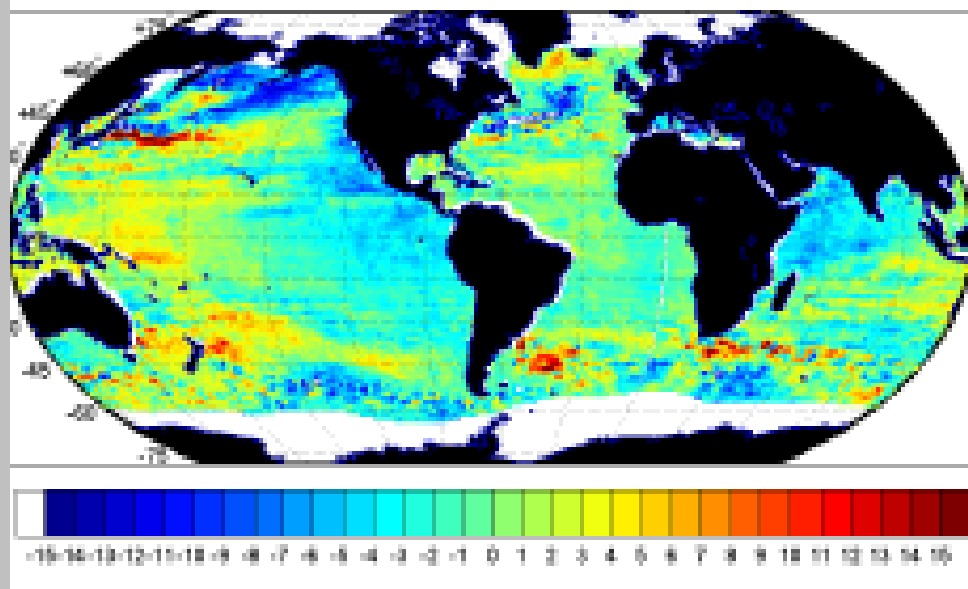
sea level trend in mm/y from
unconstrained model



surf_a Trend in sl from model and all data m/y awm: 0.00044765

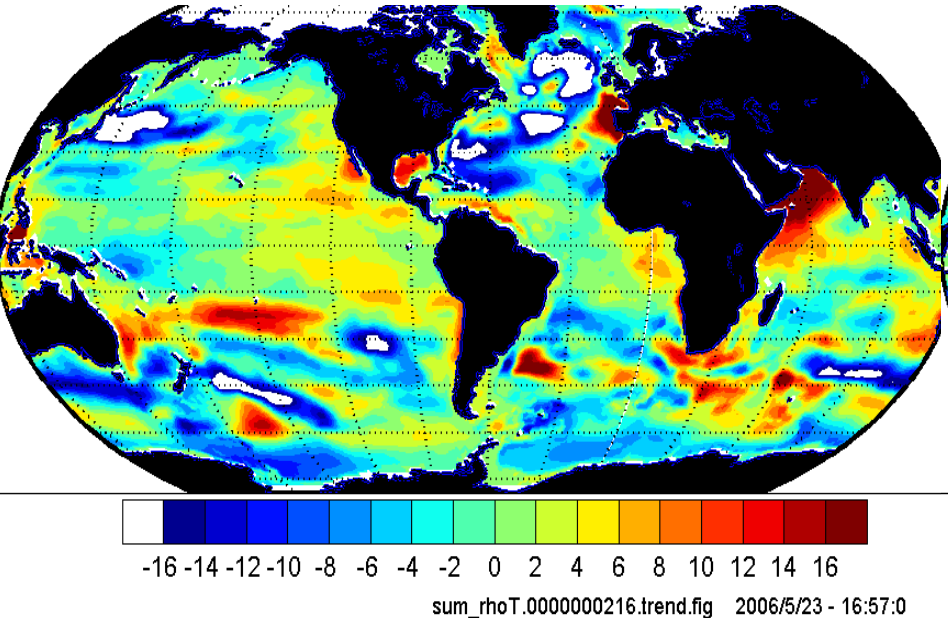


Trend in sl from altimetry data mm/y

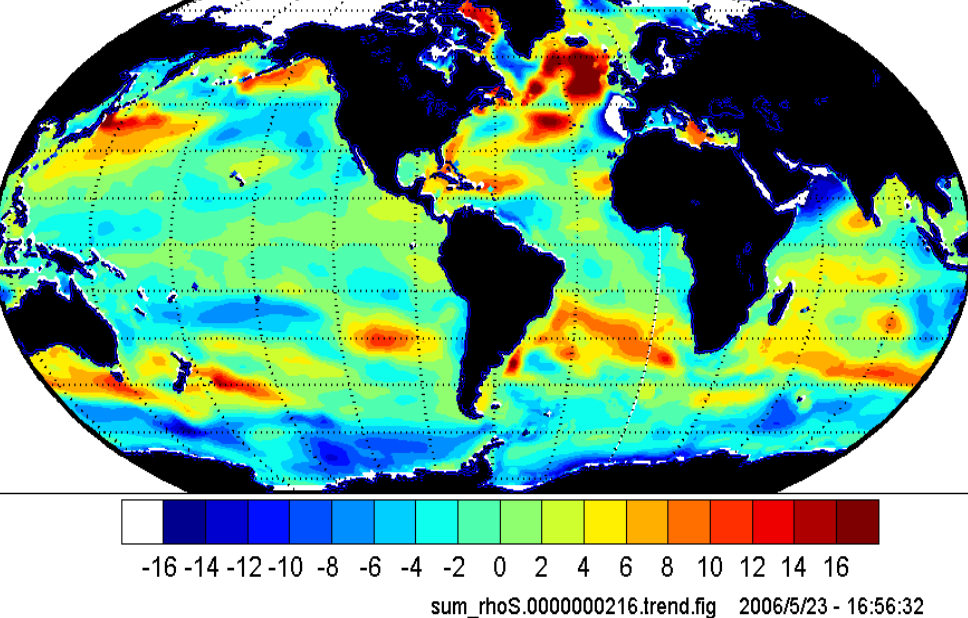


Global mean removed from the sl results

Trend in column integrated density from temperature

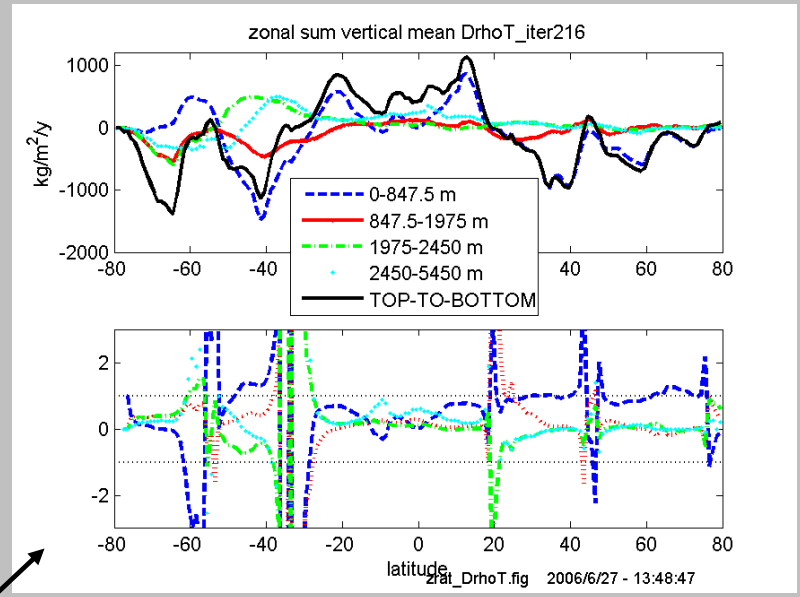
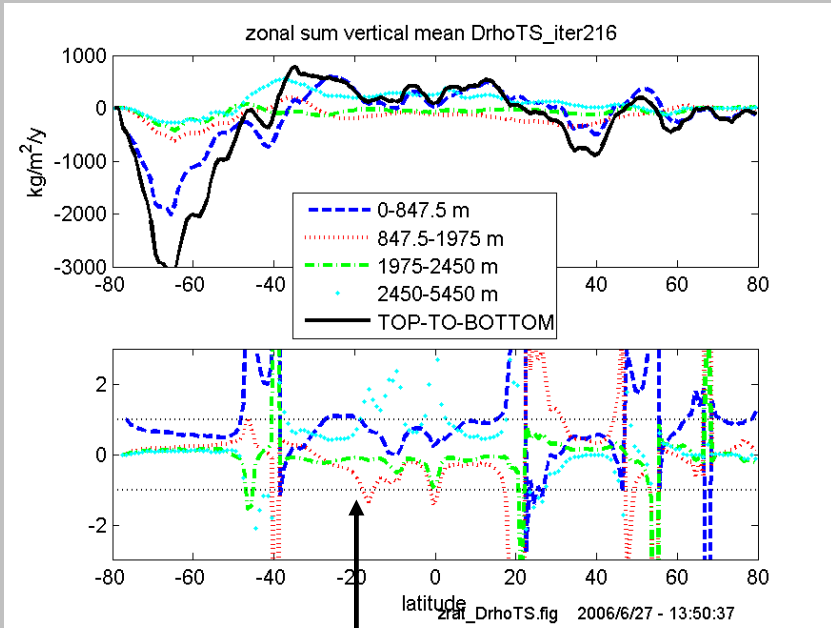


Trend in column integrated density from salt



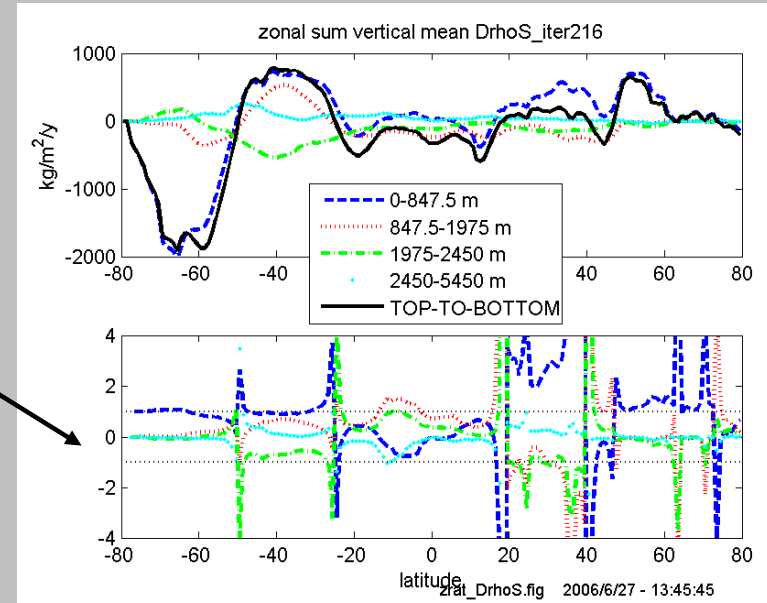
Vertical integrals of net density change from temperature

Vertical integrals of net density change

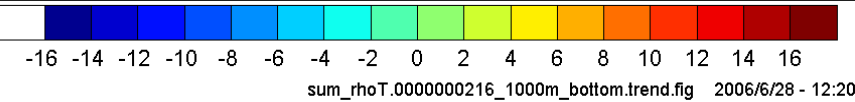
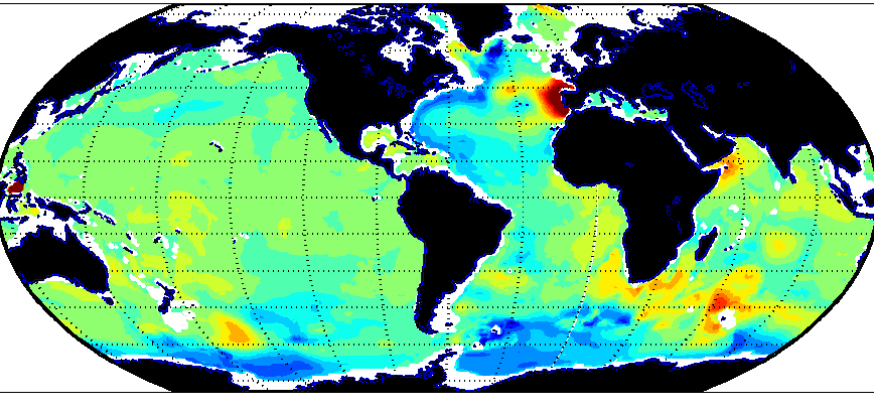


Vertical integrals of net density change from salinity

ratios to the top-to-bottom integral

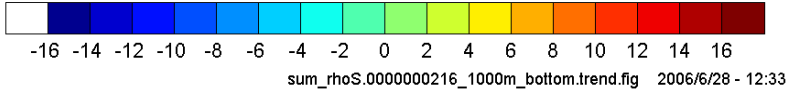
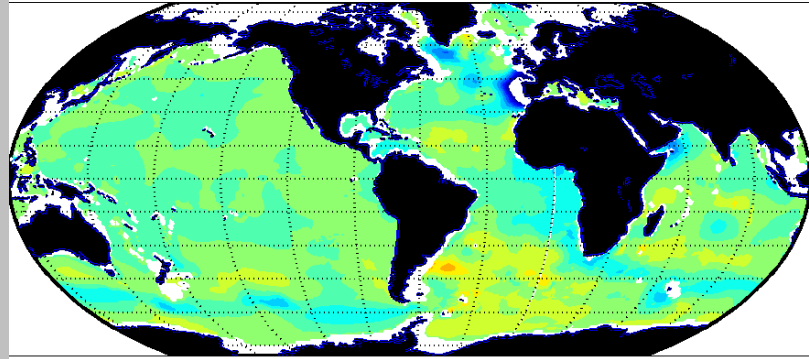


iter216 rhoT trend 1000m to bottom min: -16.8817 max: 28.8088 awm: 0.41877 rmvd



1000m to bottom temperature
contribution to density trend
mm/y

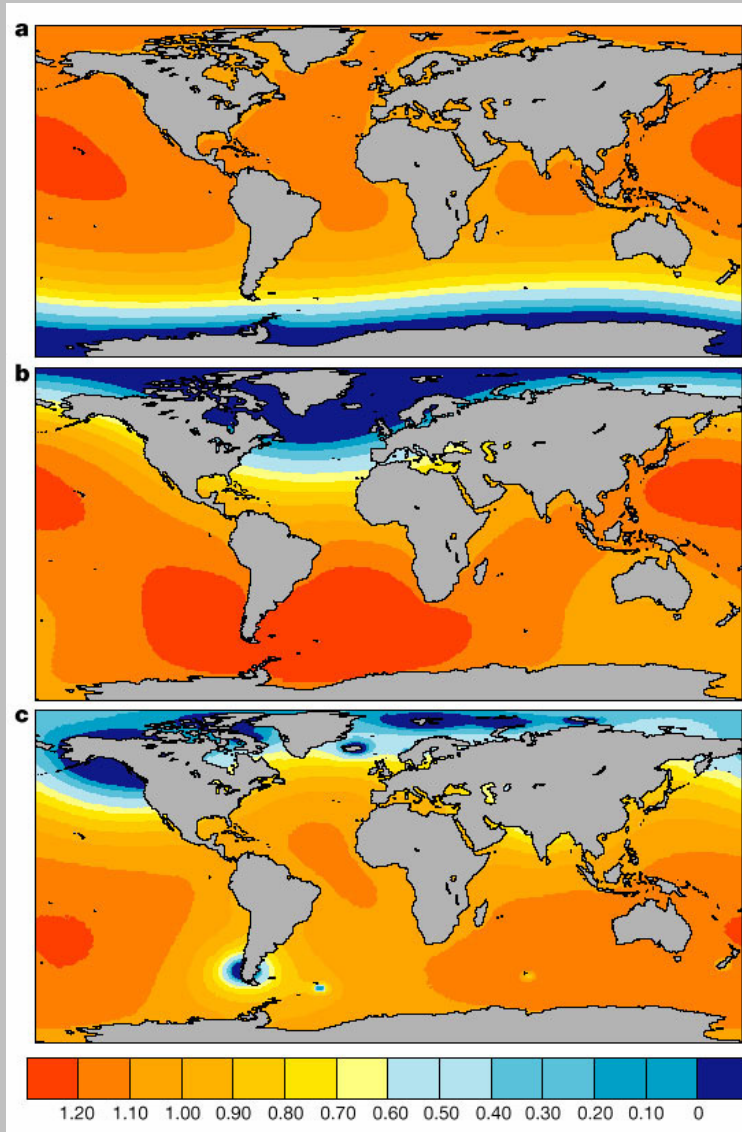
iter216 rhoS trend 1000m to bottom min: -28.5913 max: 16.7602 awm: -0.7602 rmvd



1000m to bottom salinity
contribution to density trend
mm/y

There are many more complications, e.g.:

Fractional variation of sea level from melting ice, assumed uniform over restricted areas.



← Antarctica

← Greenland

← Mountain glaciers

Mitrovica et al., Nature, 2001

A Summary

On decadal time scales, large scale patterns of sea level change are dominated by adiabatic shifts in the general circulation both lateral and vertical.

Global averages remain problematic both in general circulation models, in the data, and in the combination so that partitioning amongst heating/cooling, addition/removal of freshwater are not robust.

As time goes on, a global average signal may emerge, but one must be wary of systematic errors in both observations and models. True global averages require true global observations and models run over decades.

Not clear whether net glacial melt is better determined from oceanic or cryospheric measurements.

Thank you.