

### 230Th<sub>x</sub> Profiling method

Assumption:  $^{230}\text{Th}$  rain rate ( $F_{230}$ ) is the integrated production of  $^{230}\text{Th}$  from  $^{234}\text{U}$  in the overlying water column.

$$^{230}\text{Th}_x = P_{230} / J_m$$

with  $J_m$  = dry sediment flux ( $\text{g cm}^{-2} \text{ yr}^{-1}$ )  
and  $P$  = integrated production from  $^{234}\text{U}$  ( $\text{dpm cm}^{-2} \text{ yr}^{-1}$ )  
with  $P = \beta z$   
with  $\beta$  = production of  $^{230}\text{Th}$  from  $^{234}\text{U}$  per volume per time  
( $\text{dpm cm}^{-3} \text{ yr}^{-1}$ ),  $\beta_{230}$  ( $\lambda_{230} A_{234}$ ) is  $0.026 \text{ dpm m}^{-3} \text{ yr}^{-1}$   
and  $z$  = water depth above sediment (m)

#### Rain Rate ( $F_i$ )

$$F_i = (c_i \beta z) / ^{230}\text{Th}_{xo}$$

with  $c_i$  = concentration of component (i) in sediment.

$^{230}\text{Th}_{xo}$  is determined by correcting for decay since deposition using an independent sediment chronology.

### $^{230}\text{Th}_x$ Sediment Focusing Factor

Normalize the unsupported  $^{230}\text{Th}$  in a sediment column by the integrated production of  $^{230}\text{Th}$  from  $^{234}\text{U}$  in the overlying water column.

$$\Psi_{230} = \left( \int_{z_1}^{z_2} {}^{230}\text{Th}_{x0} \rho_b dz \right) / (P_{230} [t_1 - t_2])$$

with  $\rho_b$  = dry bulk sediment density ( $\text{kg m}^{-3}$ )

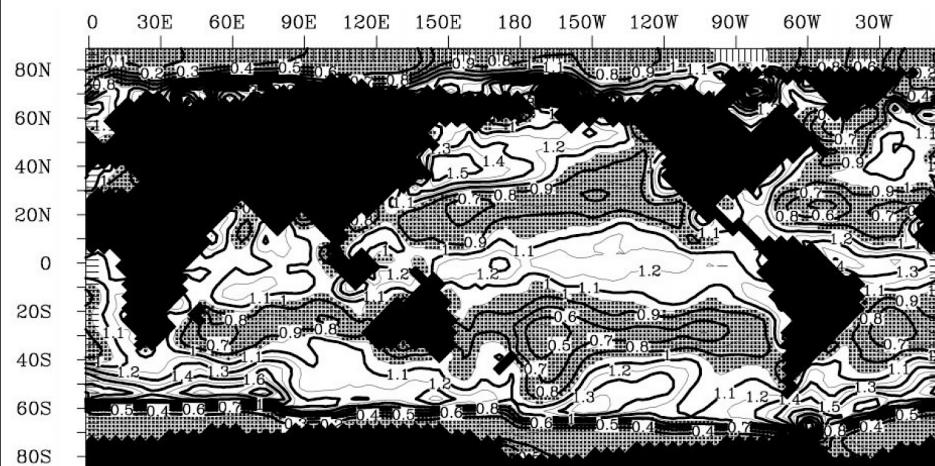
and  $t_i$  = age at depth  $z_i$  (yr)

with  $P$  =  $\beta z$

with  $\beta$  = production per volume per time ( $\text{dpm m}^{-3} \text{ yr}^{-1}$ )  
for  $^{230}\text{Th}$ ,  $\beta$  is  $2.63 \text{ dpm m}^{-3} \text{ yr}^{-1}$

and  $z$  = depth in sediment or water column (m)

Holocene  $F_{230}$  (model-derived), normalized to production in overlying water column



F = 1.0 means rain rate equals production in overlying water column

Henderson et al., 1999

